

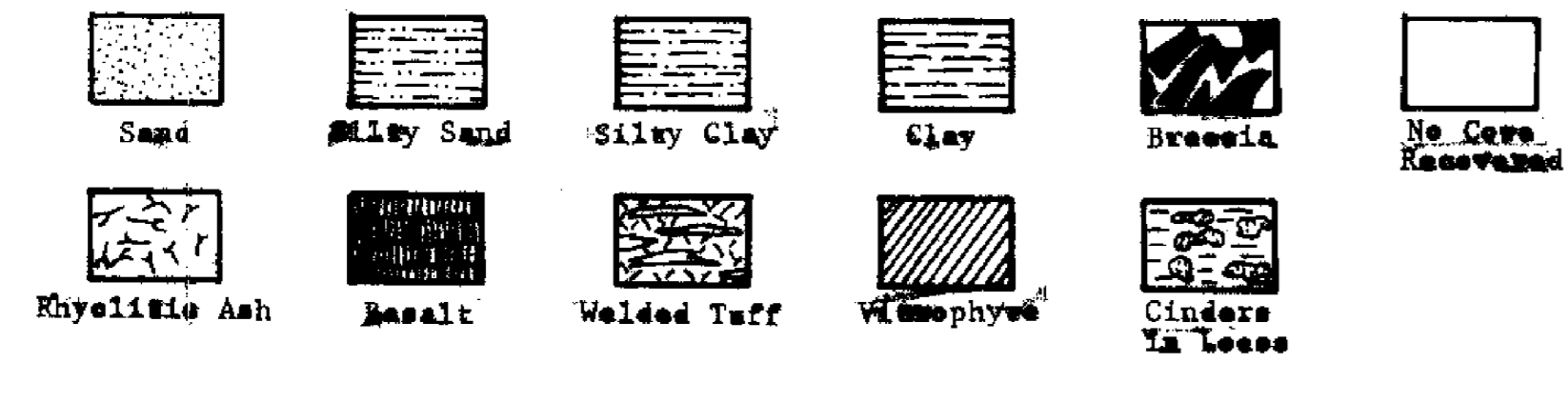
IDAHO NATIONAL ENGINEERING LABORATORY DRILL LOG

UNIVERSITY OF UTAH
RESEARCH INSTITUTE
EARTH SCIENCE LAB

GL00049-1

Duelling Dam from Exploration Well 2-2A

EXPLANATION



Structures and Textures

- Vesicular
○ Amygdaloidal
○ Lithophysal
○ Flow break (on left of column)
○ Breccia
○ No Core

Mineral and Lithologic Variations

- A Alteration (undetermined mineralogy)
B Altered basalt fragments
C Calcite
G Clay (alteration product)
G Calcite crystals in cavities
F Iron oxide staining and alteration
Op Opal
P Palagonite
S Silts and sand in fractures and openings
Wksh Wash basalt flow
S Slickensides
V Vapour phase crystallization
X Thin ash bed

Table 1.--Geophysical log information concerning drill hole 2-2A, located in the NW 1/4, sec. 15, T. 5 N., R. 31 E., Butte County, Idaho.

Table with 5 columns: Log No., Log Type, Date Logged, Test, and Interval. It lists four temperature logs and one gamma log with their respective dates and intervals.

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Logged by Charles A. Brutt and David D. Blackwell, Institute for the Study of Earth and Man, Geothermal Laboratory, 253 Marcy Building, Southern Methodist University, Dallas, Texas 75275.

REFERENCES

List of references including Armstrong, R. L., Leeman, W. P., and Malde, R. E. (1975); Chase, C. H. (1972); Ebbree, G. F., Lovell, M. D., Doherty, D. J. (1978); Malde, R. E., Pitt, A. M., and Eaton, J. P. (1971); Nace, R. L., Fogel, P. T., Jones, J. R., and Dutschak, M. (1975); Robertson, J. B., Schoen, R., and Berraclough, J. T. (1974); Ross, C. P. (1961); Walker, E. H. (1964).

The U.S. Department of Energy, in cooperation with the U.S. Geological Survey, drilled three exploration wells on the eastern Snake River Plain, Idaho, during the summer and fall of 1978. The wells were drilled to obtain subsurface geologic information pertaining to geothermal resources in these areas of the eastern Snake River Plain which have distinctly different types of geology. Two of the wells were located on property of the Idaho National Engineering Laboratory (INEL), Idaho, with the third well being located in Sugar City, Idaho. This preliminary report describes the geology and shows the subsurface distribution, and geophysical logs of well 2-2A, located on the INEL site in the NW 1/4, sec. 15, T. 5 N., R. 31 E., Butte County, Idaho. The well was located by the U.S. Geological Survey to obtain subsurface information on 1) the character and distribution of basaltic lava flows and interbedded sediments, 2) the age of the rocks and sediments, 3) the nature and extent of alteration of the rocks, and 4) an estimate of the heat flow in the area. Drilling on well 2-2A began June 13, 1978. At a depth of 412 ft (126 m), difficulties were encountered in drilling. A new well (2A) was drilled approximately 20 ft (6 m) east of the original well site, reaching a total depth of 3,900 ft (914 m) on September 23, 1978. The logs presented here are a composite of both wells (2 and 2A). The geology of the area near the well has been briefly described by Robertson, Schoen, and Berraclough (1974), Nace and others (1975), and Malde and others (1971). Walker (1964) showed that the subsurface geology of the north central part of the eastern Snake River Plain consists of at least 3,317 ft (1,011 m) of interbedded Quaternary and Neogene basaltic lava flows and beds of silt and silty clay. Walker (1964) showed evidence that the sedimentary portion of the stratigraphic section is primarily alluvial in origin, consisting of stream and lake deposits. In well 2-2A the basaltic lava flows are similar in texture and mineral composition throughout the upper 2,600 ft (610 m). There are over 14 separate basaltic lava flows in the drill core from well 2-2A. They are generally fine- to medium-grained, porphyritic, diktytaxitic olivine basalts. Phenocrysts include plagioclase and olivine in a groundmass of plagioclase, olivine, pyroxene, magnetite, ilmenite, and glass. Below about 1,900 ft (579 m), the basalts are hydrothermally altered and mineralized. This alteration and mineralization includes: filling of vesicles by calcite, zeolites, and serpentine; deposition of these same minerals along fractures; and serpentinization of the basalt, especially near some of the fractures. Many of the fractures show serpentinized slickensides. The clay-rich sediments in the lower 1,000 ft (300 m) of the well are also altered; evidence includes an overall greenish color to the clay-rich sediments, and the presence of altered grains of biotite and basalt. At least 8 thin, rhyolitic air-fall ash beds are exposed in the drill core. The glass shards in these layers appear petrographically to be unaltered by alteration. A crystal-poor, rhyolitic, lithophysal, densely welded ash-flow tuff, consisting of two flow units, occurs at a depth of 2,521 ft (768 m) to 2,538 ft (780 m). The lower flow unit consists of a black basal vitrophyre resting on a few feet of vitroclastic air-fall ash. The vitrophyre grades up into a densely welded, lithophysal, devitrified ash-flow tuff. The upper flow unit consists of about 2 ft (0.6 m) of devitrified densely welded ash-flow tuff. This ash-flow tuff is similar in texture and mineral content to ash-flow tuffs exposed at the southern end of the Lemhi Range, as described by Ross (1961), Chase (1972), and Malde and others (1971). Continuing studies on rocks and sediments from this well will investigate their age, geochemistry, alteration, paleontology, magnetic properties, and possible correlations.

