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SEISMIC STUDIES AT THE MT. HOOD VOLCANO, NORTHERN CASCADE RANGE, OREGON by

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This report is preliminary and has not been edited or reviewed for conformity with Geological Survey standards and nomenclature.

## ABSTRACT

A sixteen station telemetered seismic network was established in the Mt. Hood, Oregon area to monitor local seismicity and to study crustal and upper mantle structure. The network was in operation 13 months, and recorded 10 local earthquakes, 25 regional events, and 300 teleseisms. A series of construction blasts were recorded and used to define an average upper crustal velocity of $5.4 \mathrm{~km} / \mathrm{s}$ in the region. All local earthquakes occurred beneath Mt. Hood at shallow depths and roughly define a zone striking north-northwest beneath the mountain. The largest earthquake was a magnitude 3.4 event which had a strike-slip focal mechanism. The other events had magnitudes. (M, less than 2.0. P-wave travel time residuals from teleseismic events show a 0.5 second decrease in travel time from east to west across the Cascade Range. No travel time anomalies are associated directly with Mt. Hood.

## INTRODUCTION

Mt. Hood is the northernmost volcano in the Oregon High Cascades, the long north-south zone of Quaternary volcanism in western Oregon (Figure 1). Although the Cascades are a major tectonic province in the western United States, relatively few geophysical studies have been conducted within the range. Detailed geologic mapping of Mt. Hood (Wise, 1968) has shown that the volcano is a composite cone composed primarily of Quaternary pyroxene andesite overlying Tertiary basalt and andesite. A short-term seismic monitoring experiment (Westhusing and others, 1973) indicated some seismicity beneath the mountain.

The presence of Quaternary volcanism in the Cascade Range has raised the possibility that some geothermal potential may exist within the province. The U.S. Department of Energy, in cooperation with the State of Oregon, selected Mt. Hood as the site of a detailed, multi-disciplinary study to determine the geothermal potential of a typical Cascade volcano. Central to this evaluation was a major expansion of the existing geophysical and geological data bases at the volcano. Accordingly, a wide variety of studies were initiated at Mt. Hood during the summer of 1977. The Mt. Hood seismic network was designed for the dual purpose of monitoring local seismicity and providing data from distant earthquakes for investigating the crust and upper mantle structure beneath the mountain.

This report describes the Mt. Hood seismic network and outlines the analysis procedures used in the data reduction. We discuss the locations and focal mechanisms for the ton local earthquakes recorded by the network, and
give preliminary results of the crustal structure studies. A complete list of the local earthquakes, blasts, regional earthquakes, and teleseisms used in this work is included in tables at the end of the report.

## SEISMIC NETWORK AND DATA ANALYSIS

The Mt. Hood seismic network consisted of 16 stations deployed over an area 50 km (north-south) by 40 km (east-west) centered on Mt. Hood (Figure 2). Four stations were located on the lower flanks of the volcano, and the average spacing between stations was about 10 km . The stations began operating in late November 1977 and recorded continuously until December 1978 when 13 stations were removed. Station data is summarized in Table 1.

The network followed the standard pattern of U.S.G.S. telemetered seismographic networks (McHugh and Lester, 1978). Vertical seismometers with a natural period of 1 Hz were used at each site, and the data were telemetered from the remote sites to central receiver sites by radio. Signals were multiplexed at the receiver points and transmitted over phone lines to USGS headquarters in Menio Park, California, where the data were recorded in analog format on 16 mm film and on 2.54 cm magnetic tape. The film records were scanned daily and local, regional, and teleseismic events were cataloged.

Events were selected from the film scanning catalog for more detailed analysis. Selected events were digitized from the analog magnetic tapes using a Data General ECLIPSE $S / 200$ mini-computer, and the seismic traces were displayed on a graphics terminal for the purpose of timing the seismic phases. For local earthquakes, P-wave arrivals, the direction of the first motion, duration of the event coda, and the S-wave arrivals (when they appeared
to be clear and unambiguous) were read. Only the first arrivals were read for regional events, in most cases this corresponds to Pn. Teleseismic arrivals were timed at the first break for sharp arrivals. For emergent arrivals either the first peak or first trough was timed.

DATA AND RESULTS

## Explosions

Explosion data recorded from three construction sites (Salmon River S-curve, Trillium Lake and Wildwood) located within the Mt. Hood network (Figure 3) were used to constrain the seismic velocity model used in the earthquake location routine. Twenty-two explosions were recorded from the three blast sites, with at least two blasts recorded from each site (Table 2) 2). Origin times were recorded for the Salmon River S-curve site and the Trillium Lake blasts. For the Salmon River Shot, three temporary recorders were installed to provide additional arrivals in the first ten kilameter section of the travel time curve (Figure 3 ).

Reduced-travel time curves from the Salmon River and Trillium Lake data (Figure 4 ) indicate a thin near-surface layer with a velocity of $3.7 \mathrm{~km} / \mathrm{sec}$. overlying a deeper layer with a velocity ranging from 5.4 to $5.7 \mathrm{~km} / \mathrm{sec}$. The depth to this layer is about 1 km , although neither the velocity nor depth is well constrained (Figure 4a). Origin times are lacking for the Wildwood shots, but the slope of the reduced travel time plot (Figure 4b) is similiar to the slope determined from the other two shot points. In the composite plot (Figure 4 c ), the Wildwood origin time has been arbitrarily adjusted so that all the data can be plotted together. A simple half space crustal model, with
a velocity of $5.4 \mathrm{~km} / \mathrm{sec}$, was selected from examination of the travel time curves (Figures 4a-c). This velocity is an approximation of the actual structure in the Mt. Hood region; and from the degree of scatter in the arrival times it is apparent that the upper crust has a complex structure.

The USGS conducted a comprehensive seismic refraction experiment in the Mt. Hood region during the fall of 1978. Shot arrivals from the 4 shot points near Mt. Hood were recorded by the seismic network and were used to determine the accuracy of the simple half-space velocity model obtained from the road blasts and subsequently utilized in locating earthquakes. Treating the shot arrival times at the network as input, the shots were located using the HYPOELLIPSE computer program (John Lahr, personal communication). Comparison of the results with the actual shot locations (Figure 5) shows that the half-space model located all shots within $\pm 2 \mathrm{~km}$ of the actual locations.

The travel-time curves for the seismic refraction shots are plotted in Figure 6a-d. As with the construction blasts, an upper-crustal velocity in the range $5.4-5.6 \mathrm{~km} / \mathrm{sec}$ is indicated. A sumary of the blast data is given in Table 2.

## Local Earthquakes

A total of 10 local earthquakes were recorded by the Mt. Hood seismic network during the 13 months of operation. These earthquakes were located assuming the $5.4 \mathrm{~km} / \mathrm{sec}$ half-space model using the HYPOELLIPSE computer program (John Lahr; personal communication). The resulting earthquake epicenters are shown in Figure 7 and a list of the location parameters is given in Table 3. Event magnitudes were calculated using the relationship developed for Central California (Lee and Lahr, 1975).

The largest event ( $M_{L}=3.4$ ) occurred on September 6, 1978, beneath the western flank of the volcano (Figure 7). A single smaller event preceeded the main shock by 4 hours, and three locatable aftershocks followed within twenty minutes of the main event. All events have essentially the same epicentral position, and all are at shallow depths (less than 10 km ). The other 5 events were small ( $M_{L}<2.0$ ), but like the sequence on September 6 , they were all spatially related to Mt. Hood and occurred at shallow depths (less than 16 km ) beneath the volcano.

Fault plane solutions were determined for 5 of the earthquakes (Figure 8). Since the events are all in the center of the network, the solutions have small azimuthal gaps and are well constrained. For the September 6 mainshock, the focal machanism is nearly pure strike-slip, along either NNE or WNW striking fault planes (Figure 9). The orientation of the T-axis is NNW, with the P-axis being ENE. No consistent trend is seen in the orientation of the $P$ and $T$ axes for the other events, as several different focal mechanisms are evident (Figure 10). There is a suggestion (Figure 8) that the fault planes are striking to the northwest beneath Mt. Hood.

## Regional Earthquakes

Arrival time data from regional earthquakes can be used to model the structure of the lower crust and upper mantle in the region. Fourteen well-recorded regional events, from sources as distant as 700 km were used in this study, (Table 4). Locations for earthquakes were determined from the PDE tables published by the National Earthquake Information Service. The apparent velocity calculated for each event is shown in Figure 11. The Pn velocities
appear to be in the range of 7.5 to $8.1 \mathrm{~km} / \mathrm{sec}$. The unusually low Pn velocity from the north compared to other azimuths, though not well constrained by more data, may reflect the lower than normal Pn velocity observed in the Puget Sound area by Crosson (1976).

## Teleseisems

Earthquakes from teleseismic distances (200 or more, with $10=111 \mathrm{~km}$ ) arrive at the surface with angles of incidence that generally become more vertical as the distance of the earthquake increases. These arrivals are well suited for studying variations in the crustal and upper mantle structure beneath a small apperature array, since the rays have a nearly common travel path everywhere except near the receivers. Observed travel time differences can thus be assumed to be related to changes in the velocity structure under the recording array. Using this technique, volumes with anomalous velocities have been defined in several geothermal areas. Spectacular travel time delays have been observed at Yellowstone National Park Wyoming, and the Geyers-Clear Lake region California. In both cases the velocity change has been interpreted as being caused by the presence of rock in a state of partial melt (Iyer, H.M and R.M. Stewart, 1977).

A total of 55 well recorded teleseisems were used to study the regional variation of P -wave residuals at Mt. Hood. The teleseisems were primarily from the northwest, southwest, and southeast azimuths. The northwest and southeast events occur over a wide distance range of 230 to 910 , while the southwest events were from a narrow distance range between 660 and 900 . Three Russian nuclear events were recorded from north at a distance of about $8.1^{\circ}$. The events are listed by azimuth in Tables 5, 6, and 7. Locations of the teleseisms were determined fram the published PDE listings.

Travel time residuals were determined for each event by subtracting a theoretical travel time (TT) computed using Herrin's standard earth model (Herrin, 1968) from the observed travel time (TO) between the source and the recording station. The travel time residual ( $R$ ) at the ith station is:

$$
R_{i}=T O_{i}=T T i
$$

To correct for common source and path effects, thus isolating the contribution to the residual from local velocity variations, relative residuals were computed by selecting a single station as a reference station and subtracting the residual at the reference station (Rref ) from the residual ( $\mathrm{R}_{\mathrm{i}}$ ) at all other stations:

$$
R R i=R i-R r e f \quad .
$$

The relative residual at the reference station is zero by definition.
At Mt. Hood, station VFL (Figure 2) was used as the reference station. The relative residuals for events from the southeast, southwest, and northwest azimuths at each station are listed in Tables 8, 9, and 10 respectively. The average residuals from each quadrant at each station are plotted and contoured in Figures 12a-c.

The major feature in the relative residual pattern is the sharp gradient in the east-west direction. The relative residuals (Figure 12a-c) indicate that arrivals in the western Cascade Range are faster than arrivals to the east by almost 0.5 seconds. The gradient of residuals is nearly perpendicular to the axis of the Cascade Range. No travel time anomaly can be associated with Mt. Hood.

In Figure 13, the relative residuals have been plotted by distance and azimuth after making a simple elevation correction using a velocity of 6.0 $\mathrm{km} / \mathrm{sec}$. These results show that the east-west gradient defined by the contour plots (Figure l2a-c) is not a function of the epicentral distance (the east-west pattern is seen with both the closer and the more distant events). Using a $6.0 \mathrm{~km} / \mathrm{sec}$ reducing velocity instead of the $5.4 \mathrm{~km} / \mathrm{sec}$ velocity indicated for the upper crust by the blast data may incur a .02 sec difference in travel times.

Velocity Structure of the Mt. Hood Volcano

The blast data and the teleseismic data enable some inferences to be made on the crust and upper mantle structure in the Mt. Hood region. The crustal data indicates an upper crustal layer with a velocity in the range of 5.4 to $5.7 \mathrm{~km} / \mathrm{sec}$ overlain by a thin lower velocity layer. Our velocity structure agrees in general with the preliminary findings of the detailed seismic refraction experiment conducted in the region by the U.S.G.S. (Kohler and others, 1978). We see no evidence in the upper crust of the presence of any velocity anomalies associated with the Mt. Hood volcano. The teleseismic P-wave residuals show a 0.5 second decrease in travel time from east to west across the Cascade Range which may reflect variations in the crustal and mantle thickness and/or velocities. The blast data when combined with the teleseismic results can be used to argue against the presence of a large (5-10km in diameter), shallow (10km or less) low velocity body beneath Mt. Hood.

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Table 1: Station co-ordinates and elevation of Mt. Hood seismic network

| Station name | Latitude | Longitude | Elevation (m) |
| :---: | :---: | :---: | :---: |
|  | deg. min. | deg. min. |  |
| VBB Bald Butte | 45 N 32.25 | 121 W 31.83 | 1122 |
| VBR Bull Run | $45 \times 30.36$ | 121 W 55.55 | 0635 |
| VPP Post Point | 45 N 11.94 | 121 W 30.00 | 1128 |
| VFL Frog Lake | 45 N 13.04 | 121 W 39.92 | 16.13 |
| VWC Wolf Camp | $45 \times 14.51$ | 121 W 47.82 | 1457 |
| VHR High Rock | $45 \times 09.81$ | 121 W 53.62 | 1414 |
| VIS Indian Springs | 45 N 34.48 | 121 W 49.35 | 1305 |
| VLL Lawrence Lake | 45 N 27.80 | 121 W 40.75 | 1195 |
| VBY Bonney Butte | 45 N 15.79 | 121 W. 35.52 | 1707 |
| VTJ Tilly Jane | 45 N 24.95 | 121 W 38.82 | 1753 |
| VCR Cathedral Ridge | 45 N 25.58 | 121 W 46.96 | 1097 |
| VGC Government Camp | 45 N 18.69 | 121 W 45.08 | 1256 |
| VSR Salmon River | 45 N 18.73 | 121 W. 56.54 | 0500 |
| VOM Old Maid Flat | 45, N 22.98 | 121 W. 53.10 | 0585 |
| VHM Mt. Hood Meadows | $45 \times 19.75$ | 121 W 39.87 | 1647 |
| VFP Flag Point | 45 N 19.09 | 121 W 27.91 | 1722 |

Table 2: Blast data with locations and origin times if recorded.

| Location | Date | Latitude | Longitude | Origin Time (G.M.T.) |
| :---: | :---: | :---: | :---: | :---: |
|  |  | deg. min. | deg. min. |  |
| Salmon River | 4/10/78 | $45 \times 13.97$ | 121 W 41.99 |  |
| S-curve | 4/12/78 | $45 \times 13.97$ | 121 W 41.99 |  |
|  | 4/20/78 | $45 \times 13.97$ | 121 W 41.99 |  |
|  | 4/24/78 | $45 \times 13.97$ | 121 W 41.99 |  |
|  | 4/26/78 | $45 \times 13.97$. | 121 W. 41.99 |  |
|  | 4/27/78 | $45 \times 13.97$ | 121 W:41.99 |  |
|  | 5103/78 | $45 \times 13.97$ | 121 W. 41.99 . | 20:00:17.99 |
| - . | 5109/78 | 45 N 14.04 | 121 W 41.51 |  |
|  | 5/11/78 | 45 N 14.46 | 12.1. W 41.71 |  |
|  | 5/15/78 | 45 N: 14.97 | 121 W 42.22 |  |
|  | 6/05/78 | $45 \times 14.93$ | 121 W 42.21 |  |
|  | 6/08/78 | $45 \times 14.69$. | 121 W 42.06 |  |
| Trillium Lake | 6/09.78 | $45 \times 15.14$ | 121 W 44.84 |  |
| Quarry | 6/13/78 | $45 \times 15.26$ | 121 W 44.87 |  |
|  | 6/28/78 | 45 N 15.14 | 121 W. 44.84 |  |
|  | 7/12/78 | 45 N 15.14 | 121 W: 44.84 |  |
|  | 8/02/78 | 45 N 15.14 | 121 W. 44.84 |  |
|  | 8/15/78 | $45 \times 15.14$ | 121 W 44.84: | 04:07:04.80 |

## Table 2 (cont.)

| Location | Date | Latitude | Longitude |
| :---: | :---: | :---: | :---: |
|  |  | deg. min. | deg. min. |
| Wildwood | 3/16/78 | $45 \times 22.50$ | 122 W 00.50 |
|  | 3/17/78 | 45 N 22.50 | 122 W 00.50 |
|  | 4/21/78 | 45 N 21.56 | 121 W 59.67 |
|  | 6/03/78 | 45 N 22.50 | 122 W 00.50 |

U.S.G.S. Seismic Refraction Experiment

10/19/78
45 N 31.16
121 W 44.32
08:00:00.01
45. N 21.56

121 W 34.39
08:15:00.01
45 N 10.84
121 W 43.60
08:30:00.01
45. N 18.25

122 W. 04.94
11:00:00.01
10/23/78
45 N 31.16
121 W 44.32
08:00:00.01
45 N 21.56
121 W 34.39
08:15:00.01
45.N 10.84

121 W 43.60
08:30:00.01
45 N. 18.25
122 W 04.94
11:00:00.01

Table 3: Local earthquakes with solutions computed by the hypoellipse computer program

| Date | Origin Time (G.M.T.) | Latitude deg. min. | Longitude deg min. | Depth (km below <br> sea level | $\begin{aligned} & \text { Magnitude } \\ & \text { 1) } \quad\left(M_{L}\right) \end{aligned}$ | RMS | GAP |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 11/18/77 | 00:01:43.33 | 45N23.15 | 8121W41.86 | 0.40 | $<2.0$ | . 14 | 261 |
| 11/23/77 | 23:48:22.80 | 45N19.18 | 121W40.28 | 6.59 | <2.0 | . 19 | 64 |
| 03/30/78 | 13:35:58.92 | 45N20.37 | 121W41.42 | 7.73 | $<2.0$ | . 14 | 68 |
| 03/30/78 | 13:36:26.56 | 45N19.38 | 121W41.87 | 11.98 | <2.0 | . 12 | 87 |
| 07/14/78 | 17:25:22.31 | 45N23.24 | 121W41.74 | 15.12 | <2.0 | . 46 | 63 |
| 09/06/78 | 00:07:44.84 | 45N22.32 | 121W42.21 | 0.91 | <2.0 | . 12 | 87 |
| 09/06/78 | 04:47:54.20 | 45N22.26 | 121W42.53 | 2.08 | 3.4 | . 13 | 63 |
| 09/06/78 | 04:49:09.03 | 45N22.45 | 121W42.11 | 9.70 | <2.0 | . 28 | 99 |
| 09\%06/78 | 04:49:43.79 | 45N22.278 | 121W4 1.60 | 4.38 | $<2.0$ | . 16 | 118 |
| 09/06/78 | 04:50:40.97 | 45N21.86 | 121W42.17 | 2.01 | <2.0 | . 05 | 170 |

Table 4: Dates of regional events with locations, depth, and magnitude.

| Location | Date | Time (G.M.T) | Latitude deg. min. | Longitude deg. min. | Depth(km) Mag |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Mt. St. Helena, WA | 04/08/78 | 07:33:13.97 | 46 N 07.72 | 122 W 10.57 | 23.40 |
| Oregon | 11/01/78 | 09:14:59.97 | $44 \times 21.96$ | 122 W 07.07 | 0.00 blast |
| Arlington, Oregon | 02/20/78 | 21:23:16.10 | 45 N 53.02 | 119 W 42.13 | 0.00 blast |
| Bremerton, WA | 03/31/78 | 08:02:58.40 | 47 N 18.60 | 122 W 24.60 | 10.004 .2 |
| Bremerton, WA | 03/11/78 | 15:52:11.37 | $47 \times 24.67$ | 122 W 42.09 | 25.00.4.2 |
| Montana | 04/23/78 | 23:24:39.10 | 46 N 58.20 | 113 W 28.20 | 33.004 .9 |
| Oregon | 11/01/78 | 09: 30:00.01 | 42 N 50.89 | 122 W 30.57 | 0.00 blast |
| Mt. Shasta, CA | 08/01/78 | 09:02:34.50 | 41 N 26.40 | 121 W 46.20 | 5.004 .2 |
| Idaho | 11/27/78 | 09: $25: 55.10$ | 44 N 34.92 | 116 W. 16.38 | 5.004 .2 |
| Oregon | 04/23/78 | 20:10:04.80 | 44 N 34.20 | 127 W 13.20 | 33.004 .9 |
| Oregon | 02/16/78 | 12:00:20.30 | $42 \times 40.80$ | 126 W 07.80 | 33.004 .6 |
| Mt. Shasta, CA | 08/01/78 | 09: $46: 49.20$ | $41 \mathrm{~N}=07.20$ | 122 W 01.20 | 2.004 .2 |
| Idaho | 03/22/78 | 14:31:20.00 | $44 \times 13.80$ | 115 W 32.40 | 18.003 .7 |
| Vancouver Island | 06/02/78 | 20:41:45.40 | 50 N 19.20 | 127 W 30.60 | 15.005 .5 |

Table 5: Location and times of telesesisms from 900 to 1800 azimuths

| Location | Date | $\begin{gathered} \text { (GMT) } \\ \text { Origin Time } \end{gathered}$ | Latitude deg. min. | Longitude deg min | Depth (km) |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Panama | 04/04/78 | 21:11:41.0 | 9N 52.2 | 78W 6.0 | 33.0 |
| Peru | 07/28/78 | 14:01:29.2 | 10S 21.6 | $77 \mathrm{~W} \quad 3.0$ | 614.0 |
| Chile | 11/24/77 | 17:57:06.0 | 19S 48.0 | 68W 18.0 | 150.0 |
| Chile | 02/05/78 | 01:52:26.2 | 17500.0 | 69W 24.0 | 150.0 |
| Chile | 02/11/78 | 14:43.39.5 | 20S 54.0 | 66 W 00.0 | 100.0 |
| Chile | 03/08/78 | 18:11:15.5 | 26S 38.4 | 70W 37.8 | 51.0 |
| Guatemala | 02/22/78 | 06:07:34.1 | 14 N 12.0 | 91 W 18.0 | 73.0 |
| Guatemala | 02/22/78 | 18:31:46.0 | 14N 18.0 | 91W 12.0 | 100.0 |
| Guatemala | 02/18/78 | 19:43:09.6 | 14N 36.0 | 90W 48.0 | 282.0 |
| Guatemala | 02/22/78 | 06:07:34.1 | 14 N 12.0 | 91 W 18.0 | 73.0 |
| Argentina | 01/24/78 | 13:18:24.2 | 30S 54.0 | 68 W 6.0 | 50.0 |
| Argentina | 01/17/78 | 11:33:13.8 | 31 S 24.0 | 68 W 00.0 | 25.0 |
| Argentina | 01/28/78 | 04:42:50.3 | 34 S 18.0 | 68W 35.0 | 100.0 |
| Argentina | 11/23/77 | 09:26:22.7 | $30 S 54.0$ | 67 W 36.0 | 33.0 |
| Argentina | 11/24/77 | 02:02:36.6 | 31530.0 | $67 \mathrm{~W} \quad 3.6$ | 33.0 |
| Argentina | 10/27/78 | 10:06:45.2 | 22 S 3.54 | 65 W 51.3 | 265.0 |
| Nevàda | 07/12/78 | 17:00:00.0 | 37N 4.8 | 116W 2.4 | 1.0 |
| Mexico | 01/11/78 | 20:05:17.3 | 18 N 6.0 | 107W 12.0 | 33.0 |
| Mexico | 03/19/78 | 01:39:19.2 | 17 N 6.0 | 99W 30.0 | 85.0 |
| Mexico | 03/12/78 | 18:42:20.6. | 32N 6.0 | 115W 12.0 | 15.0 |

Table 6: Location and times of teleseisms from 1800 to 2700 azimuths

| Location | (CMT) |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | Date | Origin Time | Latitude deg. min. | Longitude deg. min. | Depth (km) |
| Pacific | 01/05/78 | 03:23:21.1 | 21542.0 | 126W 42.0 | 33.0 |
| Calif. | 11/22/77 | 21:15:51.4 | 39 N 24.0 | 123W 18.0 | 33.0 |
| South |  |  |  |  |  |
| Pacific | 11/24/77 | 16:59:54.5 | 22530.0 | 139W 18.0 | 33.0 |
| Tonga | 01/28/78 | 19:36:30.5 | 25S 54.0 | 177W 24.0 | 33.0 |
| Tonga | 01/29/78 | 21:34:00.9 | 24 S 48.0 | 175W 36.0 | 33.0 |
| Tonga | 03/01/78 | 20:33:37.6 | 15S 18.0 | 174W 00.0 | 33.0 |
| Kermadec | 11/27/77 | 12:40:04.5 | 21S 48.0 | 179E 42.0 | 550.0 |
| Samoa | 07/17/78 | 13:26:16.3 | 14543.2 | 175W 55.2 | 300.0 |
| Fiji | 01/19/78 | 10:44:03.5 | 18S 12.0 | 179E 54.0 | 500.0 |
| Fiji | 12/13/78 | 03:08:54.6 | 17 S 12.0 | 179E 18.0 | 600.0 |
| Fiji | 02/21/77 | 05:08:51.5 | 17 S 54.0 | 179E 6.0 | 600.0 |
| Fiji | 02/19/78 | 20:23:46.2 | 21S 42.0 | 179E 54.0 | 600.0 |
| Fiji | 02/19/78 | 20:23:46.2 | 21S 42.0 | 179E 54.0 | 600.0 |
| Fiji | 03/25/78 | 19:25:54.2 | 22S 28.8 | 179E 26.4 | 600.0 |
| Fiji | 11/30/77 | 10:15:48.4 | 20S 21.0 | 179E 45.0 | 600.0 |
| New <br> Hebrides | 09/19/78 | 12:45:15.0 | 19S 18.6 | 168E 43.8 | 252.0 |
| Solomon | 01/25/78 | 23:18:58.4 | 4548.0 | 151E 54.0 | 33.0 |
| Solomon | 11/22/77 | 15:56:29.4 | 11 S 30.0 | 159E 36.0 | 33.0 |
| Solomon | 04/29/78 | 04:21:08.2 | 9S 46.8 | 159E 36.0 | 33.0 |
| New |  |  |  |  |  |
| Britain | 01/25/78 | 23:18:58.4 | 4S 49.8 | 151E 51.6 | 33.0 |

Table 7: Location and times of teleseisms from 2700 to 3600 azimuths

| Location | (GMT) |  |  | Longitude deg. min. | Depth (km) |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | Date | Origin Time | Latitude deg. min. |  |  |
| Marianas | 09/26/78 | $14: 24: 37.7$ | 12 N 39.6 | 143E 9.0 | 33.0 |
| Japan | 01/11/78 | 48:10:22.3 | 31 N 18.0 | 139E 18.0 | 100.0 |
| Japan | 02/22/78 | 18:00:41.2 | 30 N 6.0 | 131E 18.0 | 60.0 |
| Japan | 02/20/78 | 04:36:57.0 | 39N 12.0 | 141 E 54.0 | 40.0 |
| Japan | 03/07/78 | 02:48:38.9 | 31 N 54.0 | 137E 30.0 | 432.0 |
| Japan | 06/12/78 | 48:14:26.8 | 38N 24.0 | 141E 57.6 | 40.0 |
| Yugoslavia | 01/24/78 | 05:54:21.7 | 45N 00.0 | 150E 00.0 | 33.0 |
| Kuril | 01/29/78 | 02:05:02.9 | 46N 18.0 | 149E 12.0 | 150.0 |
| Kuril | 12/04/77 | 11:39:08.5 | 49 N 6.0 | 146E 54.0 | 500.0 |
| Kuril | 01/10/78 | 00:50:23.5 | 47N 18.0 | 152E 30.0 | 100.0 |
| Kuril | 03/24/78 | 23:15:48.4 | 44N 47.4 | 149E 26.4 | 33.0 |
| Alaska | 11/27/77 | 15:05:02.4 | 58N 39.0 | 156W 27.0 | 125.0 |
| Alaska | 02/12/78 | 08:56:38.7 | 59N 24.0 | 152W 36.0 | 67.0 |
| Ural | 11/30/77 | 04:08:01.9 | 50N 54.0 | 78 E 12.0 | 1.0 |
| SSR | 03/26/78 | 03:57:05.1 | 51N 16.8 | 77 E 24.0 | 1.0 |
| Russia | 06/11/78 | 04:56:57.5 | 49N 54.6 | 78 E 53.4 | 1.0 |

© Table 8: Telesesimic P-delays relative to station FL. Residuals are in seconds. Events from the southeast ( $90^{\circ}-180^{\circ}$ azimuth)

| Event : | ar | CR | TJ | 4. | BR | FL | WC | FP | Is | BB | SR | HM | HR | PP | BY | GC |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Nevada |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 07/12/78 |  | -. 25 |  | . 00 |  |  |  | . 03 |  | . 05 | -. 69 |  | -. 14 | . 24 | -. 03 |  |
| Mexico |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 03/12/78 |  |  |  |  |  |  |  | . 03 |  |  |  |  | -. 34 |  | -. 04 |  |
| Mexico |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 03/19/78 | -. 47 |  | . 08 |  |  |  | -. 37 |  |  |  | -. 67 | -. 03 | -. 16 |  | . 03 |  |
| Mex1co $01 / 11 / 78$ |  |  | . 15 | . 04 |  |  |  |  | . | . 02 |  |  |  | $\ldots .10$ |  |  |
| Guatemale |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 02/19/78 |  |  | . 08 |  | -. 39 |  | -. 31 | -. 04 |  | -. 02 | -. 68 | -. 04 | -. 28 |  | . 09 | -. 35 |
| Guatemalia |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 02/22/78 |  |  |  |  | -. 36 |  |  |  | ; | -. 05 | -. 69 | -. 06 | -. 37 |  |  | -. 48 |
| Guatemala | -. 54 |  | . 08 |  |  |  | -. 31 | -. 07 |  | -. 08 |  | -. 03 | -. 31 |  | . 08 | -. 48 |
| Guatemala |  |  |  |  |  |  |  |  | $\because$ |  |  |  |  |  |  |  |
| 02/22/78 |  | -. 29 |  | -. 06 |  |  |  | -. 06 |  | -. 07 | -. 65 | -. 06 | -. 31 |  |  |  |
| Parama |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| $04 / 04 / 78$ | -. 56 | -. 22 | . 01 |  |  |  | -. 31 |  |  |  |  |  | -. 32 |  |  |  |
| Perus |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 07/28/78 |  | -. 21 | . 18 | . 02 |  |  | -. 30 | . 01 |  |  |  | -. 04 | -. 18 | . 12 |  |  |
| chile |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 11/24/77 |  |  | . 08 | -. 08 |  |  |  | -. 01 | -. 32 |  |  |  |  |  | . 09 |  |
| Chile |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 02/05/78 |  |  |  | -. 04 |  |  | -. 36 | -. 07 | $-.37$ |  |  |  |  |  | . 04 |  |
| Chile |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 03/08/78 |  |  |  |  |  |  |  | -. 04 | -. 34 |  |  |  | -. 32 |  | -. 01 |  |
| Argentina |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 11/23/77 |  |  | . 00 |  |  |  | -. 31 | -. 08 |  | -. 07 |  |  | -. 17 | . 22 |  |  |
| Argentina |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 11/24/77 |  | -. 26 |  | . 03 |  |  | -. 34 | -. 04 |  | -. 00 |  |  | -. 18 | . 18 |  | -. 52 |
| Argentim $11 / 28 / 77$ |  |  |  | -. 04 |  |  |  | -. 05 | -. 30 | -. 02 |  |  |  |  | -. 01 | -. 41 |
| argentio |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 01/17/78 |  | -. 25 | . 07 | -. 03 | -. 35 |  | -. 32 |  |  | . 02 |  | -. 06 | -. 16 | . 20 |  |  |
| Argentima |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 01/20/78 |  |  | .13 | . 02 | $-.34$ |  | -. 33 | -. 06 |  | . 00 |  |  |  |  |  |  |
| $\begin{aligned} & \text { Argentina } \\ & 01 / 24 / 78 \end{aligned}$ |  | -. 22 | . 04 | . 00 | -. 34 |  | -. 34 | . 07 |  | . 04 |  |  | -. 19 | . 16 |  |  |
| Argentina 10/27/78 |  |  |  |  | -. 37 |  |  | -. 07 | -. 26 |  | -. 72 |  | -. 24 |  |  | -. 48 |
| average | $-.52$ | -. 25 | . 09 | -. 01 | -. 34 |  | -. 33 | -. 02 | -. 33 | -. 01 | -. 68 | -. 05 | -. 26 | . 16 | . 03 | -. 46 |

(7) Table 9: Teleseispic P-delays relative to station FL. Residuals are in seconds. Events are from the southwest (180 $-270^{\circ}$ azimuth)

$\overrightarrow{\mathbf{N}}$ Table 10: Teleseismic P-delays relative to station FL. Residuals are in seconds. Events are from the northuest (270 - $360^{\circ}$ azimuth)

| Event | 04 | CH | TJ | LL | 日R | FL. | $W \mathrm{~W}$ | ation FP | IS | BB | SR | HM | HR | PP | BY | GC |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Alaska |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 11/27/77 |  | -. 19 | . 08 | -. 04 | $-.25$ |  | -.28 | . 08 |  |  | -. 45 |  | . 04 | .13 | . 02 | -. 13 |
| Alaska |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 02/12/78 | -. 28 |  | .10 |  | -. 32 |  | $-.30$ | . 02 | -. 10 | -. 18 | $-.47$ |  | . 08 | . 05 | . 08 | -. 16 |
| Yugosl avia |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 01/24/78 |  |  | . 05 | . 04 |  |  | -. 32 |  |  |  |  | . 03 |  | . 09 | -. 02 |  |
| Kuril |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 12/04/77 |  | - | .13 | -. 01 | -. 24 |  | -. 33 | -. 04 | -. 08 | -. 29 | -. 53 | . 08 | -. 05 | . 11 | -. 03 | $-.24$ |
| Kur 11 |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 01/10/78 | , | -. 23 | . 04 | -. 01 | -. 39 |  | -. 34 | -. 00 |  | -. 27 | -. 50 | . 04 | -. 07 | .17 | -. 01 | -. 20 |
| Kuril |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 01/29/78 |  | -. 22 | . 13 | .01 | -. 29 |  | -. 38 | -. 00 | . 00 | . 00 |  |  | -. 08 | .11 | -. 06 | -. 28 |
| Kuril |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 03/24/78 |  |  | . 06 | -. 01 | -. 39 |  |  |  |  | -. 22 | -. 59 | . 01 | -. 07 | . 07 | -. 12 | -. 21 |
| Japan |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 01/11/78 | -. 27 |  |  |  |  |  | -. 27 | -. 03 | . 09 |  | -. 41 | . 10 | . 00 | . 13 | -. 03 | -. 16 |
| Japan |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 02/20/78 |  | -. 19 | . 08 | -. 01 |  |  | -. 36 | -. 02 |  | -. 21 | -. 50 | .11 | -. 03 | . 09 | . 00 | -. 29 |
| Japan |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 02/22/78 |  | -. 16 |  | . 00 | -. 28 |  | -. 36 | -. 03 | . 01 |  | -. 43 | . 10 | -. 04 |  | -. 06 | -. 21 |
| Japan |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 03/07/78 | -. 29 | -. 17 | . 14 | . 05 |  |  | -. 29 | -. 06 |  | -. 22 | -. 51 | . 05 | -. 08 | . 04 | -. 08 | -. 14 |
| Japan |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 06/12/78 |  | -. 18 | . 09 | . 09 |  |  | -. 34 |  |  | -. 26 |  |  | . 02 | . 10 | -. 12 | -. 15 |
| Ural |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 11/30/77 |  | -. 25 | . 07 |  | . .31 |  | -. 35 | -. 11 | -. 02 | -. 32 | -. 44 | . 01 | -. 06 | .16 | -. 07 | -. 26 |
| Russia |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 03/26/78 | -. 31 | -. 20 | . 06 |  | -. 35 |  | -. 33 | -. 06 |  | -. 35 | -. 50 |  | . 04 | . 18 | -. 06 | -. 25 |
| Russia |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| 06/11/78 |  |  |  |  |  |  | -. 38 | -. 11 |  |  | -. 42 |  | -. 10 | . 10 | -. 13 | -. 30 |
| Marianas |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |  |
| .09/26/78 |  |  | . 06 | -. 03 | -. 28 |  | -. 32 | -. 10 | -. 03 |  | -. 48 | . 09 | -. 03 | . 05 |  |  |
| Average | -. 29 | -. 20 | . 08 | . 01 | -. 31 |  | -. 35 | -. 04 | -. 02 | -. 25 | -. 48 | . 06 | -. 02 | .11 | -. 05 | -. 21 |

Figure Captions


#### Abstract

Figure 1. Generalized geologic map of the Pacific Northwest. Stipled area indicates the High Cascades a major zone of Plio-Pleistocene volcanism in western Oregon. Adapted from McBirney (1968).


Fisure 2.
a) Mt. Hood seismic array. All stations were vertical 1 Hz seismometers with a frequency response peaked at 30 Hz .
b) Stations VCR, VTJ, VGC, and VHM were located on the flanks of the volcano.

Figure 3. Locations of 3 blast sites from quarry and construction areas used in determining a preliminary crustal structure. Origin times were recorded for blasts at Trillium Lake Quarry and at the Salmon River S-curve. Stations SSC2, SSC3 and SSC4 were temporary stations installed for the Salmon River blast.

Figure 4. Traveltime curves using a $6.0 \mathrm{~km} / \mathrm{sec}$ reducing velocity for the construction and quarry blasts.
a.) P-traveltime curve for combined Salmon River and Triilium Lake data.
b) P-traveltime curve for the Wildwood data.
c) Composite traveltime curve for all shots. Circles are Salmon River S-curve data points, triangles are Trillium Lake, squares are Wildwood.

> Figure 5. Comparison of actual shotpoint locations from October USOS refraction experiment with calculated locations from the Mt. Hood network. Dark circles are actual locations. Upen sircles are computed shotpoints for the two days of shots.

Figure $\sigma^{a}, b, c$, and $d$ ．
Reduced traveltime curves for four of the USGS uctober shotpoints．Reduced time is observed traveltime minus the staセion－shot distance divijed by a reducing reさxiンy．Eor inis figdre，a reduaing velocity of 6.0 was used．

Figure 7．Locations of the 10 local earthquakes recorded by：the Mt．Hood network during the period of November 20， 1977 to December 1978．Dark circles indicate earthquakes （M．2），the dark star indicates the mainshock of September 6，1978，the open star is the foreshock of the September 6 earthquake and the drcles are aftershocks．

Eigure 8．Eocal mechanism plots．Plots are lower hemisphere，with dark quadrants indicating compression and white quadrants dilatations．

Figure 9．Detailed focal mecnanism plot for the September 6 events．
Darkened circles are compressions，open circles are dilatations

Figure 10. Detailed focal mechanism plots for November 23, 1977, March 30, 1978, and July 14, 1978 earthquakes. Dark circles are compressions and open circles are dilatations.

Figure 11. Average apparent velocity across the Mt. Hood network for 14 regional events.

Figure $12 \mathrm{a}, \mathrm{b}$, and c .
Contour maps of teleseismic P-wave residuals using Fros Lake station (VFL) as a reference station. The azimuth of approach of the teleseism is indicated by the arrow and the times are in seconds.

Figure 13. Polar plots of the teleseismic P-delays corrected for elevation and plotted by delta and azimuth. A 6.0 $\mathrm{km} / \mathrm{sec}$ reducing velocity was used. Residuals are in tenths of seconds.


b


## TRAVEL TIMERED.JCED BY $0.0 \mathrm{KM} / \mathrm{SEC}$




$\cap$


## TRAVEL TIMEREDUCED BY 0.0 KM/SEC





eARTHQUAKE SEPTEMBER 6,1978

(F|G.o)






P-DEAYS $\mathbb{N}$ SECONDS $\because \geq 0$

$$
\begin{array}{lll}
\text { Reqo } \\
\text { Reqiad } & 0 & -.25
\end{array}
$$

$$
-.25 \text { ю }-.50
$$

$$
\text { Hole } \leq-.50
$$

$\therefore$ RELATVVE RESIDUALS IN TENTHS OF SECONDS


