GL01319

GSA un press Lipril-March release

Terrestrial Heat Flow and Crustal Radioactivity in Northeastern New Mexico and Southeastern Colorado

by

C. L. Edwards* Marshall Reiter Charles Shearer Wesley Young

Geoscience Department and New Mexico Bureau of Mines and Mineral Resources New Mexico Institute of Mining and Technology Socorro, New Mexico 87801

*Present address: Los Alamos Scientific Laboratory, University of California, P. O. Box 1663, Los Alamos, New Mexico 87545

ABSTRACT

New heat flow data obtained in northeastern New Mexico and southeastern Colorado show three regional trends. 1) A broad heat flow anomaly associated with the Southern Rocky Mountains contrasts with a narrow and 34° heat flow anomaly between 35.5° N. lat. apparently associated only with the Rio N. lat., Grande rift. 2) The high heat flow anomaly apparently associated with the Southern Rocky Mountains extends 200-300 km onto the Great Plains of northeastern New Mexico and southeastern Colorado. 3) Areas of extensive volcanic activity do not necessarily have high heat flow. In addition, measurements of crustal radioactivity in the vicinity of the Rio Grande rift suggest that the radioactive heat generation contributes uniformly to the surface heat This implies that the heat flow anomaly observed along flow. the Rio Grande rift is caused by tectonic and volcanic sources and not by anomalously high crustal radioactivity.

INTRODUCTION

The purpose of this study was to refine the heat flow map in northeastern New Mexico and southeastern Colorado and to measure crustal radioactivity at drill sites in the vicinity of the Rio Grande rift in order to estimate the contribution of crustal radioactive heating to the surface heat flow. Sixty-five temperature logs were made in northeastern New Mexico and southeastern Colorado yielding fifty-three new unreduced heat flow values (Figure 1). Nine new reduced heat flows have been made along the Rio Grande rift and vicinity in New Mexico and southern Colorado.

PRESENTATION OF HEAT FLOW DATA

Table I summarizes previously unreported heat flow data at 53 sites in northern New Mexico and southern Colorado. The location of these and other published heat flow data is shown in Figure 2.

Heat flow is calculated by multiplying measured linear thermal gradients in drill tests by the corresponding average measured thermal conductivities. A best heat flow value for each site is normally chosen by averaging the heat flow throughout the drill hole. However, if the heat fluxes vary substantially we attempt to choose the heat flow that seems most reasonable in light of the thermal conductivity control, the linearity of the gradient, possible regional groundwater movement or vertical movement of water within the well bore. The heat flow measurements are evaluated in Table I as A, B, or C, using the criteria of Reiter and others (1975).

Figure 2 shows the major basins and uplifts in northern New Mexico and southern Colorado. The hatched_areas show regions of extensive volcanic activity. Heat flow values vary considerably between various volcanic areas. High heat flow values were observed in the San Juan volcanic area (2.2-3.4 HFU) and in the Jemez volcanic field (> 3.0 HFU). Somewhat lower heat flow values were observed in the Mount Taylor volcanic area (1.5-2.0 HFU). Normal heat flow values were observed in the Raton, Clayton, and Capulin volcanic fields (1.3-1.5 HFU). Lipman and others (1973) summarize K-Ar ages of upper Cenozoic basaltic rocks in this region. The ages of volcanism range from 16 to 23 my in the San Juan Mountains, from 0.1 to >9 my in the Jemez volcanic field, from 0.1 to 3 my in the Mount Taylor volcanic field, and from 0.01 to 7 my in the Capulin and Clayton volcanic fields.

Heat flow varies from basin to basin within northern New Mexico and southern Colorado. In northwestern New Mexico and southwestern Colorado the San Juan and Blanding basins exhibit normal heat flows (1.3-1.5 HFU). In northeastern New Mexico and southeastern Colorado the Las Vegas and Raton Basins have higher heat flows (>2.0 HFU). In east central New Mexico the Tucumcari basin has a normal heat flow (1.4 HFU).

Figure 3 shows a heat flow contour map of northern New Mexico and southern Colorado. In northeastern New Mexico and southeastern Colorado this map is a revision of a previous heat flow map published by Reiter and others (1975). The contours are solid lines in areas where there is good control; that is, it appears unlikely that additional data will significantly change the character and location of the contours. Dashed contours indicate that additional data may change the location of the contours but probably not the character. Question marks are used where data coverage is sparse and considerable ambiguity exists in the location and character of contours.

The Great Plains has been considered an area of low to normal heat flow (\leq 1.5 HFU); however, this characterization does not appear to be generally valid in northeastern New Mexico and southeastern Colorado where the heat fluxes vary from 1.3 to over 2.5 HFU. A broad heat flow anomaly in northern New Mexico and southern Colorado associated with the Southern Rocky Mountain complex contrasts with a narrow heat flow anomaly between Santa Fe (about 35.5° N. lat.) and Socorro (about 34.0° N. lat.), New Mexico, apparently associated only with the Rio Grande rift. Between 35.5° N. lat. and 38.5° N. lat. the change from heat flow values of 2.5 HFU, observed near the Rio Grande rift, to heat flow values of <1.5 HFU, observed on the Great Plains, occurs over a distance of 200 to 300 km. However, between Santa Fe and Socorro, New Mexico, the change from heat flow values of >2.5 HFU to heat flow values of <1.5 HFU occurs over a distance of 20 to 30 km. Local anomalies in the Las Vegas basin, the Raton Basin, near Pueblo, Colorado, and near Questa, New Mexico, suggest that the heat flow pattern in the Southern Rocky Mountains is more complex than shown in Figure 3. More data will be necessary to determine the size and character of these anomalies.

PRESENTATION OF RADIOGENIC DATA

Reduced heat flow values were calculated according to the definition given by Roy and others (1972). The best heat flow value at each site is reduced by the amount AH where A is the radiogenic heat production at the site and H is 10 km. Roy and others (1972) and Decker and Smithson (1975) used H \approx 10 km to calculate reduced heat flows in southern New Mexico. Potassium, uranium, and thorium concentrations were measured for each site using samples from the drill hole and the heat production was calculated by the technique described by Swanberg (1972) and by Rybach (1971).

Figure 4 shows the locations of reduced heat flow sites in New Mexico and southern Colorado. The reduced heat flows and associated radioactivity data are presented in Table 2. Within the Southern Rocky Mountain region the reduced heat flows at Los Alamos (>3.2 HFU) and at Crested Butte (2.1 HFU) are higher than the intercept q* (1.3-1.4 HFU) predicted by Roy and others (1968) for the region. However the reduced heat flow values at Questa (0.52 and 0.74 HFU) are lower than the predicted intercept. East of the Rio Grande rift, between Santa Fe and Socorro, New Mexico, the reduced heat flow values at San Pedro (0.89 and 0.85 HFU) compare closely with the 0.3 + 0.1 HFU intercept for the Great Plains as suggested by Roy and others (1968) and Decker and Smithson (1975). West of the Rio Grande in southern New Mexico, the reduced heat flow value at Animas Peak (1.1 HFU) is intermediate between Basin and Range and Great Plains values. A reduced heat flow of 1.4 HFU, a typical Basin and Range value, was measured in the southern region near Orogrande. This differs from the value of 2.0-2.6 HFU that Decker and Smithson (1975) measured about 15 km to the south.

Figure 5 shows a plot of the heat flow versus heat production for the Rio Grande rift and vicinity of New Mexico and southern Colorado. The solid line represents the linear heat flow - heat production relation for the Basin and Range (Roy and others, 1968). Of the 19 values plotted, only four, Lordsburg, Santa Rita, Orogrande North, and Animas Peak plot within ±20% of this linear relation. The vertical dashed line, representing a crustal radioactive heat production of 0.38 HFU, is the average heat generation of all sites measured in this study excluding Questa and Sierra Blanca, where the data plotted more than two standard deviations from the mean radioactive heat production. This implies that the radioactive heat generation in the Rio Grande rift vicinity is relatively constant, with about 0.38 HFU of the unreduced heat flow coming from radioactive decay in the upper crust. DISCUSSION

The Southern Rocky Mountain complex in northern New Mexico and southern Colorado is characterized by a broad region of high heat flow with heat flows greater than 2.5 HFU along the Rio Grande rift and in _ the Las Vegas and Raton Basins. This broad high heat flow region narrows significantly between Santa Fe and Socorro, New Mexico, although more heat flow data would better determine the continuity of the high heat flow ribbon between Albuquerque and Socorro. This heat flow pattern suggests widely distributed and perhaps deeper thermal sources under the Southern Rocky Mountain complex of northern New Mexico and southern Colorado, as opposed to narrowly distributed, shallower thermal sources under the Rio Grande rift. Between 35.5° N. lat. and 38.5° N. lat. the transition from heat flows less than 1.5 HFU on the Great Plains occurs over a distance of 200 to 300 km. The same heat flow transition occurs over 15 to 35 km between Santa Fe and Socorro, New Mexico (for example along 34.5° N. lat.). The narrow width of the heat flow anomaly along the Rio Grande rift between Santa Fe and Socorro, 15 to 35 km, implies that the width of the thermal source at depth is probably no greater than 15 to 35 km. If one assumes that a continuous subsurface thermal source underneath the Rio Grande rift is solely responsible for the narrow ribbonlike heat flow pattern between Santa Fe and Socorro, then the associated half width of the observed heat flow anomaly suggests that the thermal source is no deeper than 15 to 35. This conclusion agrees with geochemical data presented by Lipman km. (1969) and with seismic data presented by Sanford and others (1973). The broad heat flow anomaly in the Southern Rocky Mountains implies more widely

distributed sources probably deeper than those under the Rio Grande rift.

Gough (1974) shows that areas having high heat flow generally have high electrical conductivity in the mantle. Caner and others (1967) and Reitzel and others (1970) have shown that the electrical conductivity anomalies follow closely the boundaries of structural. provinces with the electrical conductivity in the mantle increasing from the Great Plains to the Southern Rocky Mountains and then decreasing again under the Colorado Plateau. The new heat flow data presented in this manuscript show a broad heat flow anomaly associated with the Southern Rocky Mountains extending onto the Great Plains of northeastern New Mexico and southeastern Colorado (Figure 3). Caner and others (1967) discuss an electrical conductivity anomaly which appears coincident with this high heat flow anomaly. However, Porath and Gough (1971) postulate the existence of two distinct electrical conductivity anomalies superimposed in this area, one a mantle feature associated with the Southern Rocky Mountains and the other an upper crustal anomaly due to the conducting sediments of the deep Andarko Basin of southwest Oklahoma. Caner (1970) points out that such a reinterpretation of the original data is possible. The new heat flow data are consistent with the original interpretation of the geomagnetic data by Caner and others (1967) and are not consistent with the more recent interpretation of geomagnetic data by Porath and Gough (1971) and Caner (1970).

Porath and Gough (1971) show a model for a geomagnetic profile at 38° N. lat. The depth to the conducting layer is 350 km under the Great Plains, 150 km under the Southern Rocky Mountains and 350 km

under the Colorado Plateau. Porath (1971) shows an alternative model to fit the same data. The depth to the conducting layer under the Great Plains is 160 km; however, the conducting layer under the Southern Rocky Mountains is modeled as a ridge and step structure with the shallowest depth to the conducting layer as 45 km. New heat flow data presented in this manuscript show high heat flow anomalies in the Las Vegas Basin, the Raton Basin, and near Pueblo Colorado (Figure 3). Present heat flow data indicate a small decrease in heat flow under the Southern Rocky Mountains near Questa. The complex thermal structure in this area suggests that a ridge and step electrical conductivity model for the mantle is reasonable. The high electrical conductivity under the Southern Rocky Mountains implies high mantle temperatures (Madden, 1970; Gough and Porath, 1970). The broad high heat flow anomaly associated with the Southern Rocky Mountains may reflect these high mantle temperatures.

With data presently available it is not possible to determine if the source of the high heat flow anomaly that extends onto the Great Plains is a continuation of the source of the high heat-flow anomaly associated with the Southern Rocky Mountains. It is possible that two separate high heat flow anomalies exist, one coincident with the Southern Rocky Mountains and the other associated with the Raton and Las Vegas Basins. The superposition of two anomalies with the possible masking effects of deep groundwater motion could produce the observed anomaly. Suppe and others (1975) have proposed a hot-spot trace across New Mexico from the White Mountain volcanic field of eastern Arizona to the Raton-Clayton volcanic field in the Great Plains of northeastern New Mexico with the hot-spot presently near Raton, New Mexico. This hot-spot could possibly be the source of the high heat flow anomaly in the Great Plains.

Heat flow data measured in volcanic areas (rigure 2) vary significantly from one area to another. The volcanics of the Raton-Clayton fields are derived from the mantle with essentially no crustal contamination (Jones and others, 1974), the heat flows measured in this area range from 1.2-1.5 HFU. The tholeiitic basalts within the Rio Grande rift in northern New Mexico and southern Colorado probably fractionated at a depth of 15-20 km (Lipman, 1969; Lipman and others, 1973). This area is characterized by heat flows greater than 2.5 HFU. The volcanic rocks in the San Juan volcanic area of southwestern Colorado erupted from a cluster of central vent volcanoes (Eardly, 1962; Lipman and others, 1973). The western and central portion of the San Juan volcanic area contain abundant cauldron subsidence structures (Luedke and Burbank, 1968; Steven and Epis, 1968). This area has high heat flows (2.2-3.4 HFU), but the values are not consistently as high as those within the northern Rio Grande rift. The Jemez volcanic area in north-central New Mexico is the result of multiple eruptions which have produced an immense cauldron structure. The heat flows on the western side of the Jemez volcanic area are greater than 3.0 HFU (Reiter and others, 1976b.) The Mount Taylor volcanic field in north-central New Mexico is the result of a large central vent volcano. The heat flows in the Mount Taylor area range from 1.5 to 2.0 HFU (Reiter and others, 1975).

The heat flow data within the basins of northern New Mexico and southern Colorado also show considerable variation between basins. The anomalously high heat flows in the Las Vegas and Raton Basins contrast with normal heat flows in other basins within the region, the San Juan, the Blanding and the Tucumcari Basins. The Las Vegas and Raton Basins have been extensively intruded by igneous rocks (Johnson and others, 1966;

Johnson, 1968). The Tucumcari Basin in the Great Plains of eastern New Mexico, and the San Juan and Blanding Basins within the Colorado Plateau, have had little igneous or volcanic activity (Eardley, 1962).

The results of the crustal radiogenic measurements in the Rio Grande rift vicinity (Decker and Smithson, 1975; this study) suggest that radioactive decay in the upper crust contributes about 0.38 HFU to the unreduced heat flow within the area (Figure 5). This implies that the heat flow anomalies observed in New Mexico and southern Colorado along the Rio Grande rift, (this study; Reiter and others, 1975; Decker and Smithson, 1975) are not caused by lateral variations in the concentrations of potassium, uranium and thorium within the crust. The Rio Grande rift high heat flow anomaly observed by Decker and Smithson (1975) in southern New Mexico and by Reiter and others (1975) throughout New Mexico and southern Colorado is probably a result of non-radioactive thermal sources underneath the Rio Grande depression.

The low reduced heat flow values at Questa, 0.52 HFU and 0.74 HFU, are below the 1.4 HFU intercept of the Southern Rocky Mountain-Basin and Range heat flow province. The reduced heat flow value of 0.59 HFU at Sierra Blancais below the 0.8 HFU intercept for the Great Plains heat flow province. The heat generation measured at these sites is three times greater than the average for the Rio Grande rift vicinity. It is possible that the samples measured at these sites are not representative of the upper crust. The reduced heat flow values of 0.89 and 0.85 HFU at San Pedro suggest that this area is part of the Great Plains heat flow province. Reduced heat flow values of 2.1 HFU at Crested Butte and >3.2 HFU in the Jemez Mountains suggest the presence of additional nonradioactive, thermal sources within the crust and upper mantle at these sites. The reduced heat flow value of 1.1 HFU at Animas Peak is somewhat low for the Basin and Range. The value of 1.4 HFU at Orogrande is appropriate for the Basin and Range. The uncertainty of all reduced heat flow measurements in the area is quite high, and caution should be exercised when considering the significance of a value at a single site.

REFERENCES CITED

Birch, F., 1947, Temperature and heat flow in a well near Colorado Springs: Am. Jour. Science, v. 245, p. 1-18
-----, 1950, Flow of heat in the Front Range, Colorado: Geol. Soc. America, Bull., v. 61, p. 567-630
Caner, B., 1970, Electrical conductivity structure in western Canada and petrological interpretation: Jour. Geomagnetism and Geoelectricity, v. 22, no. 1-2, p. 113-129

Caner, B., Cannon, W. H., Livingstone, C. E., 1967, Geomagnetic depth sounding and upper mantle structure in the Cordillera region of western North America: Journ. Geophys. Research, v. 72, no. 24, p. 6335-6340

Costain, J. K., and Wright, P. M., 1973, Heat flow at Spor Mountain, Jordan Valley, Bingham, and La Sal, Utah: Jour. Geophys. Research, v. 78, p. 8687-8698

Decker, E. R., 1969, Heat flow in Colorado and New Mexico: Jour. Geophys. Research, v. 75, p. 550-559

Decker, E. R., and Smithson, S B., 1975, Heat flow and gravity
interpretation across the Rio Grande rift in southern
New Mexico and West Texas: Jour. Geophys. Research,
v. 80, no. 17, p. 2542-2552

Eardley, A. J., 1962, Structural geology of North America: New York, Harper and Row, p. 553-582

Gough, D. I., 1974, Electrical conductivity under western North America in relation to heat flow, seismology, and structure: Journ. Geomagnetism and Geoelectricity, v. 26, p. 105-123 Gough, D. I., and H. Porath, 1970, Long-lived thermal

structure under the Southern Rocky Mountains:

Nature, v. 226, p. 837-839

- Herrin, E., and Clark, S. P., 1956, Heat flow in west Texas and eastern New Mexico: Geophysics, v. 21, p. 1087-1099
- Johnson, R. B., 1968, Geology of the igenous rocks of the Spanish Peaks region Colorado: U. S. Geol. Survey, Prof. Paper 594-G, 47 p.

Johnson, R. B., Dixon, G. H., Wanek, A. A., 1966, Late Cretaceous and Tertiary stratigraphy of the Raton Basin of New Mexico and Colorado: New Mexico Geol. Soc., Guidebook 7th field conf., p. 88

- Jones, L. M., Walker, R. L., Stormer, J. C., Jr., 1974, Isotopic composition of strontium and origin of volcanic rocks of the Raton-Clayton district, northeastern New Mexico: Geol. Soc. America, Bull., v. 85, p. 33-36
- Lipman, P. W., 1969, Alkalic and tholeiitic basaltic volcanism related to the Rio Grande Depression: Southern Colorado and Northern New Mexico, Geol. Soc. America, Bull., v. 80, p. 1343-1353

Lipman, P. W., Bunker, C. M., Bush, C. A., 1973, Potassium, thorium, and uranium contents of upper Cenozoic basalts of the Southern Rocky Mountain region, and their relation to the Rio Grande depression: Jour. Research, U. S. Geol. Survey, v. 1, no. 4, July-Aug. 1973, p. 387-401

1/5

Lovering, T. S., 1948, Geothermal gradients, recent

climatic changes, and rate of sulfide oxidation in the San Manuel district, Arizona: Econ. Geology, v. 43, p. 1-20

Luedke, R. G., and Burbank, W. S., 1968, Volcanism and

cauldron development in the western San Juan Mountains, Colorado: Colorado School Mines Quart., v. 63, p. 175-207 Madden, T., 1970, Geolectric upper mantle anomalies in the

United States: Journ. Geomanetism and Geoelectricity v. 22, no. 1-2, p. 91-95

- Porath, H., 1971, Magnetic variation anomalies and seismic low-velocity zone in the western United States; Jour. Geophys. Research, v. 76, no. 11, p. 2643-2648
- Porath, H., and Gough, D. I.,1971, Mantle conductive structures in the western United States from magmetometer array studies: Royal Astron. Soc., Geophys. Journ., v. 23, p. 387-398
- Reiter, M., Edwards, C. L., Hartman, H., Weidman, C., 1975, Terrestrial heat flow along the Rio Grande rift, New Mexico and Southern Colorado: Geol. Soc. America, Bull., v. 87, p. 811-818
- Reiter, M., Simmons, G., Chessman, M., England, T., Hartman, H., Weidman, C., 1976a, Terrestrial heat flow near Datil, New Mexico: New Mexico Bureau Mines Mineral Resources Ann. Rept. 1975-1976, p. 33-37
 - Reiter, M., Weidman, C., Edwards, C. L., Hartman, H., 1976b, Subsurface temperature data in Jemez Mountains, New Mexico: New Mexico Bureau Mines Mineral Resources, Circ. 151, p. 1-15

Reitzel, J. S., Gough, D. I., Porath, H., Anderson, C. W. III, 1970, Geomagnetic deep sounding and upper mantle structure in the western United States: Royal Astron. Soc., Geophys. Jour., v. 19, p. 213-235

- Roy, R. F., Decker, E. R., Blackwell, D. D., Birch, F., 1968, Heat flow in the United States: Jour. Geophys. Research, v. 73, p. 5207-5221
- Roy, R. F., Blackwell, D. D., Decker, E. R., 1972, Continental heat flow, <u>in</u> The nature of the solid earth: New York, McGraw Hill, p. 506-544
- Rybach, L., 1971, Radiometric techniques, in Modern methods of geochemical analysis, Wainerdi and Uken, Eds.: New York-London, Plenum Press, p. 271-318
- Sanford, A. R., Alptekin, O., Toppozada, T. R., 1973, Use of reflection phases on micro-earthquake seismograms to map an unusual discontinuity beneath the Rio Grande rift: Bull. Seis. Soc. America, v. 63, p. 2021-2034

Sass, J. H., Lachenbruch, A. H., Monroe, R. J., Greene, G. W., Moses, T. H., 1971, Heat flow in the western United

States: Jour. Geophys. Research, v. 76, p. 6376-6413 Spicer, H. C., 1964, Geothermal gradients and heat flow in the

Salt Valley anticline, Utah: Boll. Geofisica Teor. ed Appl., v. 6, p. 263-282

Steven, T. A., and Epis, R. C., 1968, Oligocene volcanism in south-central Colorado: Colorado School Mines Quart., v. 63, p. 241-255

Suppe, J., Powell, C., Berry, R., 1975, Regional topography, seismicity, quaternary volcanism, and the present-day tectonics of the western United States: Am. Jour. Science, v. 275-A, p. 397-436

Swanberg, C. A., 1972, Vertical distribution of heat generation in the Idaho batholith: Jour. Geophys. Research, v. 77, p. 2508-2513

Warren, A. E., Sclater, J. C., Vacquier, V., Roy, R., 1969, A comparison of terrestrial heat flow and transient geomagnetic fluctuations in the southwestern United States: Geophysics, v. 34, p. 463-478

ACKNOWLEDGMENTS

This work was supported by National Science Foundation Grant GI-32482, New Mexico Energy Research and Development Program Proposal #5, and New Mexico Institute of Mining and Technology - the Geoscience Dept. and the New Mexico Bureau of Mines and Mineral Resources. Some of the equipment used was purchased on Bureau of Reclamation Contract #14-06-500-1875. Salary for student assistance with the gamma ray spectrometer was supported by New Mexico Water Resources Institute Grant 3109-405 funded by New Mexico Energy Research and Development Program.

The following organizations and individuals gave permission to log boreholes on their property and/or under their supervision. Their help and cooperation is gratefully acknowledged: American Fuels, AMOCO, CONOCO Minerals, Ben Donegan, Duval, Inspiration Development, Kaiser Steel, Molycorp, National Park Service, USGS.

			TABLE I	SUMMARY OF HE	AT FL	UW DATA	15 0 1 4 1		DECT	01141 574
LOCALITY AB301/AST (NM) AB301/ADTH (NM) AD301/ADTH (NM) ALUM CZEK (C) ARKUYD TETILLA (NM) BLUEMATER(NM) BUCKMAN 2 (NM) BUCKMAN 2 (NM) BUCKMAN 2 (NM) CANDW CITY EHAPMT (C) CANDW CITY EHAPMT (C) CARAILOS 42 (NM) CERRILLOS 42 (NM) CERRILLOS 42 (NM) CHAPELL SPADE (NM) CHAPELL SPADE (NM) CHIMINEY CREEK (NM) COLMOR/WEST (NM) DIXON 42 (NM) GRAMEROS RD (C) JACKSON 4 (C)	NOR 1 6 6 2 4 5 5 3 3 6 6 1 2 6 4 5 5 3 3 3 5 5 2 7 7 6 5 3 3 5 5 2 7 7 6 5 5 3 3 5 5 1 6 6 1 5 6 0 1 5 0 1 5	WEST ELE 104 0.6 1.87 104 1.4 201 105 1.0 205 104 1.0 205 104 1.0 205 104 4.2 2.1 106 0.9 1.80 107 0.8 1.83 106 2.9 2.31 106 2.9 2.31 106 0.6 1.84 106 0.6 1.84 106 0.6 1.84 106 0.7 1.82 103 5.1 1.22 104 5.2 2.43 104 4.1 1.82 105 4.8 2.07 104 4.1 1.82 105 4.8 2.03 104 4.7 1.52 104 4.7 1.52 104 4.7 1.52 102 3.2 1.25 <	TABLE I TABLE I PERKS O PERKS O INTETE4605 O INTETE4750 O INTETE1800 O INTETE1800 O INTETE1800 O INTE1800 O INTE1800 O INTE1800	SUMMARY OF HEA THERMAL GRADIENT DEG.C/XN9 25.35+00.127 30.89911.13 36.19120.499 33.9442.322 48.4910.11 31.2810.34 35.9140.11 31.2810.34 36.2210.46 24.39410.11 31.2810.34 39.2010.46 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 24.3910.20 25.3100.20 26.3910.20 26.3910.20 26.3910.20 26.3910.20 27.441	AT FLU N 26 802783048739366182857456556	JH DATA TYPE CON DPF SEC FRAGMENTS 5.	IERMAL DUC I IVI TY -DEG C 2040.19 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.40 3010.20 300.20 3000.20 3000.20 300.20 3000.20 3000.20 3000.20 3000.20	ATW 0626013139000000000000000000000000000000000	AEST FLOH HEAT FLOH ESTIMAT FLOH HEUA FLOH 1.62 2.00 2.60 2.44 1.73 2.04 1.96 1.83 1.20 1.20 1.21 40 2.88 2.87 5.47 2.42 2.41 2.23	QUALITY DE FLO HEAT FLO VALUE AA BC BC CC CC A CC B C C C B C C C C B C C C C
LAS VEGAS #1 (NA) LAS VEGAS #2 (NM) LITTLE BEAR #14 (NM) LITTLE CROW CR (NM) MAES (NM) MI DORA (VM)	35 35 35 34 34 17 36 49 36 09 36 28	105 15 194 105 15 195 107 15 190 104 41 243 104 42 197 103 34 173	9 80-150 150-198 100-170 1 100-170 2 60-210 200-290 9 9 110-225 2 20-90 9U-170 9U-170 8 190-290 4 190-290	# 17.78±0.48 27.25±0.52 22.5±0.544 # 29.15±0.58 29.15±0.36 38.12±0.46 36.12±0.46 43.12±0.46	0439675858	FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 6. FRAGMENTS 7. FRAGMENTS 7.	2210.19 5010.40 4510.48 9010.28 9010.28 9010.12 5610.57 3210.03 9310.61	1.00100.04 1.50100.04 1.50100.14 1.45100.10 1.72100.06 2.13100.25 2.44100.00 1.64100.00 1.64100.00	1.50 1.45 1.72 2.13 2.35 1.30	С вс вс вс
NOLAN/EAST (N4) NOLAN/WA (N4) ORTIZ MIN 2 (N4) ORTIZ MIN 3 (N4) ORTIZ MIN 3 (N4)	36 58 36 39 36 11 35 19 35 18 36 53	104 53 192 104 35 192 104 40 186 106 10 239 106 10 221	4 80-100 115-180 1 80-140 140-180 170-240 9 140-720 8 130-280 370-420 9 40-160	* 30.85±1.17 30.85±1.17 43.37±0.29 55.31±1.11 43.17±0.17 64.62±0.80 18.18±0.11 16.03±0.13 16.03±0.13 22.00±0.39 38.00±0.73	8885856738 26738	FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS* 5. FRAGMENTS* 5. CORE 7. CORE 7. COR	23+0.15 23+0.60 23+0.60 23+0.20 23+0.20 32+0.20 32+0.29 32+0.77 36+0.10	1.9310.02 2.7010.26 2.8010.27 3.2710.17 3.2710.17 1.3310.06 1.2510.12 1.4010.17	2.75 2.98 1.33 1.33 2.69	C C B 4
PUEBLO/AEST (C) PUEBLO/AEST (C) QUESTA #2 (NH) QUESTA # 3 (NH)	38 11 38 18 36 42 36 42	104 44 157 104 46 152 105 31 293 105 32 290	160-280 0 80-130 130-230 4 80-130 190-240 3 310-450 590-690 590-690 0 310-410	50.58±0.92 51.69±1.13 25.31±1.10 51.77±0.70 30.36±1.17 18.52±0.17 22.70±0.32 21.22±0.32 20.32±0.16	555552472	FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. CORE 7. CORE 8. CORE 8. CORE 8.	35:0.04 15:0.57 71:0.23 71:0.23 71:0.23 77:0.23 77:0.24 25:1.09 25:1.0.28 25:1.0.28	2.71±0.0 2.60±0.30 2.71±0.30 2.60±0.30 2.71±0.30 3.65±0.20 1.71±0.20 1.76±0.10 1.76±0.10 1.76±0.20	2.69 2.85 1.74	C C A
RED CREEK (C) RIO CUCHARAS (C) San Ckistobal (N4) San PEDRO 4 (A4)	38 14 37 32 36 38 35 15	105 00 213 104 54 210 105 39 231 106 11 210	430-510 510-590 4 50-1590 4 100-240 290-450 7 140-290 5 190-410 0 120-330	23.32*0.30 21.21*0.41 22.14*0.41 31.91*0.30 34.62*0.59 35.52*0.59 25.74*0.15 30.69*0.25	4747 11 20 14	CORE 7. CORE 8. FRAGMENTS 10. FRAGMENTS 0. FRAGMENTS 5. FRAGMENTS 8. FRAGMENTS 6. FRAGMENTS 6. FRAGMENTS 6.	77-0.26 19-0.34 84+0.28 09-0.13 76-0.31 79-0.15 26-1.56 96+0.43	1.81±0.0 1.74±0.1 2.40±0.1 1.99±0.1 3.12±0.1 1.42±0.3 2.44±0.1	2.40 7 1.97 3.12 7 1.42 5 2.44	C B 4
SHALE HILLS/NE (C) SIERRA DEL DJITO (C) SILOAM ROAD (C) SIMMS (NM) SOCORRU/S (NM)	37 46 37 16 38 14 36 08 33 57	103 37 137 105 14 230 104 56 213 104 43 192 106 56 155	2 50-100 100-145 3 250-350 350-470 5 40-100 5 40-100	$\begin{array}{c} 32.70 \pm 0.70\\ 64.36 \pm 0.84\\ 0.64.20 \pm 0.35\\ 58.30 \pm 0.57\\ 538.74 \pm 0.59\\ 643.62 \pm 0.34\\ 0.34 \pm 0.59\\ 0.34 \pm 0.62\\ 0$	-66875677	FRAGMENTS 4. FRAGMENTS 4. FRAGMENTS 4. FRAGMENTS 4. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5. FRAGMENTS 5.	79-0.13 79-0.13 03-0.17 17-0.05 15-0.57 13-0.12 03-0.40	1.57±0.0 2.22±0.1 2.59±0.1 2.43±0.0 2.00±0.2 2.67±0.0 1.13±0.1	3 1.90 2 2.51 5 2.00 7 2.71 1 1.73	C A C B C
TETILLA PEAK (NM) TRES MENTOSA (NM) TURQUDISE MIN 1 (NM) TURQUDISE MIN 2 (NM) VAN BREMMER CR (NM)	35 35 34 05 35 30 35 31 36 48 36 38	106 13 188 107 22 209 106 06 19 106 07 19 106 57 232	10-140 90-200 200-240 200-250 200-250 200-240 20-240 200-250 200-240 200-250 200-250 200-250 200-250 200-250 200-250 200-430		103285667 24	FRACMENTS 5 FRACMENTS 4 FRACMENTS 4 CORE 4 CURE 4 CURE 5 FRACMENTS 5 FRACMENTS 5	55-56 66-10.07 91-10.09 70-10.09 15-10.29 14-10.23 14-10.30 69-10.30	2.01+0.0 2.721+0.0 2.321+0.0 1.01+0.0 1.01+0.0 2.13+0.0 2.13+0.0 2.13+0.0 2.13+0.0 2.13+0.0 2.13+0.0 2.13+0.0 2.13+0.0	9 2.00 5 1.96 3 1.32 5 1.22 8 2.04 7 1.44	С С А А
N IS HUMBER OF THERMAL 0 INFU = 1 UCAL/CMO 0 CONDUCTIVITIES OF 0 OUTCADP SAMPLES 1 THIS DEPTH INTERVA 1 INDICATES THE STAM 1 C2, SITE IN CULURADD: ELEVATIONS ARE ± 20 ME	C (NDUCI CM-SEC FRASHENI L NOT US DARO DEI (NM), S TERS	TIVITY SAMPLE TSAMPLES HAN SED IN DETER: VIATION SITE IN NEW P	E BEEN COM IN ING BEST EXICO	RRECTED FOR PUR F HEAT FLOW EST	ROSIT	r E				

[ab]	le	2.	Summary	of	Radioactivity	and	Reduced	Heat	Flow	Data	
------	----	----	---------	----	---------------	-----	---------	------	------	------	--

Well Name	Rock Type	No. of Samples	Potassium (%)	Uranium (ppm)	Thorium (ppm)	Heat Generation (HGU)	Reduced Heat Flow (HFU)
Crested Butte+	Monzonite	5	4.10±.20	1.95± .15	5.41±.51	3.05	2.10
Questa/East ⁺	Granite	3	0.78±.55	13.02±2.01	31.71±.57	12.85	0.52
Questa #2	Granite	5	1.49±.24	13.61±1.23	25.07±1.43	12.97	0.74
Los Alamos*	Granite	4	3.52±.17	0.92±.39	15.95±3.23	3.09	>3.20
San Pedro #3 ⁺	Monzonite	6	2.55±.18	2.51± .30	10.95± .80	3.98	0.89
San Pedro #4	Monzonite	4	3.33±.11	3.39± .07	12.02±.24	4.88	0.85
Sierra Blanca ⁺	Monzonite	4 ;	3.36±.63	9.97±3.24	28.65±6.78	11.76	0.59
Animas Peak ⁺	Granite	5	3.96±.20	5.18± .33	14.50±.58	6.55	1.05
Orogrande ⁺	Granodiorite	6	3.18±.31	1.96± .20	7.37±.25	3.18	1.43

⁺unreduced heat flows from Reiter and others (1975).

· · ·

.

11

""" (1976b).

.

Figure 1. Heat flow sites in the Four Corners States. Solid diamonds indicate data from Birch (1947, 1950), Lovering (1948), Herrin and Clark (1956), Spicer (1964), Roy and others (1968), Warren and others (1969), Decker (1969), Sass and others (1971), Costain and Wright (1973). Open diamonds indicate data from M. Reiter and others(1976a, 1976b). Solid circles indicate data from Reiter and others (1975). Open circles in Arizona, Utah, western and southern New Mexico, indicate unpublished data of Reiter and others. Open circles in northeastern New Mexico and southeastern Colorado, within the hatched area, indicate data from this manuscript. Sites where data are disturbed such that no heat flow information is obtainable are indicated by X.

Figure 2. Heat flow sites in northern New Mexico and southern Colorado. Data in parenthesis from Reiter and others (1975) and Reiter and others (1976a, 1976b). Data in brackets from Decker (1969) and from Sass and others (1971). Other data presented in this manuscript.

Figure 3. Terrestrial heat flow contour map of northern New Mexico and southern Colorado. Contour interval, 0.5 HFU. Plus signs indicate control sites measured by New Mexico Institute of Mining and Technology; Xs indicate control sites measured by other investigators.

Figure 4. Reduced heat flow sites in New Mexico and southern Colorado. Reduced Heat flow data in parenthesis from Decker and Smithson (1975). Other reduced heat flow data presented in this manuscript.

Figure 5. Heat generation versus heat flow in New Mexico and southern Colorado. (+) reduced heat flow site from this manuscript; (x) reduced heat flow site from Decker and Smithson (1975)



Figure I.





FIGURE 3





· · · · · ·

• • •

: .