UNIVERSITY OF UTAH **RESEARCH INSTITUTE**

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Ed Sammel Water Resources Division USGS 345 Middlefield Rd., MS 67 Menlo Park, CA 94025

Dear Ed:

Attached are the Petroleum Information data I promised to send you. Included are descriptions of 71 wells in South Dakota, which comprise all of our well temperatures in that state, and a few wells in south-central North Dakota.

Also attached are a format tabulation and formation codes to help you with the data listings.

I hope these attachments will be useful to you. Please call or write if you have any questions.

Sincerely,

Ross W. Whipple

RWW/smk

Encl.

cc: Mike Wright

PET. INFO. - SOUTH DAKOTA WELLS WITH BHT'S BY COUNTY No. WELLS COUNTY WITH BHTS 4 BUTTE CORSON 7 CUSTER 1 DEWEY 3 9 FALL RIVER 4 HAAKON 30 HARDING HUGHES 2 JACKSON / JONES 1 MELLETTE / PENNETON 2 3 PERKINS STANLEY 2 SULLY 1 TOTAL FOR STATE 71

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PAGE 861		06/08/78	PETROLEUM INFORMATION, CORP. PAGE 848		
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		40 N 6 E 21	10002 0445043310153493 0445010510153481		33 N 163 N
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		1	105 SPUD 11/15/1965 COMP 12/20/1965 POT567 05055 00		
		1	107 DTO 4323 40101 DST 01 4001- 4083 0001UNKN 4042 DC2 VV		
		10 N 6 E 21 26	40131 FINGL OP 1H IFF 26 FFP 31'BHT 1425		
	1	40 N 6 E 21 2	10002		33 N 163 H 1
			101 SDAX (HARXON) A 660 FSL 660 FHL C SH SULHF HF 102 EXETER DRLG ETAL 29-13 BIERINGEN		
		•	103 2340 KB 2330 GR WILDCAT 104 1880 DF 1872 GR F LSE NO AP1 40-055- 20008-00		
		· · · · · · · · · · · · · · · · · · ·	105 SPUD 05/24/19/0 CONF 05/25/19/0 KOTHAY ULH 107 DTD 2575 SPUD 02/24/19/0 CONF 05/25/19/0 KOTHAY ULH	-	· ·
	· .		40131 FINAL OP 1H IFP 782 FFP 818 BHT 120F?		
		40 N 10 E 6 25	10002 0447946010334147 0447941410334103 10010 019 20010273 9999950 9006020K0T 2900471		
· · · · · · · · · · · · · · · · · · ·			10021 HAP N HULL ROLE & SEC 25 1991HER HULLS 101 SDHX BUTTE A 660 FSL 660 FHL CC H SD HF HF 102 KOCH FYDORRUNN		33 N 163 H 1
			103 2959 KB 2949 GR HILDCAT 104 2956 DF 2949 GR P API 40-019- 20010 00		t i
	i		105 SPUD 06/29/1969 COMP 07/02/1969 ROTARY DBA 107 DTD 2900	-	
			40101 DST 01 2553-2626 6027009 3 3 7 6 002 40131 FINAL 0P 0450M IFP 45 FFP 60.8MT 92F		
	/	40 N 12 E 22 4	10002 0450319210144288 0450311610144353 10010 041 20005273 9999950 710203RDRV 5055472		•
			10021 THP N +2 ROLE 22 SEC 4 1981ACH HULLS 101 SDAY DEHEX A 2345 FNL 660 FEL SE NEWF HF	-	33 N 163 H 1
		i e _e e	102 INVESTERS DALG VENTURES 3 HOLLOGARY 103 2357 KB 2347 GR HILDCAT 104 2762 DF 2378 GR PLSE NO OPL HULDCAT		
· · ·		· · · · · · · · · · · · · · · · · · ·	105 SPUD 07/07/1971 COMP 08/07/1971 ROTARY TA-0 107 DTD 5055 FM/TD 203/07/1971 ROTARY FM/TD 203/07/1971		
		· • •	40101 DST 01 4975-5045 203RDRV 50 10 002 40131 FINAL OP 2H IFP 65 FFP 230 BHT 130F		
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			101 SDAK DEHEY A 3665 FNL 660 FHL NH SH D D 102 INVESTORS DRLG7 HOLLOWAY	1 · t ·	
			103 2358 KB 2348 GR UNNAMED 104 API 40-041- 20009 00 105 SPUS 6241043970 COMP 0140741971 POTOPS		
	1	•	105 STUD 02/17/17/17/17/17/17/18/18/18/18/18/19/17/17/17/18/18/18/19/17/17/17/17/17/17/17/17/17/17/17/17/17/		
			40201 DST 02 5006-5066 203RDRV S004 40210 WTR TS IN 1H50M		·
			40231 FINAL OP 2H IFP 1162 FFP 2203 BHT 175F 50102 203RDRV PERF JET 5051-5051 006		
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· · · · · · · · · · · · · · · · · · ·	40.N 14 E 4 29	107 DTD 3241 LTD 3240 FM/TD 602SKCK 40101 DST 01 3174- 3241 602NCSL 017 40102 OVERLAPS 602SKCK 40131 FINAL OP 1H IFP 47 FFP 94 BNT 117F	.	33 N 163 H 8 33 N 163 H 8
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	40 N 14 E 11 20	40131 FINAL OF 200 17F 1030 FFF 1110 BM 987 10002 0451565010280602 0451571810280593 10010 105 20006273 8900550 901419MN S 10021 TMP N LH RGE E 11 3EC 20 1980 RK-HULLS 101 SDRK CERKINS A 660 FSL 1980 FML SE SHAFF MF /S		<u>33 n 163 h 8</u>
		102 1240 1M 1 W LWHAR OND 103 2599 KB 2588 GR 4 LSE NO API 40-105- 20006-00 104 259 KB 2588 GR 4 LSE NO API 40-105- 20006-00 105 SPUD 09/17/1971 COMP 09/30/1971 ROTARY 08A-G 107 DTD 5000 LTD FH/TD 4199HNLS 40101 DST 01 3240-3301 602007 32.7/ 5003		
· · · · · · · · · · · · · · · · · · ·	40 N 14 E 15 24	40131 FINAL OP 450 1FP 1104 FFP 1208 BH1 1375 26 40201 DST 02 4708-4791 419904LS 47 50 004 40231 FINAL OP 14304 IFP 242 FFP 598 BHT 1675 10002 0451599410222981 0451599410222911 10010 105 2001 277 0495950 04051599410222911	.	<u>33 N 163 N 8</u>
· · · · · · · · · · · · · · · · · · ·	/	10010 10021 TUF N_LLL-RGE E 155750-20020380240 1001 SDAK GERK INS A 1980 FSL 05570-RAHNCIPRI 101 SDAK GERK INS A 1980 FSL 1980 FL 102 SHOKEY OTL 33-24 ACKERNAR-FD LIF HF /6 103 2433 KB, 2423 GR HILDCAT 104 RPI 40 105 20011 00 105 SPUD 11/27/1976 COMP 02/20/1976 ROTRRY D&A 107 DTD 6503 FM/TD 203RDRV	-	•
· · · · · · · · · · · · · · · · · · ·	40 N 14 E 16 8	40101 DST 01 6290-6311 203RDRV 63 507 002 40131 FINAL OP 2H IFP 1860 FFP 2590 BHT 172F 3 10002 0451878610218947 0451876510218923 10010 20010 9026050 900203RDRV 6454775 10010 20010 9026050 900203RDRV 6454775	-	33 N 163 H 8
		101 SDAK (PERKING A 1880 FSL 1995 FEL SL 1995 FEL SL IF HF / 7 102 TRUE OTE 33-8 HILEER 103 2381 KB 2369 GR HILDCAT 104 RPI 40 105 20010 00 105 SPUD 06/14/1975 COMP 06/28/1975 ROTARY D&A		
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PAGE 864	06/08/?8	F 2 PETROLEUM INFORMATION, CORP PAGE 851	-	
E GEOTHERMAL LISTING		SUB FILE GEOTHERMAL LISTING	1	06/08/78
	40 N 14 E 16 8	107 DTD 6448 LTD 6454 FM/TD 203RDRV 40101 DST 01 6220-6230 203RDRV 002 40131 FINHL OP 3H IFP 34 FFP 54 BHT 158F 40201 DST 02 6290-5340 203RDRV 003 40231 DST 02 1H IFP 2014 FFP 2062 BHT 170F	· · ·	33 N 163 H 87 33 N 163 H 87
	40 N 15 E 2 29	10002 10010 063 20017273 10010 063 20017273 10010 063 20017273 10021 THP N 15 ' RGE E 2 SEC 29 19BLACK HILLS 101 SORK HERDING A. 1980 FSL 1980 FSL 0 HUSE HE		· · ·
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			40101 DST 01 6220-6230 2038084 6225 002 40131 FINSE OP 34 IFP 34 FFP 54 BHT 158F		33 N 163 N 8
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			103 3548 KB 3536 GR HILDCAT		
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		40 N 15 E 5 21	10002 C452437110352527 C452448310352497		· .
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			40231 FINAL OP 1H30H IFP 242 FFP 598 BHT 167F		[
		40 N 15 E 6 36	10002 0452147210333148 0452152910333116 10010 063 20016273 9120550 900553hRSN 3959969		
			10021 THP N 15-RGE E 6 SEC 16 1980 RCK-HHLS 101 SDAK SUBPOLND A 735 FSL 1980 FEL SU SEJUF UF QO		33 N 163 H 8
			102 UNION OTE OF CALIFORNIA 1 STATE=217 103 2975 KB 2966 GR HILDCAT		
			104 2985 DF 2977 GR P LSE KO API 40-063- 20016-00 105 SPUD 08/19/1969 C0HP 08/24/1969 R0TARY D&A		
	- I		107 LTD 3959 2568 FH/TD 553hRSN 40101 DST 01 3343-3393 602hDDY 002		
			40102 OVERLAPS 602SKCK 40131 FINAL OP 1H IFP 835 FFP 1115 BHT 128F		1 \$ ~ ~
		40 N 15 E 7 4	10002 0452860210327029 0452861210326968		-
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		40 N 15 E 74	107 DTD 4250 FM/TD 553MRSN 40101 DST 01 3650- 3716 602MDDY 002		77 1 167 11 01
			40131 FINAL OP 11430M IFP 1222 FFP 1267 BHT 130F 40201 DST 02 4110- 4190 602FLRV 003	•	22 H 102 H 81
			40231 FINAL OP 1H IFP 1084 FFP 1354 BHT 163F	· · -	
		40 N 17 E 1 1	10002 0454701610394358 0454708410394328 10010 063 20065273 5651350 9015045NNN 1533773	i .	
		· · · ·	10021 THP N 17 RGE E 1 SEC 198LACK HILLS 101 SDAK HARDING A 1120 FNL 1980 FEL 52 NH NE HF	. i	
			102 MONT-DAKOTA UTILITIES 31X-1 STATE 103 3553 KB 3548 GR HIDCAT	, 	·
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			107 DTD 1533 FM/TD 6045NNN 40101 DST 01 1328-1455 6045615 S002		
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 DTD 8535 LTD 8537
 FM/TD 203R0RV

 40101 DST 01
 8236-8310 203R0RV
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 40102 OVERLAPS 259SLRN
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 152 FFP 329 ANT 180F
 33 N 163 F 40131 FINAL OP 2H IFP 132 FFP 329 BHT 180F 40 N 19 E 25 10002 0456331710104267 0456339110104133
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PAGE 866	06/08/78	PETROLEUM INFORMATION?	CORP PAGE 853 SUBFILE GEOTHERMAL LISTING		06/08/78
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	102 007 103 3072 104 105 SPUD (107 010 1 20102 20380	KB 3060 GR BUFFRLO 07/17/1975 GO GPI 40 063 20090 00 07/17/1975 COMP 10/12/1975 ROTRRY 01L 01L 8652 PB 8558 FM/TD 203R0RV RV RV PERF W/ 2/FT 8426- 8468 GROSS 006			33 N 163 H
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 	✓ 40 N 20 E 4 10 10002 10010 063 09 10021 THP N 101 SDRK 101 SDRK 102 PENNZ 103 2018	0457144010361571 0457155710361357 5121273 6000 9999961 110203RDRV 203RDRV 8420166 299 REE E 4 SEC 10 198LRC+H4LLS 4RRDING A 1980 FN± 2180 FEL SH NE 0 D0 51L C0 32-10A TILUS 122-10A TILUS	27		
·····	104 2038 105 SPUD 107 CTU 5 20102 203R0 40101 DST 01	OF 2023 GR P LSE NO 29 RP1 40-063- 05121 00 11/21/1965 COMP 01/01/1966 ROTARY 01			M 201 N 27
· · · · · · · · · · · · · · · · · · ·	40131 FIXed 50102 203RD 50202 203RD 50302 203RD 50302 203RD	0P 3H IFP 33 FFP 429 BHI 1964 RV. PERF H/ 4/FT 8375- 8376 004 RV PERF H/ 4/FT 8375- 8376 005 RV PERF H/ 4/FT 8375- 8376 006 0456992610351297 0456988510351187			· . I
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PAGE 867 LE GEOTHERMAL LISTING	06/08/78	I 2 PETROLEUM INFORMATION,	CORP PAGE 854 SUBFILE GEOTHERMPL LISTING		06/08/72
·····	40 N 20 E 5 16 50202 203RDP 40 N 20 E 5 22 10002 10010 063 20	V PERF H/ 4/FT 8362- 8368 008 0456815410349554 0456828910349630 0492273 2342950 911203R0RV 8500572		-	32 N 163 H 8
	10021 THP N 101 SDAK H 102 UEPCO 103 2926 104 2727	20 RGE 5 SEC 22 198LRCK HILLS MRDING A 1980 FSL 2029 FML NE SH HF INC 23-22 FEDERAL-GRUSE KB 2914 GR HILDCAT DF 2718 GR HILDCAT DF 2014 00 <			·,·
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	40 N 20 E 18 32	104 2381 DF 2371 GR P API 40-0317 20014 00 105 SPUD 10/23/1970 COMP 11/17/1970 ROTARY D&A 107 DTD 7515 FFL/TD 2024MPG $($ 40101 D\$1 00 24 F5490 3520550 5 455 002.	33 N 163 33 N 163
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	- 40 N 21 E 2 26	10002 043 20001273 5480505 9002730808 8600167 10021 TWP N 21 RGE E 2 SEC 26 1981 RCK-HILLS 101 SORK GRR0ING A 1980 FSL 660 FEL ME SEJ WF WF 33 102 NIRAI OTT FROULERS 1 PAINTERSTIC 103 3179 KB 3168 GR HILDCAT	
		104 3176 ES 3168 GR P RP1 40-063-20001 00 105 SPUD 07/04/1967 COMP 09/02/1967 ROTRRY DBA 107 DTD 3600 \$550-8580 2139029 \$550-450 \$550-450 40101 DST 01 8550-8580 2139029 CUSH 650 HTR 002 40131 FINRL 0P 24 IFP 303 FFP 677 BHT 184F	
	40 N 21 E 3 22	10002 0457713210373543 0457734410373416 10010 063 05126273 6000 1371510 911203808V 8805369 10021 TUP N 21 REFE 3 SEC 22 198LACK-HULLS 101 SDAK (HERDING) A 1980 FNL 1980 FEL (SH KE) 40 34 102 CHOPPLIN PETROLEUM 1 CLARKSON ESTATE	<u>33 N 163</u>
		103 3280 K8 BUFFRL0 104 3288 ES 3280 38 P RPI 40-063- 05126 00 105 <u>SPUD 06-07/1958 (OHP 12/31/1958 ROTRAY</u> DRA-06 107 0T0 8805 P8 8559 EN/TO 203808V 40101 DSi 01: 6840- 6910 <u>353(RLS</u> 6 6 C DD2 40101 DSi 01: 6840- 6910 <u>353(RLS</u> 6 6 C DD2 40101 DSi 01: 6840- 6910 <u>153(RLS</u> 6 6 C DD2 40101 DSi 01: 6840- 6910 <u>153(RLS</u> 6 6 C DD2 40101 DSi 01: 6840- 6910 <u>153(RLS</u> 6 6 C DD2)	· · · · ·
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		42631 FIRML OF IN IFP 23 FFP 848 BHT 1500 50102 203R0RV PERF 8572-8576 009 50103 203R0RV PERF 8584-8602 010 50202 203R0RV PERF 8572-8602 GROSS 011 50302 203R0RV PERF 4/-3747-8529-8541 013 50400 - 203R0RV PERF 1/-3747-8529-8541 013	
		SOBO2 ZOJRDRV PERF I/ 3/FT 8529- 8541 015 SOBO2 ZOJRDRV PERF 8529- 8541 016 SO702 ZOJRDRV PERF 8529- 8541 016 SO702 ZOJRDRV PERF 8529- 8541 016 SO702 ZOJRDRV PERF 8529- 8541 017 SO802 ZOJRDRV PERF 8529- 8541 018	33 N 164
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PAGE 869	06/08/78	K 2 PETROLEUM INFORMATION, CORP PAGE 856	06/08/78
	40 N 21 E 3 22	SUB FILE GEOTMERTIAL LISTING 50902 203RDRV PERF 8529- 8541 019 51002 203RDRV PERF 8529- 8541 020 51102 203RDRV PERF 8529- 8541 020 51102 203RDRV PERF 8529- 8541 020 51102 203RDRV PERF 8529- 8541 020	33 N 164 I
	40 N 21 E 21 5	51202 203R0RV PERF 8529-8541 022 51302 203R0RV PERF 8529-8541 023 10002 0458161410154121 0458166210154161 10010 031 20018273 1270550 90115300kD 7400876 10021 TUP N 21 PGE F 21 SEC 5 0557h PR UK (PGI	
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	-	4.	107 DTD 2400 40101 DST 01 4867-4885 JS2HSHC 4876 005 005		
			40201 DST 02 6325-6362 2038080 (244 S006 40231 F1N94 0P 1H IFP 569 FFP 1554 BHT 174F		
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			105 SPUD 06/06/1969 COMP 06/16/1969 ROTARY D8A-G 107 DTD 4039 LTD 4040 FM/TD 353MIDL		
			40101 DST 01 1521- 1551 603x9RR HISRUN 002 40201 DST 1A 1521- 1551 603x9RR HISRUN 003		
			40401 DST 1C 1521-1551 602MBRR /536 005	1	
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			104 2098 DF 2090 GR P LSE NO API 40-031- 20007-00 105 SPUC 06/03/1969 COMP 06/06/1965 ROTARY DOA		
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		. ··· .	104 3245 DF 3237 GR P LSE NO API 40-063- 20030-00 105 SPUD 12/04/1969 COMP 01/22/1970 ROTARY D&A-OG 107 DTD 9065 JTD 9066 SM/TD 2020994		
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		103 2924 KB 2911 GR IIILOCAT 104 2921 DF 2911 GR P LSE NO API 40-063- 20073-00	
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	40 N 22 E 21 36	10002 C458204810146456 0458208310146576 10010 031 05056273 8325050 911203R089 6680966 10°21 THP N 22 RGE E 21 SEC 36 1981ACK_HULLS 101 SCORE CREECIL SEC 1980 EUR	
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		104 2247 DF 2238 GR P API 40-031- 05056 DD 105 SPUD 08/23/1965 C019 11/01/1965 R0TARY D8A-0G	í .
	•···	107 LTD 6680 FM/TD 203RDRV 40101 DST 01 6464- 6495 203RDRV S003	• .
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	40 N 23 E <u>3 24</u>	10002 0459413410368740 0459401310368710 10010 063 20099273 99999 6665010 911203R0Rv 9187776	· · · · · · · · · · · · · · · · · · ·
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	40 N 23 E 3 31	100C2 0453196810379194 0459182610379234 10010 063 20091273 3466050 900203RDRV 8897K75	40 N 3
		10021 MARTIN 63 RULE 3 SEL 31 LYDLMLK HILLS	;
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		105 SPUD 10/29/1975 COMP 11/30/1975 ROTARY // 004	
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		10010 063 20144273 50000 666501 110203R0RV 203R0RV 9246377 10021 THP N 23 RGE F 4 SFC 19 (RSFH 924637	
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	40 N 23 E 3 31	10002 10010 063 20091273 3466050 900203R0RV 8897K75 10021 THP H 23 RGE E 3 SEC 31 1981 RCK HILLS 101 S FORM GENERING & CAD FM CAD FEL & K WE THE UF 47		40 N 3 E
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		105 SPOD 107 27 173 CONF 11730 1973 KONKET FH/TD 203RDRU 107 LTD 8897 FH/TD 203RDRU 40101 DST 01 <u>2669-8695 203RDRU</u> 868-3 002 40131 FINAL OP 1H IFP 62 FFP 291 BHT 209F	-	40 N 3 E
	/ 40 N 23 E 4 19	10002 10010 063 20144273 50000 66501 11020300RV 2038080 9246377 10021 Tup N 23 RGE E 4 SEC 19 055 <u>TH_PRIMCIPAL</u> 10, 2040 HARDING A 1600 FSL 2400 FEL SEL NH SED HO HOE		1 1 1
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	;/·····	40210 DST 02 9090- 9142 20 JPDRV 9/1/C S003 40210 GAS TS IN 1H52H 40231 FINAL OP 2H IFP 387 FFP 1082 BHT 215F		40 N 4 E
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		102 HOROUGER PLANNING 41-22 OTERNESS 103 2956 KB 2942 GR UNNINED 104 2968 DF 2958 GR P LSE NO API 40-063- 20064-00 105 SPUD 05/06/1973 COMP 08/27/1973 ROTARY OIL 107 DT 9230 FM/TD 203RDRV		
¥ - Ø-	· · · · · · · · · · · · · · · · · · ·	20102 203R0RV PERF W 2/FT 8970-9150 GROSS 015 40101 DST 01 8945-9075 203R0RV 90 / 0 S002 40131 FINAL OP 1H15M IFP 135 FFP 234 BHT 222F 19.64 40201 DST 02 9080-9170 203R0RV 9/ 0 S003	-	40 N 5 F
		40210 GAS TS IN 40H NOT GAUGED 40231 FINAL OP OHSSM IFP 172 FFP 247 BHT 224F 19.61 50102 203RDRV PERF JET H/ 2/FT 8970- 8974 003 50202 203RDRV PERF JET H/ 2/FT 9038- 9050 GROSS 009		
	· ·	50302 203RDRV PERF JET H/ 2/FT 9094- 9150 GROSS 010 50402 203RDRV PERF JET H/ 2/FT 8970- 9150 GROSS 011 50502 203RDRV PERF 8970- 9150 GROSS 012 50602 203RDRV PERF 8970- 9150 GROSS 013	· · .	1
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PAGE 874	06/08/78	PETROLEUM INFORMATION, CORP PAGE 861 SUBFILE GEOTHERMALLISTING		06/08/78
	10 N 27 5 5 22			
	40 N 23 E 5 30	10002 0459302210354447 0459291910354359 10010 053 20126273 60000 2940010 911203RDRy 9200178		40 N 5 E 40 N 6 E
		10021 HE M 20 KNE E D SEL SU UDDHERKINCHML 101 SDAK HARDING A 1800 FNL 660 FEL N2 SE NE HO D 102 HEBB RESOURCES 30-8 NJOS-JANVRIN 103 3048 KB 3038 GR STATE LINE 104 001 HO 163 20126 DD		
		105 SPUD 06/03/1977 COMP 06/14/1977 ROTARY 084-06 107 DTC 9200 LTD 9191 FH/TD 203RDRV 9002 1010 DST 01 : 8986-8998 203RDRV S002 40131 DST 01 : 8986-8998 203RDRV S002		·
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			10021 TWP N 23 REFE 5 SEC 30 US51H_BALLE IPAL 101 SDRX GERVING A 1800 FNL 660 FEL GR2 SE NED HO D 5 0 102 WEBB RESOURCES 30-8 NJ0S-JARVEN N 103 3048 KB 3038 GR STATE LINE 104 RPI 40 063 20126 00 /		· · · · · · · · · · · · · · · · · · ·
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		40 N 23 E 6 27	10002 0459285610336328 0459266810336052		40 N 6 E
			10010 053 20084273 99999 5025051 110203R0RV 203R0RV 9270775 10021 THP N 23 RGE E 6 SEC 27 199LRCK HHLLS 101 SDAK HARDING A 2640 FNL 1950 FEL E2 CHF HFD 5		
			103 2907 KB 2895 GR UNNAMED 104 2904 KB 2895 GR UNNAMED		
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		40 N 23 E 6 32	107 DTD 9110 PB 9060 20102 203RDRV PERF H/ 2/FT 8858- 9018 GR055 006 40101 DST 01 8818- 8886 203RDRV S002		40 N 12 E 2
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			103 2812 K9 2800 GR WILDCAT	1	



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	l. ,	40 N 111 H 79 1	10002 044442510028374 0444427010028343 10010 065 20002273 150070 9016020K0T 1667K75		
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			101 30445 1027 COP 1 1510 13C 1510 14C 534 1C 5475 3		40 N 14 E
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K6778W	506 0021	Forest Service	CP-N	Perkins	45.885°/102-195°	300
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	591-0073	Forest Service	GR-N_	Perkins	T20N R16E S6	308
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T.	595-0058	Forest Service	GR-N	Perkins	T21N R16E S2	298
<u>172</u>				Dealtre	45.6860/102.1640	2/1
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K67-76W	THE CONTRACTOR				45.643°/102.503°	
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	590-0063	Forest Service	GR-N	Perkins	T20N R14E S3	385
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<u>_K6778Y</u>	594-0090	Panada Carrier	CR-N	Perking	$45.83/0/102.3/2^{\circ}$	357
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PI REF. #	WELL NAME	COMPANY/OPERATOR NAME	TYPES OF LOGS RUN	COUNTY	LATITUUL,	
		AUSTIN QUAD				
G7199A	#2 Tillie Juergen	Burns Petroleum	GR-N	Fayette	30°/97° 03.75'	5000
071000		CORPUS CHRISTI QUAD		······································		
<u>G/1990</u>	2 H.J. Mosser	Texas Energy Exploration	GR-N	Jim Wells	27°54.56'/97°57.50	5334
G7199C	- 	CRYSTAL CITY QUAD			28° 41'/98° 23.20'	
	#2 Smith	Tipperary Oil & Gas	GR-N	Frio		5010
G7199E		SAN ANTONIO QUAD			29° 16.70'/98°02'	2660
a	#1 Patillo Higgins Jr.	Holloway Oil Co. SEOUIN OUAD	<u> </u>	Wilson		2660
G7199B	#3 Robin_B. Unit	T.D. Coffman	GR-N	Fayette	29 ⁰ 45.25'/97 ⁰ 7.85'	5000
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Resistance: RS, Resistivity-R, Gamma Ray-GR, Neutron: N, Density-Den, Temperature-T, Caliper-Cal, Spectral Gamma-KUT, Self-Potential-SP

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L 1994 **(** 1994 - 1997) AREA SD UNIVERSITY OF UTAH VICINITY of Philip Haakon research institute Biblio. Haakon County, 5. Dakota_ EARTH-SCIENCE LAB. \mathcal{D} Aldophson, D.G., Le ROUX, E.F., 1968, Temperature Variations of deep sowing wells in S. Dakota: USGS PP 600 P, p. D60-62. Ŀ Canda, D.J., 1975, A Study of the Radium (content or the Graundwaters in western s. Dak witemphasis on the Madison (Pathasopa) Limestene & Ph. & Dissentation, S. Dak Schold Minesd Technology, 58p. l' Davis, R.W., Dyer, C.F., Pawell, J.F., 1961, Progress rept on wells penetrating artesian aquifers in 5. Dakotaus65, WSP 1571 Æ Freeman R., Meier, R., 1978, Potential Appl. of Madison Fm. Waters For community Heatings: Johns Hepkinsu, Appleed Physics Lob, QM-78-042 (2 (?)تھ ا 5. Dak Acad Sci. Proc., vol. 50, pp 61-65. Gues, John P., 1977, Crothermal Applications on the Madison (Pahasapa) Squifer system in S. Dakota: S. Dak. Sch & nuneso Tech, Final Rept oct 1,4776 - Sept 30, 1977 - Sept (A) (25) Konikau C.F., 1976, Frel. Digital Model or orainduater Flow in (25) the madison Bp., Powder River Basin + Adj Areas, Wy., Mt., S. Dak. N. Dak, & Nebaska: USGS, Water Res ances Sources (3-75) (T) Mississippian Linestone, 5. Dakota S. Dak Sch. M+T, Backelon's Thesis. (MAD. Ģ LUM, Daniels, 1961, Gravity measurements east of the Black Hills and along a ine from Rapid Uty to Scourt Fails, 5 Dakota 5 Dak Gool Survey, Rept. IN. 88, 26 p. SPAN) F Mutfler, L. J. P., 1978, psot. of Geothermal Perances of the U.S. -1978. USG 5, Enc. 790, pp 92-93, 131. KSU Petsch, B.C., 1967, Magnetemeter map of South Dakota: 5. Dak beel.


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GEOTHERMAL POTENTIALS IN SOUTH DAKOTA Robert A. Schoon and Duncan J. McGregor

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> For, a geothermal reservoir to exist, there must be a heat source, a suitable reservoir rock above the heat source, water in the rock to transmit heat in the form of hot water or steam, and usually a cap rock above the geothermal reservoir rock to prevent rapid dissipation of heat. This last requirement does not appear to be absolutely necessary in that Allen and Day (1927) measured temperatures of 101.5°C $(214.7^{\circ}F)$ at a depth of 3 feet at Steamboat Fumarole in the geyser area of California. In this area there is no apparent cap rock (McDonald, 1972, p. 342).

> In areas of active geothermal production, all have some like similarities in that they are found in regions of geologically young mountains, usually associated with active or past volcanic activity or where existence of faults are the rule rather than the exception. This is not to say, however, that every area meeting the above criteria overlies geothermal reservoirs. The relationship between any of the geographic or geologic variable is not all that clear cut. For examples, at Larderello the nearest volcanic rock lies 15 miles to the north and the Hawaiian volcanic areas seem to possess only minimal potentials for geothermal energy. In South Dakota the Homestake Gold Mine has a rather low thermal gradient in that at a depth of 6800 feet the temperature is only 122°F (personal communication, Mr Olin Hart, Chief Geologist, Homestake Mining Company) This is in a structurally deformed area that exhibits evidence of nearby Tertiary volcanic activity but has some of the lowest geothermal gradients in the State.

SOURCES OF HEAT

Sources of heat to create geothermal reservoirs involve much speculation but there appears to be general concurrence that the reservoir rock must have a direct relationship to a cooling igneous mass." According to McNitt (1963, p. 37) those areas throughout the world where geothermal energy is used are all located in regions where Cenozoic volcanism has occurred. This indicates that a direct relationship exists between thermal areas and processes of volcanism and magmatic intrusions. However, there is not complete agreement on this point Levorsen (1967, p. 423) deals with aspects of the causes of the phenomenon by stating, "The source of heat of the upper few miles of the earth's crust may be in the outward flow of heat from the central core of the earth, in the presence of igneous magmas that are cooling, in the disintegration of radioactive elements, or in the heat of subcrustal

thermal convection currents. Lesser amounts of heat include the frictional heat formed during diastrophism ... and exothermal chemical reactions that take place within permeable reservoir rock, both of which sources, if present, are temporary and local in their effects." Levorsen previously suggested (p. 419) that in some cases change in geothermal gradient is best explained by a change in thermal conductivity of the rocks.

It is a known fact that temperature increases with depth in bore holes and that the rate of temperature increases or the thermal gradient varies considerably from place to place. There are a number of ways to formulate the phenomenon of the geothermal gradient, perhaps the most common is given by:

$$G = \frac{T \cdot tF}{D}$$

where G = geothermal gradient, T = formation temperature (°F), tF - mean annual temperature (°F), and D = depth in hundreds of feet. The formula is self-explanatory and is the same as that used in computing the thermal gradients found in this report with only minor variations.

In this report the mean annual temperature was considered to be a constant 45°F over the entire State. Also, the annual mean temperature affects subsurface temperature down to a depth of 60 feet Because of this, in computation of thermal gradients, the depth of the well was considered as 60 feet iess than factual to account for the influence of the annual mean temperature. The effect of these deviations is minimal and for practical purposes corresponds to the above mentioned formula.

Mr. O. M. Phillips (1968, p. 138) states, "Gradients as small as 1°C per 140 m are measured in some locations and as large as $1^{\circ}C$ per 10 m in others. but in spite of this, the average over many such drillings in many countries of the world is very close to 1°C per 30 m." If the foregoing is converted into degrees Fahrenheit and feet we find that some localities have geothermal gradients as low as .36°F/100 feet, some as high as 5.1°F/100 feet, with the average world geothermal gradient of 1.7°F/100 feet.

Schuster (1973) reports the average worldwide geothermal gradient is 87°F per mile (1.6°F/100 feet). At this gradient the boiling point of water would be reached at a depth of about 2 miles. Any area having a geothermal gradient several times that of the worldwide average certainly wairants investigation as a potential geothermal area.

A casual inspection of the isogradient map of the State (fig. 1) reveals that large areas have geothermal gradients considerably higher than the world average.



Figure 2. Photograph of the municipal well at Midland, S. D. The small building in the foreground houses the Midland city well. Hot water from this well heats the gymnasium in left background and the classroom building in right background.



Figure 3. Photograph of the municipal well at Philip, S. D. The well yields water at a temperature of 158° Fahrenheit.

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In northern Gregory County the $7^{\circ}F/100$ feet isogradient is present. Theoretically, this isogradient indicates that if one were to drill a hole 2400 feet deep, the temperature at the bottom of the hole (if the annual mean temperature is taken to be $45^{\circ}F$) would be approximately $212^{\circ}F$, the boiling point of water under atmospheric pressure. If this same hole were continued to a depth of 27,600 feet, the temperature should be about $1976^{\circ}F$, equivalent to the temperature of molten lava being ejected from an active volcano.

POSSIBLE CAUSES OF HIGH GEOTHERMAL GRADIENTS IN SOUTH DAKOTA

The study of geothermics is in its infancy, but much attention currently is being focused on the subject. In part this is because our nation is faced with a heating oil and natural gas shortage and additional heating sources are being sought. Then also, the exploration geologist has always been intrigued by the "anomaly" or the departure from the norm. From previous pages it is known that geothermal anomalies do exist in South Dakota. The cause of the hot artesian wells in South Dakota is speculation at this time. However, mechanics or conditions suggested by previous authors are given brief consideration.

Inspection of the Vertical Intensity Magnetic Map of South Dakota (Petsch, 1967) reveals a marked change in the general configuration of the contours. This change occurs in an area a few miles west of, and parallel to, the Missouri River. Many scientists contend that magnetic maps reflect structural trends or Precambrian rock types and younger magmatic intrusives. In South Dakota the magnetic map (fig. 9) is believed to be a reflection of basement rock types and/or structural trends and may mark a boundary of Precambrian Provinces. In North Dakota Laird (1964) separated these rock types into the Peace River, Superior, and Churchill Provinces of the Canadian Shield area. Similar to North Dakota, much of the structure and sedimentation in South Dakota may be governed by relative movements of two blocks or provinces of the shield area. Any relative movement between these two blocks would generate heat and create fracture zones which are two main requisites for an exploitable heat or power source. If this interpretation, based on the magnetic map, is correct then the logical area to search for geothermal energy is near the boundary of the Peace River and Superior Provinces This boundary drawn by Laird in North Dakota has been projected through South Dakota and appears as figure 10 of this report.

Unstable areas are often accompanied by volcanic activity and hot thermal waters. It has been shown that the thermal gradient in south-central South

SOUTH DAKOTA

Dakota is higher than the national average; however, it is more difficult to establish volcanic activity in the area during or subsequent to Precambrian time Data concerning Precambrian rocks are limited, but Steece (1961) indicates that extrusive rocks are present in northwestern Hyde County. This extrusive is approximately 40 miles east of the inferred boundary of the Peace River Superior rock masses, which is a rather extreme distance to postulate for a lava flow. This extrusive rock does strongly suggest the existence of an unstable area during, or subsequent to, Precambrian Time.

A comparison of the isogradient map with the gravity map of the State is less striking. The gravity map of South Dakota is guite generalized but in the area of Murdo (see fig. 11) a negative anomaly of -110 milligals is present in Jones, Mellette, and small portions of northwestern Tripp and southwestern Lyman Counties. Lum (1961, p. 6) in his gravity traverse of the State recognized this negative regional anomaly just east of Murdo, but he also (p. 7) discovered the existence of a broad positive anomaly of 11 milligals superimposed on the regional gravity minimum. He speculated this 11 milligal positive anomaly (fig. 12, this report) is possibly related to the Stanley County magnetic high. The authors do not disagree with Lum but also recognize that the positive anomaly could, be a reflection of a post-Precambrian volcanic intrusive body. If this is the case the thermal gradient should be higher than that of the surrounding area. From observation of the isograident map it is apparent that some of the highest recorded temperatures in the State are present near the gravity anomaly east of Murdo. This area may be an early target in the event of future exploration.

An additional reason for singling out the area is to show the reader a difference of 11 milligals does exist between two surveys. If differences such as this do occur on a state-wide basis, the interpretation of the gravity map of the entire State could vary greatly and very conceivably agree quite closely with the magnetometer map in suggesting that the boundary of the Peace River-Superior Provinces exists in South Dakota. If the gravity map does indicate the Peace River-Superior Provinces then the boundary probably occurs along the -70 milligal contour which roughly follows the course of the Missouri River (see figs. 10 and 11).

According to McNitt (1963, p. 41) gravity and magnetic surveys do not appear to be particularly helpful in locating production fissures within structural depressions. He does note that detailed land magnetic surveys reveal magnetic lows over thermal areas due to hydrothermal alteration of magnetite to pyrite. In south-central South Dakota, Petsch (1967) shows two rather extensive magnetic

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SOUTH DAKOTA

lows situated in eastern Mellette and Todd Counties and in northern Gregory and southern Brule. Counties. Lum (1961, p. 6 and 7) notes that the broad positive gravity anomaly just east of Murdo, may be a reflection of basement rock that is more mafic in character than the surrounding basement rock. This gravity anomaly lies just to the north of the magnetic low in eastern Mellette and Todd Counties. Lum's gravity traverse passed to the northof the magnetic anomaly of northern Gregory and southern Brule Counties. However, Lum (p. 5) does state "a large positive anomaly of 15 milligals is located immediately west of Chamberlain. This anomaly is too large to be caused by either basement topography or by structure in the sedimentary rocks. Therefore, it is probably caused largely by a density contrast in the basement rocks." The proximity of this gravity anomaly to the magnetic low in southern Brule County warrants additional study.

From Levorsen (p. 11, this report) it is stated that heat is evolved from the disintegration of radioactive elements. If such a source is present near the area of Midland, the source must be in rocks of Precambrian Age, because Gamma-Ray logs of sedimentary rocks in the area exhibit no intervals of extraordinarily high gamma activity. However, there is no evidence, such as cores of Precambrian rocks with which to lend credence to the preceding statement. On the other hand, Koenigsberger (1910) states that in the locality richest in radium, the pitchblende deposit in Joachimstal, the temperature gradients are normal. At this time the heating effect that radioactive decay has upon thermal gradients is problematical.

Any evolutionary process that produced structure in South Dakota would no doubt generate heat. Due to the masking effect of the Pierre Shale in the central and south-central part of the State, no structural evidence of sufficient intensity to generate a great amount of heat has been recognized in the area of higher thermal gradients. However, earthquakes are results of structural adjustments. Inspection of the earthquake locations in South Dakota (fig. 13 and table 2) suggests that structural adjustments are occurring in the vicinity where high thermal gradients are present. Heck and Eppley (1958, p. 39) reported an earthquake on May 9, 1906, in eastern Washabaugh County that was noted all along the Niobrara Valley from Rushmore (Nebraska) to Valentine (Nebraska). This quake was felt over an area of 7,000 to 8,000 square miles. The preceding authors also reported an earthquake occurring on July 23, 1946, at Wessington, South Dakota. This quake was generally felt from Pierre to DeSmet eastward and northward to Redfield, South Dakota. Agnew and Tullis (1962) report the occurrence of an earthquake 6 miles west of Pierre that had its focus at a depth of 10 miles. This report of the earthquake near Pierre is interesting in that the

depth of origin or focus was located at a depth of 10 miles or 52,800 feet. From the isogradient map (fig. 1) the thermal gradient is at a rate of approximately 3.5°F/per 100 feet of depth. Multiplying the thermal gradient by the depth in hundreds of feet (3.5°F x 528) the temperature at the focus of the quake is projected to be 1848°F. This compares quite closely with the temperature of molten lava (1976°F). Thus, if the geothermal gradient is correct, and if this gradient is projected downward to a depth of 10 miles it is readily apparent that the temperature of the rocks is sufficient to cause these rocks to be in a plastic state and to yield to structural adjustment by flow rather than by rupture with subsequent shock. Perhaps a knowledge of the geothermal gradient of an area may enable one to accurately forecast the focus of earthquakes in any given area.

At any rate, the foregoing examples do not constitute a complete history of earthquakes in South Dakota, but do indicate that the State is not a totally stable area. These structural adjustments may be, in part, a cause of the warm artesian waters.

Another type of heating mechanism which may be the cause of the hot spots in South Dakota is that produced by convection cell currents existing in the mantle. Figure 14 shows the possible correlation between the stages of an orogenic cycle and a hypothetical convection current cycle. If convection currents were responsible for creating the energy that elevated the Black Hills area, the general fall-out of data becomes simple to explain. For instance, in stage 2 of figure 14, cooler temperatures exist below mountainous areas than in surrounding plains areas which is the reverse of what one would expect if high temperatures were evolved from structural deformation. Data in appendix I and plotted on the isogradient map indicate that the thermal gradients are indeed lower in the Black Hills area than in the plains area in south-central South Dakota. Stage 2 of figure 14 may also lend credence to the high geothermal gradient in northern Gregory County as being a true gradient and not a result of migrating hot water as discussed in the following paragraphs.

Closely related to the stages of orogenic cycle and the hypothetical convection current cycle illustrated by Bullard (see fig. 14, this report) is the idealized cross-sectional view of a geothermal reservoir (fig. 15, this report) illustrated by Schuster (1973). The illustration by Schuster shows a smaller convection cell with the currents flowing counter to the illustration in figure 14. This type of smaller convection cells may in practice be superimposed on a convection cycle such as illustrated in figure 14. Due to the masking of structures in South Dakota by glacial drift and Upper Cretaceous Shales, faults such as illustrated in figure 15 are not known to occur outside of the Black Hills area. However, it is possible



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that differences in geothermal gradients in small areas are clues to the discovery of such structures.

As previously mentioned it is surprising that the periphery of the Black Hills, an area of deformation, does not exhibit as high geothermal gradients as the area in the eastern south-central part of the State which in the opinion of most geologists has been a relatively stable area throughout geologic time. The preceding discussion of convection currents in the mantle appears a logical explanation of the high geothermal gradients in the eastern south-central part of the State.

It is apparent from Schoon (1971; fig. 16, this report) that the $7^{\circ}F$ isogradient is suspiciously near the area where the Madison Red River recharge to the Dakota Formation takes place. If the two do in fact coincide then it is logical to assume that the water from the Madison-Red River is heated by a thermal gradient of $3^{\circ}F/100$ feet a distance of 20 miles to the west of the recharge area and upon rising to the Dakota Formation increases the thermal gradient of the area that is characterized by the $7^{\circ}F$ thermal gradient.

The majority of geologic opinion holds that water movement is generally from west to east in sedimentary rocks in the southern half of South Dakota. Therefore, it is logical to search for a source of heat in a westward direction from the Kucera No. 1 Bartels Oil Test (SW SE section 23, Township 100 North, Range 77 West). From the isogradient map (fig. 1) of the State there is another large area with a $5^{\circ}F$ thermal gradient which surrounds the $6^{\circ}F$ and 7° F gradients and it logically follows that this area should be discussed further. This area is located in northwestern Tripp County and northeastern Mellette County For instance, at the Kucera No. 1 Bartels Oil Test (for location, see fig. 17) the thermal gradient is 3 1°F, approximately 28 miles to the east the thermal gradient is 7°F At the location of the Kucera No. 1 Bartels Oil Test the bottom hole temperature was 117°F at a depth of 2387 feet. If it is assumed the water from near the bottom of the test is transferred to the area of high thermal gradient with no heat loss, and using the average depth of 1100 feet in the wells located in the area of high thermal gradient, the new assumed thermal gradient would be 7°F/100 feet. If the foregoing assumptions are true then perhaps the effect of circulating water on the thermal gradient should also be added to Levorsen's list of heat sources that contribute to thermal gradients.

The weak point of the foregoing assumption is, why should the temperature remain relatively constant between the Kucera No. 1 Bartels Oil Test and the area of the high thermal gradient and suddenly cool from the area of the high thermal gradient in an eastward direction? It is possible the area of recharge of the Madison-Red River interval to the Dakota Formation is quite restricted areally and as this warmer water enters the Dakota Formation and fans out in all directions the effect of warm waters soon becomes negligible.

According to Heald (1930, p. 4) some students of earth temperatures feel that variations in these temperatures are in large part due to another means of circulating ground water. This school of thought believes that as the sediments pack down, water is squeezed out of the clays and shales and enters porous sandstones and limestones. The water thus released moves updip through these porous beds until it finally escapes where the beds crop out at the surface. Thus, waters deeply buried in a syncline would move updip and result in the temperatures in the updip areas being somewhat higher at a given depth below the surface of the ground than in the lower limbs of the syncline from which the water originated. Although this has been observed in Oklahoma and the regional picture seemed to support the conception, Heald noted that measurements of individual oil fields and individual wells seem to refute this theory.

Examination of the isogradient map of South Dakota tends to convince the observer that the theory is not at work in the State. For example, the axis of the Williston Basin passes through central Jackson and central Perkins Counties. If the theory were at work in the State the isogradient contours would be steeper on the west side of the axis of the Williston Basin than on the east side because the flanks of the basin are steeper on the west than on the east. Thus, the hotter fluids should migrate faster to the west. A cursory review of figure 1 shows the reverse to be true.

Unfortunately the data in the area of this high thermal gradient come from water wells which have not been logged. Thus, reliable data points upon which to base a structure are not numerous. About all that can be stated is that the possibility does exist. If a structure exists, the western end of the structure may be located in section 2, Township 4 South, Range 28 East. At this point, the sea level elevation of the top of the Minnelusa is -290 feet (see fig. 17). A well immediately to the south has the Minnelusa at -440 feet and another well approximately 7 miles to the northeast the Minnelusa is at 415 feet below sea level. However, there are two reasons against postulating structure on this data: (1) the Minnelusa has been quite severely eroded in this area, and (2) the structure, if present at all, is more subdued on the Precambrian surface.

In view of the aforementioned possibilities it is not possible with present data to identify the cause of the high geothermal gradients that exist in east and the second a star a

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Figure 15 Cross section of northern California's geothermal electrical generating plant. (from Schuster, 1973.)



Idealized cross-sectional view of a geothermal reservoir. Hot igneous rocks at depth supply heat to the water-filled reservoir rock above. The hot water is less dense and rises until the impermeable cap rock is reached. If fissures are present in the cap rock, part of the geothermal fluid may escape and form hot springs or fumaroles at the surface. When the water reaches the area of the cap rock, it begins to move outward, cool, and become more dense. The greater density or weight causes the cooler water to move downward and be recycled, along with recharge water that might enter the reservoir along faults or fractures. A power plant is shown, drawing steam and/or hot water from the upper part of the reservoir.



GEOTHERMAL

ENERGY

SEPTEMBER 1974

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Aquiclude

Skull Creek denotes base of Dakota Formation

Figure 16. Schematic drawing illustrating the proximity of recharge to the Dakota Formation to the area of high geothermal gradient shown in Figure 1. (after Schoon, 1971)

south-central South Dakota. Perhaps the theory involving convection currents originating at depth is most plausible.

Farther west, in the area of the Black Hills, the thermal gradient is consistently guite low. However, the reader is advised that the condition may be true only at comparatively shallow depths. From the isogradient map it is clear that there is virtually no data in the crystalline area of the Black Hills.¹ Also, the number of wells that penetrate to the Precambrian rocks in the periphery of the Hills are few. These two facts in conjunction with the belief of numerous geologists that outcrops of younger sedimentary rocks take on significant quantities of surface runoff may explain the lower than expected thermal gradient. If the conjectured recharge of warm water from the Madison-Red River interval in Gregory County is responsible for an increased thermal gradient in that area then, conversely, cool water in the form of surface runoff could lower the thermal gradient in the area of the Black Hills.

OTHER APPLICATIONS

In the search for oil, Levorsen (1967, p. 423) states that the crests of anticlines appear to have a small but measurable increase in thermal gradients in relation to the flanks of these same anticlines. Data points on known structures in South Dakota are not sufficient to confirm Levorsen's statements. The thermal gradients appear quite uniform in Jeeper portions of the Williston Basin. However, if Levorsen's statement holds true in South Dakota, undue attention should not be given to areas with contours at the expense of large areas with no contours. Because of the selection of the designated contour intervals it is quite possible that small areas within contour intervals exhibit greater variations than that suggested by wide spaced contour intervals. For this reason all data points in appendix I are plotted in areas that are considered to have oil potential.

Increased temperature lowers the viscosity (resistance to flow) of liquids. With this in mind the isogradient map may be instrumental in locating migration paths of oil.

In northwestern Haakon, northeastern Pennington, and southeastern Meade Counties, there is an area with a geothermal gradient in excess of 3° F/100 feet (fig. 1). Immediately to the west, in an updip direction, the geothermal gradient is 2° F or less per 100 feet. Although the variation is not great it is sufficient to give a 5000-foot oil test a temperature differential of 50° F (195° F - 145° F). From Levorsen (1964, fig. 5-33; fig. 18, this report) an accumulation of 32 API gravity oil subjected to a change in temperature from 90° F to 140° F could decrease in viscosity from 75 to 50 Saybolt Universal seconds.

California Edison And Geo-Energy Sign Steam-Purchase Pact

By a WALL STREET JOURNAL Staff Reporter ROSEMEAD, Calif. – Southern California Edison Co. said it signed a contract with Geo-Energy Systems Inc., Los Angeles, to purchase geothermal steam upon the successful completion of a test geothermal well program in California's Imperial Valley.

According to the agreement, the utility would build a 55,000-kilowatt generating facility at the well site upon successful sustained production of a specified minimum amount and quality of steam required for economical electricity generation. Depending upon additional sustained production of steam, the company would build subsequent generating units, it said.

Geo-Energy, under terms of the agreement, must complete a test well within seven months to verify the technology to be utilized and the geology of the Imperial Valley area selected-five miles south of the Salton Sea. Additional wells will be produced to verify adequate reserves, Southern California Edison said.

Geo-Energy will be using a system called the Van Huisen Downhole Heat Exchanger, a patented heat-transfer system that is designed to extract heat from geothermal reservoirs without bringing any of the underground materials to the surface, the utility said. Allen T. Van Huisen, who invented the process, is the director of research and development for Geo-Energy, a Southern California Edison spokesman said.

Southern California Edison said it couldn't estimate the price of the steam to be purchased or the cost of the generating facility at this time. Geo-Energy officials couldn't be reached for comment.

This will be the third geothermal project for Southern California Edison, which is seeking steam elsewhere in the Imperial Valley and in Northern California.

GEOTHERMAL REPORT

BI-WEEKLY NEWSLETTER DEALING WITH GEOTHERMAL

DEVELOPMENTS

RICHARD A. SMITH P.O. Box 35-K Tracey's Landing, Md. 20869 Given the right combination of viscosity, pressure, and permeability this temperature variation could conceivably halt migration. Because permeability is, affected inversely with the viscosity, and temperature increases result in a decrease in viscosity the possibility of geothermal trapping mechanisms are very apparent. Perhaps additional study may find that some so-called permeability traps are in reality thermal traps or viscosity traps.

CONCLUSION

The cause of high geothermal gradients in South Dakota is not fully understood. Dutcher et al. (1972, p. 25) gives criteria for distinguishing between conduction and convection by saying that a rule of thumb is, if the gradient is uniform, thus permitting an extrapolation to a reasonable mean surface temperature, heat flow is assigned to conduction. Whereas, if high temperatures with small gradients extrapolate to high surface temperatures, heat flow is attributed to convection. Whether this rule of thumb holds true in an area of extensively water flushed formations at depth is not known.

In the area of the $7^{\circ}F$ isogradient the Precambrian surface is approximately at sea level, or depending upon drilling site, at a depth of 1500 to 2000 feet. If a test hole were continued a few hundred feet into the Precambrian rocks a temperature log could be run to determine whether or not the high geothermal gradient continues below the Cretaceous sands. If the geothermal gradient decreases, one could assume a false geothermal gradient is caused by hot water moving into the area from another area. On the other hand, if the geothermal gradient is continuous, a geothermal reservoir suitable for generating power might be present at a depth of 7800 feet.

Burnham and Stewart (1971) report that an ideal power site should have rock temperatures above 600°F at depths ranging from 6,000 to 10,000 feet, large quantities of surface water and an area remote from population centers. Using these criteria as requirements for a power site, the nearby Lake Francis Case and the rural nature of the area satisfy two requirements and there is left only one questionable requirement. This is determination of a completely reliable measurement of the geothermal gradient.

One possible requisite not mentioned in the preceding paragraph is the existence of a permeable, interval at sufficient depth to insure a means to utilize high temperatures. Many wells drilled to and in Precambrian rock in South Dakota fail to yield water. However, if the change of the character of the contours on the magnetic and gravity maps of South Dakota represent the Peace River-Superior Provinces, it is quite possible that fractured Precambrian rocks are present in the area.

In view of the projected fuel shortage faced by our nation, the area of high geothermal gradients should be researched as a potential power supply. Along the same line more attention should be given by exploration geologists to the possible effects that subsurface changes in temperatures have on the migration and accumulation of oil.

Perhaps there are other areas in South Dakota with high geothermal gradients. These may in time be discovered by carrying out an extensive drilling program. However, a drilling program is a relatively expensive means of exploration. In south-central South Dakota including the area of Shannon, Washabaugh, Bennett, Mellette, and Todd Counties, there is a scarcity of data. The fact that all geothermal data points plotted in this area (with one exception) are higher than the world-wide average, suggests that this rather large area merits further study.

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Copies of Report of Investigations No. 110 may be obtained from South Dakota Geological Survey, Dept. of Natural Resource Development, Vermillion, South Dakota.



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LEGEND



Areas indicating geothermal gradients of less than 3° F per 100 Areas indicating geothermal gradients ranging from 3° F to 5° F Areas indicating geothermal gradiants in excess of 5° F per 100 Numbers adjacent to data points indicate geothermal gradient in Isogradient intervals at 2° , 3° , 4° , 5° , 6° , and 7° per 100 feet (For county names see back of this map.

Figure I. Isogradient map of South Dakota.



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GEOTHERMAL ENERGY DEVELOPMENT BY THE OGLALA SIOUX AT KYLE, SOUTH DAKOTA

> UNIVERSITY OF UTAN RESEARCH INSTITUTE EARTH SCIENCE LAB.

R. C. Eberhart R. A. Freeman

April, 1978

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APPENDIX B

A preliminary study has been completed which assesses the feasibility of developing the geothermal resources of the Madison formation for the community of Kyle, South Dakota, located on the Pine Ridge Indian Reservation. The key question considered is whether geothermal space heating can be costcompetitive with conventional space heating systems in the Kyle area. To answer this question, a cost model for a community geothermal space heating system was developed for Kyle, and used to compare costs of using geothermal energy with costs of using fuel oil and propane.

SUMMARY

The design analysis assumes that 125 homes in the Kyle area are converted so that each can use geothermal energy for space heating. The analysis also assumes that the new high school planned for Kyle is heated by geothermal energy.

The result of the analysis is that geothermal space heating does appear to be cost-competitive with conventional space heating systems. In fact, substantial monetary savings are projected, whether or not a reinjection well is required, provided only that the geothermal resource exists in a form similar to that which has been assumed.

Moreover, the installation of a geothermal space-heating system would result in other important benefits for the Oglala Sioux:

- 1) The redirection of money back into the reservation,
- 2) New jobs,

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- 3) A move toward energy self-sufficiency for the tribe, and
- 4) The availability of large additional amounts of energy (in addition to that used for space heating) for other purposes, such as greenhouses and fish hatcheries.

The analysis which has been performed has led to several recommendations:

 The potential economic and social advantages of developing geothermal energy resources in the Kyle area are extremely significant, and it is urged that this development be given serious consideration.

It is recommended that a feasibility study, similar to that carried out for this report, be completed for the Wanblee, S. D., area.

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- 3) It is recommended that the design of standardized heat exchangers, suitable for either solar or geothermal applications in BIA and HUD Native American housing throughout the United States, be pursued as soon as possible.
 - It is recommended that the utilization of low-grade heat from the discharge of the primary side of the central heat exchanger and the utilization of offpeak energy from the transmission/distribution system at Kyle be studied in detail.

5) It is recommended that, as a part of any effort to determine the magnitude and temperature of the Madison waters near Kyle for use as a geothermal resource, close attention be paid to the determination of water quality. If the water quality is sufficiently high, consideration can be given to using water from the geothermal well for drinking water and/or for irrigation.

I. INTRODUCTION

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A large percentage of the western half of South Dakota appears to be underlain by a geothermal energy resource. This resource, an aquifer known as the Madison Limestone Formation (the Madison) varies in temperature, water quality, depth and thickness from place to place.

The Madison contains vast amounts of hot water at temperatures between 120°F and 160°F (49°C and 71°C). Much of the water is potable and has been used by some communities and individuals for many years. A few isolated projects exist or are underway to employ the waters for space heating. These projects, however, will use only a small fraction of the energy resource.

The geology of the Madison is reasonably well known and this knowledge is being advanced further by studies now being carried out. Thus it can be stated that the waters of the Madison represent a major geothermal resource that is available, today, to the people of western South Dakota. Furthermore, no new technology is required to tap this energy source, and so it is possible to develop it in a relatively short time.

It is especially significant that approximately one-third of the land over the Madison Formation in South Dakota lies within the boundaries of Indian reservations. Although the exact areal extent of the formation is unknown, all or portions of up to six Sioux Indian Reservations lie over the Madison.

The lifestyles, beliefs, dependence upon high priced energy sources, and economic plight of the Indian people all contribute to make the development of geothermal resources on Indian land in South Dakota a unique opportunity. The close involvement of various federal agencies such as the BIA and EDA in tribal affairs, and the tribal governmental structure which exists, contribute to make the institutional mechanisms of geothermal development on Sioux lands unique. And the current disputes relative to sovereignty of Sioux Indian people, particularly with respect to water and mineral rights, serve to make the potential problems which surround the possible development of Madison Formation geothermal resources unique.

This report is intended to examine, in a preliminary way, some possibilities which exist for the utilization of geothermal resources by the Sioux people. The outline of a plan is presented for resource utilization at Kyle, South Dakota, a town on the Pine Ridge Reservation of the Oglala Sioux.

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The outline is not meant to be all inclusive; not all information needed to prepare a detailed plan for geothermal resource utilization at Kyle has been assembled. Likewise, this preliminary plan is not meant to be exclusive; Kyle is not the only location on Sioux land in South Dakota where geothermal development is possible.

The material in this report should be interpreted only as representing possible options for those who wish to consider geothermal energy for local applications. Since certain resource data are still being investigated, this report must be considered as preliminary. Comments and suggestions are actively solicited.

II. CHARACTERISTICS OF THE GEOTHERMAL RESOURCE

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The Madison is a well-known aquifer that underlies portions of South Dakota, North Dakota, Montana and Wyoming. In South Dakota it extends beneath the western half of the state. The entire aquifer covers an area of about 25,000 square miles and stores an estimated one billion acre-feet of water. The water temperatures are considered "moderate", ranging from about 120°F to 160°F (or 49°C to 71°C). Although its exact extent is unknown, it is believed that the Madison underlies most of the northern and western portions of the Pine Ridge Indian Reservation. The Black Hills, Big Horn and Laramie Mountains are believed to recharge the Madison aquifer at the rate of 150,000 acre-feet per year.

The estimated age of the Madison waters is between 15,000 and 30,000 years. They reside in a limestone formation averaging 400 feet in thickness, at depths ranging from 3,000 to 6,000 feet in South Dakota. The porosity of the formation averages 8 per cent; the transmissivity is estimated to be 0.013 $ft^2/$ sec and the storage coefficient is estimated to range from 0.0001 to 0.00025 (South Dakota School of Mines Report, "The Geothermal Applications on the Madison Aquifer System (Pahasapa) in South Dakota", 1976).

The ability of the aquifer to transmit water is quantitatively described by its transmissivity. Transmissivity is the rate at which water passes through a unit width of the aquifer under a unit hydraulic gradient (Lohman and others, 1972). The storage coefficient is a dimensionless number that is the ratio of the volume of water that an aquifer releases from or takes into storage per unit surface area of the aquifer per unit change in the hydraulic head (pressure) (Lohman and others, 1972).

The southern region of the Madison, including the area around the Pine Ridge Reservation, contains water that is generally potable and with dissolved solids of 1000 to 3000 ppm. To the north the water becomes more saline, less potable. Dissolved solids may be as high as 20,000 ppm near the border with North Dakota, and this presents problems in the direct use of the water.

Detailed discussions of the Madison's geology and its geothermal potential are available in a series of publications by the South Dakota Geologic Survey (of particular interest is the Report of Investigations No. 110) and by the South Dakota School of Mines and Technology in its report: Geothermal Potential of the Madison (Pakasapa) Limestone, 1976.

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Some information relative to the geothermal potential of the Madison in the vicinity of the Pine Ridge Reservation is contained in a listing of potential geothermal resources done by the Bureau of Indian Affairs (BIA) for The Department of Energy (DOE). Page 166 of the listing, which contains information relative to Pine Ridge, is attached as Appendix A. The information from Appendix A related to Pine Ridge is summarized in Table I.

TABLE I

PARTIAL LIST OF POTENTIAL GEOTHERMAL RESOURCES IN THE VICINITY OF THE PINE RIDGE INDIAN RESERVATION*

	Location		Temperature		Flow
Name	Longitude	Latitude	Type	(Degrees F)	(GPM)
Hot Spring	103°23.2'	43°33.7'	Spring	81	5000
Hot Brook	103°31.2'	43°33.7'	Spring	90	50
Cascade Spring	103°34.4'	43°23.4'	Spring	68	7200
Buffalo Gap Spring	103°12.0'	43°39.4'	Spring	?	?

Taken from a listing prepared in 1977 for DOE by the BIA

As can be seen from Table I, only four warm springs are listed. Since all potential resources within 60 miles of the reservation, including oil wells, are supposed to be included in the listing, it is obviously incomplete. For example, the towns of Midland and Edgemont, both within 60 miles of the reservation, currently use geothermal energy. In addition, four oil wells in the immediate vicinity of the reservation have yielded geothermal-related data. (These wells are discussed later in this report).

Although the Madison can be used for space heating and other applications in South Dakota much more widely than it is today, several important factors must be better known before extensive geothermal development can be shown practicable and economic. These factors, listed below, require the collection of data relating to the Madison, itself, and other aquifers that may interact with the Madison.

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1. The heat source and its life expectancy.

It is considered probable that the thermal properties of the Madison waters will remain relatively constant. However, the exact nature of the source of heat that raises the water's temperature above that found at equal depths in other regions of the United States in not known yet. If the source is a concentration of trace radioactive elements in the Precambrian granite which covers the Madison Formation, as has been suggested, the natural diminution of water temperature is expected to be very slow, measured in tens of thousands of years. The USGS/WR Madison Study, in progress, should resolve the question of heat source and better predict the expected life of the Madison as a geothermal resource.

2. Madison parameters

Physical factors that affect the ability to withdraw water and the drilling for water are well known in some areas, but not for the parts of the Madison which underlie the Pine Ridge Reservation. These factors include the aquifer's thickness, porosity, permeability, potentiometric head (artesian pressure), temperature, depth below ground level and the thickness of Precambrian base rocks that must be penetrated to reach the Madison's hot waters. USGS/WR work and planned work of the School of Mines should help to develop these data. Additional assistance is planned through the resource engineering program of the Division of Geothermal Energy of the Department of Energy.

3. Effects of substantial withdrawal

Extensive use of the water can affect the aquifer's thermal properties, quality of water and artesian pressure. In turn, such changes would affect the annual cost of a heating system by demanding higher flows, more rapid cleaning or replacement of components, and possible installation of additional pumps. Current USGS/WR work is developing a detailed and quantitative model of the hydrology in the Madison aquifer and other aquifers that can effect its flow. The area of particular interest is currently the Powder River Basin because of the Wyoming proposal to use Madison water to move coal via a slurry pipe line. DOE is negotiating to have the study extended so that proposed Madison usage in the South Dakota area would be included.

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Thermodynamic model of aquifer

A local (western South Dakota) model of the aquifer needs to be developed so that various rates of withdrawal and locations of withdrawal can be examined with respect to cooling the aquifer, well to well artesian effects, or the need to reinject water after use. The USGS model should be carefully examined to see if it would be adequate, or if it needs to be modified for purposes of utilization in the Pine Ridge Reservation area.

To estimate the total geothermal potential of the Madison for community space heating (or other purposes where the assumptions apply) in western South Dakota the following assumptions are considered: the average water temperature is 140°F, water is withdrawn at the natural recharge rate of 150,000 acre-feet/ year, and the load rejection temperature (water temperature at the end of the process) is 100°F. On this basis, the quantity of energy available is 16 x 10^{12} BTU, or the equivalent of 2.7 million barrels of oil, per year.

The above calculation, while providing a very rough estimate of total geothermal potential in western South Dakota, does not provide a particularly

useful estimate for the amount of geothermal energy potential on the Pine Ridge Reservation. Several factors contribute to make the situation on much of Pine Ridge significantly more attractive than much of the surrounding area of western South Dakota.

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One factor arises from the end uses to which the geothermal energy may be applied. An increased emphasis by the Sioux people on uses such as fish hatcheries, hydroponics and greenhouses would mean that the average load rejection temperature would probably be less than 100°F, resulting in more energy being recovered from a given volume of water. For example, the water temperature after using it for the space heating of homes, then for heating a greenhouse, might be 70°F, resulting in 240 BTU's per gallon more energy being recovered than if the load rejection temperature is 100°F, as above.

Another factor arises from an examination of geothermal gradient data for the Pine Ridge area. The geothermal gradient is defined as the increase of temperature per unit depth below the earth's surface. Its average value in the United States is about 1.4 or 1.5 degress Farenheit per 100 feet of depth. (That is, if a 1000-foot hole is drilled, the bottom-hole temperature will be, on the average, 14 or 15 degrees warmer than the temperature near the surface). Areas with geothermal gradients of $1.8 - 2.0^{\circ}$ F/100 ft. are often considered as areas of potential geothermal resource. Areas with a gradient of over 2.0° F/100 ft. are almost always considered as promising geothermal resource areas. (The Ocean City, Maryland, area with a gradient of 2.1° F/100 ft, is considered as one of the most promising areas in the Eastern United States).

A series of geothermal gradient maps for the United States was published in 1976 by the American Association of Petroleum Geologists. A portion of one of the maps, with the boundaries of the Pine Ridge Reservation and the location of Kyle and Wanblee added, is reproduced as Plate I. It is clearly evident from Plate I that the geothermal gradient for almost all (over 90 percent) of the reservation is over two degrees Farenheit per 100 feet. Even more significant is the fact that a large area in the north-eastern part of the reservation (and part of the Rosebud Reservation, to the east), including Kyle, apparently has a geothermal gradient of at least 2.4 degrees F per 100 feet.



In fact, as can be seen from the map from which Plate I was taken, this is the largest area in the southern two-thirds of South Dakota and the northern one-half of Nebraska with a gradient of 2.4 or higher.

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The existence of these relatively high geothermal gradients on the Pine Ridge Reservation is supported by data from four oil wells which have been drilled on, or in the immediate vicinity of, the reservation. The locations of these wells, with gradients of 2.47, 2.35, 2.30 and 2.18, are shown on Plate I.

Using a gradient range of 2.35 - 2.45 degrees F per 100 feet, and a depth range of 3500 - 3800 feet, the resulting water temperature range which might be expected in the Kyle area is 147 - 158 degrees F.

III. PROPOSED GEOTHERMAL APPLICATION OF THE MADISON AQUIFER AT KYLE

Sec. C.

INTRODUCTION

The sections above have defined the Madison aquifer as a geothermal resource that is already being used for space heating and which has the potential for much broader application in the Pine Ridge area.

In addition, geothermal energy is a clean source with few environmental disadvantages and, since it is located on the reservation, its use could reduce the problems that accompany the almost complete dependence on imported fuels: shortages, distribution and increasing prices.

Therefore, it is suggested that the Oglala Sioux, possibly in conjunction with assistance from federal agencies, consider developing local systems that tap the Madison and use the energy of the hot waters for heating homes and other buildings.

The key question in considering this proposed conversion is whether or not a community area can afford the costs. Unless the monthly costs (capital, amortization, maintenance, etc.) to a household (or other user) can be less than current and projected costs of conventional fuel systems, the household (or other user) cannot be expected to be interested in converting to geothermal.

To answer this question, a cost model for a community geothermal space heating system was developed for western South Dakota and used to show, in terms of town population, where the system is cost competitive with fossil fuels (Ref. 1). This model was then modified to more clearly reflect the conditions which exist on the Pine Ridge Reservation near Kyle, South Dakota, and used to make a similar cost comparison with fossil fuels for the Kyle community.

The design analysis which follows assumes that 125 homes in the Kyle community are converted so that each can utilize geothermal energy for space heating. Ninety of these houses are assumed to be currently designed to burn propane; 35 to burn oil. The analysis also assumes that the new high school planned for Kyle is heated by geothermal energy.

The system design parameters, as well as most of the design and cost calculations, are in Appendix B. The calculations are considered to be conservative. For example, it is assumed that the geothermal well yields water with a temperature of only 147°F, the lowest temperature expected; and that the well is drilled to a depth of 3800 feet, the greatest depth expected. ESTIMATED COSTS FOR OIL AND PROPANE

Current Average Heating Fuel Cost for Homes in Kyle, S. D.

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Assuming an average heating load of 1.725×10^8 Btu/year for a house in Kyle, the monthly average (over the entire year) is 14.4 x 10^6 Btu/mo (see Appendix B). Using a cost for propane of $\$.39/gallon (\$4.24/10^6$ Btu), and a cost for heating oil of $\$.45/gallon (\$3.26/10^6$ Btu), the average cost of heating a home in Kyle with propane is \$733 per year (\$61/mo.); with oil is \$563per year (\$47/mo.). Table II summarizes the estimated current cost of heating the 125 homes in Kyle which are being considered for conversion to geothermal energy. As can be seen, the total estimated average yearly fuel cost is \$85,675.

TABLE II

· · ·	Existing System					
· ·	One Home	No. of Homes	Total Yearly for Fuel Type			
Oil Heat	\$563	35	\$19,705			
Propane	\$733	90	\$65,970			
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ESTIMATED YEARLY RESIDENTIAL COSTS

Existing System

TOTAL \$85,675

Projected Heating Cost for New High School (011 Heat)

Making cost calculations (and comparisons) for a heating system for the new high school is difficult. For the purpose of this report, however, it is assumed that the cost of the system to transfer and distribute heat from a central location to individual rooms in the school is the same regardless of whether the heat is obtained from an oil-fired boiler or a geothermal resource.

The three factors which then enter into a cost comparison for the school

- 1) The cost of an oil-fired boiler for the school;
- The school's share of the cost of a geothermal well/ distribution system; and
- 3) The cost of fuel oil to fire an oil-fired boiler.

Assuming an average heating load of 13×10^9 Btu/year for the new high school, the monthly average (over the entire year) is 1.1×10^9 Btu/month (see Appendix B). Using a cost for heating oil of \$.41/gallon (\$2.97/10⁶ Btu), the average cost for fuel of heating the new high school would be \$38,610/year, or \$3217/month. (This price reflects the lower price of fuel oil to the school than the \$.45/gallon charged to homeowners). In addition, the amortization of the \$50K oil boiler results in a yearly cost of \$7,500, or \$625/month. The total cost, for purposes of this report, then, is estimated to be about \$46,110 per year or \$3,842 per month. (Note that the cost of the heat distribution system within the school has not been included, and is assumed to be the same as a distribution system for geothermal energy.)

DESCRIPTION OF GEOTHERMAL COMMUNITY HEATING SYSTEM FOR KYLE

Introduction

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The geothermal space heating system is divided, for cost modeling purposes, into five main components:

- (1) geothermal well,
- (2) central heat exchanger facility,
- (3) two-way transmission from central heat exchanger to community,
- (4) reinjection well and transmission from central heat exchanger to reinjection well, and
- (5) community distribution system (two-way pipeline, residence hookup and conversion of home heating system).

The Supply Well (and reinjection if required)

Based on current prices the cost of drilling, casing and enclosing a 7" diameter well drilled 3800' into the Madison formation was estimated by Francis-Meador-Gellhaus, Inc. to be about \$155,000. This figure is used to calculate total community costs for Kyle in the example below. Such a well is assumed to be artesian and have a flow rate of 933 gpm.

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The Central Heat Exchanger

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A central heat exchanger is used to transfer heat from the well to a secondary, closed-loop system containing treated water and thus, increase the overall system reliability. A stainless steel, plate-type heat exchanger with removable covers permitting ready access for periodic inspection and cleaning was selected. This type is considered to be extremely reliable and so a backup exchanger is not considered to be required.

To allow for growth it is assumed that the central exchanger for Kyle is designed with approximately ten percent excess capacity.

Town Distribution System for Kyle

The required cross-sectional area of the transmission pipe line is directly proportional to the population it supplies. Similarly, the cost per mile of installed pipe line varies linearly with the cross-sectional area of the pipe. This relationship is shown in Figure 1 for the closed-loop, twoway distribution line and in Figure 2 for the one-way line to a reinjection well, which might be considered desirable or required.

It follows, then, that pipe-line costs are directly proportional to the population served and costs per house can be determined.

The costs of transmission and distribution lines are based on the use of cement-asbestos pipe, on-site installation of 1" foam insulation for the feed line, burying the pipe 5' deep and compacting the refill soil only where the line crosses roads.

Individual Connections to Municipal System

Residential hook-up charges and conversion, including the local heat exchanger using the closed-loop heating fluid, are fixed costs per household.





COST MODEL CALCULATIONS FOR KYLE

Introduction

 Figure 3 illustrates the total costs for Kyle, S. D., with the production well one mile from town and one reinjection well located one mile from the production well. (Whether or not a reinjection well is required has not been determined.)

Energy Demands

The design requirements, both peak heating and average seasonal demands, were determined for weather conditions existing in western South Dakota and the type of housing in Kyle (see Appendix B). The peak heating demand was estimated to be 11.89×10^6 BTU/hr for 125 homes and the new high school. The distribution system consists of a six-inch main pipe and six three-inch feeder lines. Houses were assumed to be uniformly distributed through the town area for estimating the required pipe lengths. Homes were considered to be converted from a conventional forced-air system by installation of a heat exchanger.

COST MODEL RESULTS FOR KYLE

Using the estimated component costs, monthly costs for a geothermal heating system were calculated.

In these calculations it is assumed that the 125 homes to be converted have forced-air systems. (Conversion from oil or propane fired, forced-air furnaces, is relatively inexpensive since existing ducts can be used for hot air distribution.)

The conversion cost includes bringing the hot water into the house, the home heat exchanger, inserting the heat exchanger into the distribution ducting, thermostat and wiring for automatic control.

The monthly costs, assumed to be 15 percent per year, include capital amortization, maintenance, and services.

Table III, then, provides an answer to the question, "Can Kyle, South Dakota, consider geothermal space heating to be cost-competitive with conventional space heating systems?" The answer is yes, whether or not reinjection is required, providing only that the geothermal resource exists in a form similar to that which has been assumed.


Moreover, the installation of a geothermal space-heating system would result in other important benefits for the Oglala Sioux:

- 1) The redirection of money back into the reservation
- 2) New jobs

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- A move toward energy self-sufficiency for the tribe, and
- 4) The availability of large additional amounts of energy (in addition to that used for space heating) for other purposes, such as greenhouses and fish hatcheries.

TABLE III

SUMMARY OF ESTIMATED YEARLY COSTS

	125 Homes	New High School	TOTAL	Yearly Savings
Propane/011	\$85.7K	\$46.1K	\$131.8K	-0-
Geothermal with Reinjection	\$68.1K	\$34.8K	\$102 . 9K	\$28.9K
Geothermal with- out Reinjection	\$48.8K	\$23.0K	\$ 71.8K	\$60.0K

It is anticipated that the energy would be sold by a public utility type of organization which is owned by the tribe. Thus, revenues received (or at least a portion of them) will remain on the reservation. This is in marked contrast to the current cash flow off the reservation for fuel oil and propane.

It is anticipated that most (if not all) of the labor required to install the transmission and distribution systems, the hook-ups, and the heat exchangers, would be done by Oglala Sioux people. This would result in perhaps 10-15 jobs (rough estimate) being created for the period of time it takes to get the system on-line (perhaps two years). In addition, it is anticipated that perhaps 3-5 full-time long-term jobs will be created (for example, maintenance persons, and bookkeepers).

It is anticipated that a large amount of additional heat energy will be available at minimal cost for additional applications. The discharge temperature of the central heat exchanger is designed to be 107°F. Assuming that energy is extracted such that the water temperature is lowered from 107°F to 87°F, an additional 8.9 million BTU's per hour will be available from the well side of the central heat exchanger. In addition, during most days, the transmission and distribution systems will be working at considerably less than peak design capacity. Significant energy (several million BTU's per hour) would therefore be available for uses which might be considered to be "interruptable", i.e., could be disconnected from the geothermal source for a few days or a few weeks each year.

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Underscoring the job and cash flow advantages described above is the potential for saving significant amounts of oil and propane. Conversion to geothermal energy in Kyle could result in a yearly savings of about 138,000 gallons of oil (43,800 in homes; 94,200 in schools) and 169,000 gallons of propane.

In summary, the potential economic and social advantages of developing geothermal energy resources in the Kyle area are extremely significant, and it is urged that this development be given serious consideration.

IV. OTHER IMPLICATIONS OF GEOTHERMAL DEVELOPMENT AT KYLE

INTRODUCTION

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The positive nature of the results of this study suggest several other projects which appear quite promising. Four of the most promising are discussed below.

UTILIZATION OF GEOTHERMAL ENERGY AT WANBLEE

As can be seen from Plate I, Wanblee is situated almost in the center of the highest geothermal gradient contour on the Pine Ridge Reservation. It is possible that a geothermal gradient as high as 2.5°F per 100' exists, which would result in water temperatures of about 153 - 160°F at depths of 3500-3800 feet. The population of the Wanblee area (about 500) may be sufficient to support the development of geothermal resources. It is recommended that a feasibility study, similar to that carried out in this report, be completed for the Wanblee area.

DESIGN OF HEAT EXCHANGERS

The heat exchangers which are designed for use at Kyle may have a much wider potential application than just at Kyle. Indeed, the potential application may be nation-wide.

The heat exchangers are being designed for Native American housing built by HUD and the BIA, and are designed for an inlet water temperature of about 140°F. Thousands of BIA and HUD houses exist on Indian Reservations all over the United States; thousands more planned or under construction. The design inlet temperature of about 140°F is appropriate for many potential geothermal resources, and it is expected that a number of other tribes, mainly in the West and Southwest, could utilize the heat exchanger design in developing their geothermal resources.

Possibly even more significant, however, is the fact that most solarthermal flat plate collectors have outlet water temperatures of about 140°F. It would therefore seem possible to utilize the same basic heat exchanger design for all solar installations on BIA and HUD Native American Housing. It would seem that a relatively large industry, hopefully organized and run with Native American labor, could be established to manufacture these heat exchangers. It is recommended that the design of standardized heat exchangers be pursued as soon as possible.

UTILIZATION OF LOW-GRADE HEAT

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As was outlined earlier in this report, it is anticipated that a large amount of "low-grade" heat energy will be available at Kyle, at very minimal cost, for additional applications. The design discharge temperature of the central heat exchanger is 107°F.

Temperatures as low as 60°F are useful for fish hatcheries and greenhouses. Assuming only that energy is recovered such that the water temperature is lowered from 107°F to 87°F, an additional 8.9 million BTU's per hour will be available from the well side of the central heat exchanger; lowering the temperature to 77°F would make 13.4 million BTU's available.

In addition, during most days, the transmission and distribution systems will be working at considerably less than peak design capacity. Significant amounts of energy (several million BTU's per hour) would therefore be available for uses which might be considered to be "interruptable", i.e. could be disconnected from the geothermal source for a few days or a few weeks each year. Such interruptable customers could either supplement the geothermal source with conventional sources, such as electricity and fuel oil, or shut down completely during the coldest periods of the year. It is recommended that the utilization of low-grade heat from the discharge of the central heat exchanger and the utilization of off-peak energy from the transmission/distribution system be studied in detail.

UTILIZATION OF DISCHARGE WATER

Although the quality of the Madison water in the Kyle area is unknown, its quality in other communties in southwestern South Dakota is relatively high. This relatively high quality might make it possible to use the water from the geothermal well (before or after extraction of low-grade heat) for drinking water and/or for irrigation. The design flow of 933 gpm in the well would mean that up to 1.3

million gallons per day could be available for these (or other) uses. It is recommended that, as a part of any effort to determine the magnitude and temperature of the Madison waters near Kyle for use as a geothermal resource, close attention be paid to the determination of the water quality.

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V. FINANCING GEOTHERMAL PROJECTS

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As discussed, previously, the Madison geothermal resource is known to exist and known to be readily available in specific locations. However, for the broad utilization of the Madison, more precise reservoir and well data are needed in many areas. Department of Energy assistance in obtaining these data is available in PRDA's and PON's which are defined and discussed below.

In areas where the resource availability is proven, communities must obtain financing for the conversion project. Total project costs for conversion to geothermal energy space heating for 125 homes and a high school in the Kyle area, for example, would be approximately 500K-700K dollars. Various federal departments offer assistance in the grant, loan, and loan guarantee areas and some of the more promising possibilities are discussed below:

A. Department of Energy

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1. Grant and Cost Sharing Programs

The DOE provides grants and participates in cost sharing programs for geothermal projects in the private sector. This is done through two vehicles: The Program Research and Development Announcement (PRDA) and the Program Opportunity Notice (PON).

a. PRDA

Each such announcement solicits proposals for studies and analyses that will lead to new and improved technology for extracting and utilizing energy from geothermal resources. PRDA's are issued from the department's San Francisco Operations office and can provide total funding for approved projects or sharing of costs when a proposer could benefit independently from participation in a project. State, municipal or noncommercial applicants are chosen on a competititve basis.

b. PON

This type of notice solicits proposals for geothermal field experiments and applications that will demonstrate

adequacy of the reservoir as well as provide technical and economic data, and address legal, environmental and institutional issues for assessing the practicability of further resource usage. PON's are issued from the department's San Francisco Operations office. Applicants are selected competitively and projects are funded through federal and local cost sharing.

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Of twenty-two proposals submitted in reponse to a 1977 PON eight were selected, four in South Dakota. These were to the School of Mines and Technology for heating ranch buildings and agriculture uses; to the community of Box Elder for heating the Douglas School complex; to the Haakon School District for heating school buildings in Phillip; and to the St. Mary's Hospital in Pierre for space heating the hospital and neighboring business structures

2) Loan Guaranty Program

This program is intended to assist lenders in the private sector by guarantying them against loss of principal or interest on loans made for evaluating economic potential of geothermal reservoirs, for research and development in the technology of extracting and utilizing resources, for obtaining rights in resources, and for developing, constructing and operating geothermal energy producing facilities.

The San Francisco office is responsible for supplying information on the program and for analyzing guaranty applications from South Dakota. (Mailing address: Department of Energy, Geothermal Loan Guaranty Program Office, San Francisco Operations Office, 1333 Broadway, Oakland, California, 94612, Telephone: (415) 273-7151).

B. Department of Housing and Urban Development

Title 1 Community Development Block Grants are currently available in at least three areas, two of which appear to be applicable to the Kyle, South Dakota, area.

1. Small Cities Grants

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The first type of grant is the Small Cities Grant, which is a discretionary grant limited to non-SMSA areas. (An SMSA is a Standard Metropolitan Statistical Area, of which there are none on the Pine Ridge Reservation). These grants are awarded on a scoring system which includes factors such as the need for the project and the types of impacts on various types of people. It appears that the Kyle area is eligible for this type of grant.

2. Urban Development Action Grants

The second grant type is the Urban Development Action Grant, which is a one-year grant available to non-SMSA's. This type of grant is appropriate when private investment money is also available. Kyle would appear to be eligible for this grant type.

3. Entitlement Grant

The third grant type is the Entitlement Grant, which is available only to SMSA's. This is a formula grant, based upon factors such as population and income levels. Kyle does not appear to be eligible for this type of grant.

An application for a grant by the Oglala Sioux will, of course, be in competition with other applications within the HUD Region (Region 8 for South Dakota). A major consideration in awarding grants is the inclusion of low-incoming housing. The Oglala Sioux Housing Authority is currently involved in, and aware of, a number of HUD-sponsored programs.

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SYSTEM DESIGN PARAMETERS FOR SPACE HEATING OF OGLALA SIOUX HOMES AND NEW HIGH SCHOOL AT KYLE, SOUTH DAKOTA

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Average Volume per House (ft ³)	7700
Heat Loss Factor (Btu/hr, ft ³)	7.7
Based upon: 70°F inside design temp. -20°F outside design temp. Assume roof insulation	
Peak Heat Load per Home (Btu/hr)	5.93×10^4
Residential System Peak Design Size (Btu/hr)(125 homes)	7.41 \times 10 ⁶
Estimated Volume of New School (ft ³)	640,000
Heat Loss Factor (Btu/hr,ft ³)	7.0
Based upon: 70°F inside design temp. -20°F outside design temp. Assume roof and wall insulation	
New High School Peak Heat Load (&Peak Design Size) (Btu/hr)	4.48×10^6
Total Kyle System (125 homes & new high school) Peak Design Size (Btu/hr)	11.89 x 10 ⁶
Well Flow Rate (gpm)	933 gpm
Well Temperature (°F)	147°F
Well Depth (ft)	3800'
Assume 6" Main Dist. Pipe: 0.20 ft ² area	
If $v = 7$ ft/sec in pipe, then 1.4 cfs is flow	
84 cfm = 628 gpm = 5024 lb/min in central heat exchanger	output
Assume ∆t at home∜school of 137° - 93° = 44°F	
Then heat avail = 221,056 Btu/min = 13.26×10^6 Btu/hr. Av	AIL
And peak flow/home = 2.8 gpm	
In secondary of main heat exchanger:	
Assume 628 gpm, $\Delta t = 54^{\circ}F$; then heat = 16.28 x 10 ⁶ Btu/hr	· · · ·
In primary:	• • •
Assume 90% efficient main heat exchanger, then heat in = 17	.90 x 10 ⁶ Btu/hr
Assume Δt_{in} for main heat exchanger = 50°F, flow = 933 gpm	in well.

Assume main heat exchanger is sized for <u>full_load</u> (with respect to well and distribution system) to allow for future growth:

Service Service

Area (m²) =
$$\frac{\theta}{K\Delta t_m}$$

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where θ = peak heat load, Kcal/hr (0.252 Kcal/Btu)

 $k = 3050 \cdot Kcal/m^2, h, °C$

∆t_m= 8.4°C

$$\theta = (16.28 \times 10^6 \frac{BTU}{hr}) (0.252 \frac{Kcal}{BTU}) = 4.10 \times 10^6 \frac{Kca}{hr}$$

Area = $\frac{4.10 \times 10^6}{3050 \times 8.4}$ = 160m²

Cost = $(160m^2)(\$150/m^2) = \$24,000$ 7,000 15,000

\$46,000 TOTAL for central heat exchanger facility

Distribution, Hookup and Conversion Costs for Residential part of Kyle, S.D.

Installation Building

Distribution

7200 ft (1.70 mi) 3" (.05 ft²) pipe, 2-way, insulate 1-way = \$72K(~\$10/ft) 2500 ft (0.47 mi) 6" (0.2 ft²) pipe, 2-way, insulate 1 way = <u>42K</u> \$114K Total

Hookup

\$210 per house for hookup in Philadelphia District Heating with 50' connection length - Brookhaven National Lab. Study - Science v. 195, p951

Use: \$250 per house x 125 houses =\$31K

Conversion

(Assume forced-air heat exists now) Heating coil = \$270/home 8hr. installation @ \$10/hr = \$80

Heating coil and installation = $350/house \times 125$ houses = 44K

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AVERAGE ANNUAL HEATING DEMAND FOR HOUSES AND NEW HIGH SCHOOL IN KYLE, SOUTH DAKOTA

Heating season: 287 days (6888 hrs).

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Average heating season outside temperature: 32.8°F (use 32°F) Design outside temperature: -20°F

Average heating season inside temperature: 70°F

sing
$$H = \frac{24hD(T_1-T_a)}{(T_1-T_0)}$$
;
 (T_1-T_0) ;
 (T_1-T_0) ;
 $H = annual heating requirements
 $h = hourly design peak heat load
 $T_1 = inside temperature (70°F)$
 $T_a = average outside temperature (32°F)$
 $T_0 = outside design temperature (-20°F)$
 $D = number of days in heating season$$$

For each house, on the average,

$$H = \frac{24(5.93 \times 10^4) 287(70-32)}{70-(-20)}$$

= 1.725×10^8 BTU/year average for each house = 14.4 x 10⁶ BTU/month average for each house

For the new high school

H

$$= \frac{24(4.48 \times 10^{6}) \ 287(70-32)}{70 - (-20)}$$

= 13.0 x 10⁹ BTU/year (average)

= 1.1×10^9 BTU/month (average)

TOTAL GEOTHERMAL SYSTEM COST FOR KYLE, S.D. (without reinjection)

4

Well	\$155K
Central heat exchanger facility	46K
Double transmission pipeline	89K
Kyle distribution system	<u>114K</u>

\$404K for well and distribution system

Based upon relative sizes of peak heat loads $(7.41 \times 10^6 \text{ Btu/hr} \text{ for homes}, 4.48 \times 10^6 \text{ Btu/hr} \text{ for new high school})$ assign 62% of the \$404K to homes and 38% to the new high school.

SUBTOTAL

Residential System

\$250K	well/distribution
<u>75K</u>	hookup & conversion for 125 home
\$325K	TOTAL RESIDENTIAL COST
\$48.8K	15% yearly cost
\$390	per year per house (average)
\$ 33	per month per house (average)
1. A	

System for New High School

\$154K well/distribution

\$ 23K 15% yearly cost

\$1,920 per month (average) (does not include distribution system inside school, which is assumed to cost the same as one for oil heat). TOTAL GEOTHERMAL SYSTEM COST FOR KYLE, S.D. (with reinjection)

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\$155K
155K
52K
46K
89K
<u>114K</u>

SUBTOTAL

\$611K for wells and distribution/ reinjection system

Based upon relative sizes of peak heat loads $(7.41 \times 10^6 \text{ Btu/hr} \text{ for homes}, 4.48 \times 10^6 \text{ Btu/hr} \text{ for new high school})$ assign 62% of the \$611K to homes and 38% to the new high school.

Residential System

\$379K well/distribution/reinjection 75K hookup & conversion for 125 homes \$454K TOTAL RESIDENTIAL COST \$68.1K 15% yearly cost

\$545 per year per house (average)

\$ 45 per month per house (average)

System for New High School

\$232K well/distribution/reinjection

35

34.8K 15% yearly cost

\$2,902 per month (average) (does not include distribution system inside school, which is assumed to cost the same as one for oil heat) AREA SD Gthm RshPro

university of utah Research institute Earth Science lab. 3-21-7%

PROPOSAL FOR GEOTHERMAL RESEARCH

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Robert Schoon and Duncan McGregor undertook a study of the geothermal potential in South Dakota. This study culminated by publication of Report of Investigations No. 110 of the South Dakota Geological Survey. Since publication several inquiries have been made as to the validity of the high geothermal gradient (see pp. 17 to 25, and fig. 1 of Report of Investigations No. 110, enclosed). This question cannot be concretely answered using current data.

Thus, the primary objective of this proposal is to determine the reliability of the geothermal gradients as illustrated in figure 1 of Report of Investigations No. 110. In other words, does the geothermal gradient continue as high in the Precambrian rocks as in the overlying sedimentary rocks. If the gradient is continuous, temperatures equal to the boiling point of water could be encountered at the relatively shallow depth of 2,500 feet. Basically, this submittal requests financial aid in the drilling of a geothermal test well to a depth of 2,500 feet to determine if the geothermal gradient is true.

The foregoing constitutes stage 1 of this submittal and cost estimates include only those relative to stage 1. If stage 1 is successful, i.e., finds the geothermal gradient true as illustrated in figure 1 of Report of Investigations No. 110, the project will have been successful and completed at the end of stage 1 and the heating mechanism can be assumed as either due to convection currents within the earth's interior or to radioactive decay in Precambrian rocks.

Perchance the test well proves the gradient to be false, i.e., due to preheated, circulating ground water derived from the west, a stage 2 submittal will be made at a later date. This stage 2 proposal will have as its primary objective the location of the heat source that causes abnormally high temperature of ground water underlying an area of 14,000 square miles. From Report of Investigations No. 110, if the high geothermal gradient existing in Gregory, Lyman, and Tripp Counties is false, them the geothermal gradient is due to circulating ground water and the heat source is located basinward in southern Haakon, Jackson, and Washabaugh Counties. However, implementaticm of stage 2 is directly dependent on the findings of the stage 1 program.

A stage 3 of this proposal might have as its objective the drilling of a geothermal production well. However, this stage 3 is also premature because it hinges on the outcome of stages 1 and 2.

The projected implementation of this proposal is June 1, 1975. Drilling and testing procedures should be completed no liter than September 15, 1975. A completed technical report is envisioned as ready for final editing within four months of cesssation of drilling and testing programs.



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the entire area are of Recent age, but, in general, are composed of $\mathfrak{a}_{cathered}$ and reworked Pleistocene deposits. Their total volume probably is no more than one one-hundredth as great as that of the glacial deposits.

Hopkins, W.B.; Petere, L.R. 1963 - Geol Svevey - W.S.P. 15397

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The thickness of the unconsolidated Quaternary deposits ranges from a featheredge around the periphery of bedrock exposures to more than 256 feet in the Lords Lake buried channel a few miles southwest of Aberdeen. Available logs of holes drilled to bedrock indicate that the average thickness of the deposits is 76 feet.

About 30 percent of the wells shown on plate 1 tap the Quaternary deposits. However, not all these wells are sources of water supply, many being used only for observation of water-level fluctuations or tot used at all. Unlike the wells tapping the Dakota sandstone, these wells do not flow.

GROUND WATER

Except for the very few wells that tap Cretaceous rocks other than the Dakota sandstone, all wells in the Lake Dakota plain area either tap the Dakota or the glacial deposits. Because the two aquifers differ the so many respects, they are discussed separately in this report.

WATER IN THE DAKOTA SANDSTONE

The presence of an artesian aquifer beneath the James River lowand was known at least 10 years before South Dakota became a State. A well at Aberdeen, 1,066 feet deep and drilled in 1881, was among the first in South Dakota to tap this aquifer. Hundreds of deep wells and been drilled by 1920, and while still new most of them had a frong flow. The availability of such an ample supply of ground that without the need for pumping equipment prompted early writfrs to describe the James River lowland as a "Garden of Eden."

RECHARGE

Because the Dakota sandstone in this area is overlain by a thick excession of beds that are practically impervious to water, the places there water enters the formation obviously are outside the area. Also, because water in the Dakota in this area is under sufficient presare to flow from wells, the places where water enters the formation and be at a higher altitude. Hydrologists long have recognized that to water in the formation throughout the James River lowland is not local origin and generally have agreed that the formation is rearged principally where it crops out in the vicinity of the Black ills in western South Dakota. However, G. A. LaRocque and J. R. nes, of the Geological Survey, who made a hydrologic study of the unit as originally proposed, believe that only a small part of 654868-63-5

CONTRIBUTIONS TO THE HYDROLOGY OF THE UNITED STATE

the recharge enters the Dakota where it is exposed and much it was T28the recharge enters the Landau formations that underlie the Dakoa in most, of the recharge is from formations that a higher altitude most, or the recharge is from to rop out at a higher altitude. The the western part of the State but we phasapa limestone and Minnelus underlying formations, notably the Pahasapa limestone and Minnelus underlying formations, notatoly the charge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the Dakota free sandstone, receive much greater recharge than does the does the does than does the doe sandstone, receive much group in the Black Hills region (Brown, 1914) p. 1), and, because they thin eastward to a knife edge west of the p. 1), and, because any and the mercolates under pressure into James River lowland, one mane and Jones believe also that its Dakota may be recharged by downward movement of water in the Dakota may be recharged by the pressure in the Dakota is less day enough to raise the water in wells to the top of the zone of saturation in overlying rocks.

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MOVEMENT

According to Darton (1909, p. 60-61), the hydraulic gradient (ast therefore, the direction of percolation) of water in the Dakota is erally eastward throughout South Dakota. The rate of eastward percolation toward the James River lowland in T. 129 N. (the south ernmost east-west row of townships in North Dakota) was estimated by Meinzer and Hard (1925, p. 90) to be 400 to 500 gpm (gallons in minute), and this estimate probably is equally applicable to each est west row of townships in the Lake Dakota plain area. The average rate of discharge from the Dakota that could be maintained indefinitely without a regional diminution of hydrostatic pressure would be equal to the rate of eastward percolation across the west boundary of the

If the casing of a flowing well tapping the Dakota were to be enarea of withdrawal.

tended upward to a level that flow no longer occurred, the water level in the casing would coincide approximately with the piezometric, or pressure-indicating, surface of the water in the Dakota at that point, Because some wells tap only one water-bearing stratum and some up two, three, or possibly more, the water level in the extended casing A one well would not necessarily represent precisely the same piezone ric surface as the water level in another well. However, they prod ably would differ in such a small amount, all other factors being en at that in this report the water in the Dakota is considered to have only one piezometric surface. This imaginary surface is continuous with the water level in nonflowing wells that tap the Dakota outside the report area, provided the failure to flow is not due to plugging of U. well screen or escape of water from the well into rocks overlying its Dakota. Because maps showing the configuration of the piczometre surface, prepared by Erickson (1954, 1955), show that a trough 12 that surface coincides with the James River lowland and that the

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TEMPERATURE VARIATIONS OF DEEP FLOWING WELLS

By D. G. ADOLPHSON and E. F. LeROUX, Rapid City, S. Dak., Huron, S. Dak.

Work done in cooperation with the South Dakota Water Resources Commission

Abstract.—Measurements from about 200 deep artesian wells in South Dakota indicate that temperature differences in water flowing from wells of similar construction are related to the depth of wells and volume of discharge. Geothermal gradients at wells in the Dakota Sandstone east of the Missouri River range from 0.7° C per 100 feet in the southeast and 1.1° C per 100 feet in the northeast to 1.6° C per 100 feet along the Missouri River. Immediately west of the river, geothermal gradients average 1.5°C per 100 feet. In a "hot water belt" farther west, average geothermal gradients of 2.2°C per 100 feet may be due to deep high-temperature recharge to the Dakota Sandstone. Relatively low geothermal gradients in pre-Cretaceous rocks in the Black Hills may be due, in part, to rapid downward movement of recharging water in very porous formations.

Artesian aquifers tapped by thousands of flowing and nonflowing wells underlie much of South Dakota. Wells have been developed in both the pre-Cretaceous and Cretaceous Systems.

Temperatures of water flowing from about 200 wells of similar construction have been found to be related to the depth of the well and the volume of discharge. For large volumes of discharge, the temperature of the water discharging at the surface is very nearly that of the water in the producing formation. For wells of low flow, the temperature of the discharging water has been cooled during the relatively slow movement of water up the casing and is not as representative of formation temperature. For example, in 1960 a well in western South Dakota, drilled to a depth of 2,225 feet in the Minnelusa Sandstone, flowed 75 gallons per minute at a temperature of 39°C (Celsius). By 1962 the flow had decreased to 24 gpm and the temperature to 36°C. In 1965 the flow was 7 gpm and the temperature 32°C, and in 1967 the flow was 3 gpm and the temperature 27°C.

The southeastern, northeastern, and western areas of the State, shown on-figure-1, designate units-within_

which water-temperature variations with depth and volume of discharge from flowing wells are characteristically similar. Temperature plots indicate that temperatures for flows of less than 20 gpm are not representative of formation temperatures. Lines in the graphs, computed by the least-squares method, show that most of the plots for small flows fall between 13° and 17°C, regardless of the depth of the well, whereas for wells flowing more than 20 gpm there is an increase in the temperature of water with increase in depth of wells.

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SOUTHEASTERN AREA

Depth of well, temperature of water, and relative volume of discharge for 67 flowing wells that yield water from the Dakota Sandstone in the southeastern area are shown in figure 2. The measured temperatures of water discharged by flowing wells in the area ranged from 11°C for a well 200 feet deep, in Hutchinson County in the center of the area, to 24°C for a well





↓ ∧ U.S. GEOL, SURVEY PROF. PAPER 600-D, PAGES D60-D62



FIGURE 2.—Variations in water temperature with depth of well and volume of discharge for flowing wells in the Dakota Sandstone in southeastern South Dakota.

1,240 feet deep, in Buffalo County near the Missouri River. The average geothermal gradient in the southeastern area is 0.7°C per 100 feet.

NORTHEASTERN AREA

The 31 wells in the Dakota Sandstone represented by temperature and depth plots on figure 3 range in depth from 530 to 1,450 feet. Water temperatures range from 11° to 34°C. The water temperatures and geothermal gradients decrease eastward from the Missouri River. Temperatures as high as 34°C, and geothermal gradients of about 1.6°C per 100 feet are recorded for wells in Hughes and Buffalo Counties along the Missouri River. Farther east in Spink, Beadle, and Sanborn.

Counties, temperatures average about 20°C and geothermal gradients about 1.1°C per 100 feet.

WESTERN AREA

In the area immediately to the west of the Missouri River, known depths of wells flowing more than 20 gpm from the Dakota Sandstone range from 720 to 1,500 feet, and water temperatures range from 21° to 33°C (fig. 4, group *B*). Geothermal gradients at 22 wells average 1.5°C per 100 feet. Farther west there is a "hot water belt" (fig. 1) where well depths range from 1,180 to 1,830 feet and water temperatures from 36° to 54°C (fig. 4, group *A*). Geothermal gradients at 14 wells average 2.2°C per 100 ft. This "hot water belt" is in an area in which many of the deeper pre-Cretaceous formations are wedging out (Sandberg, 1962) and may be recharging the Dakota Sandstone (Hopkins and





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FIGURE 4.—Variations in water temperature with depth and volume of discharge for flowing wells in the Dakota Sandstone in western South Dakota. A, wells in "hot water belt;" B (shaded), other wells.

Petri, 1963) with high-temperature water. The municipal well at Philip, which is west of the "hot water belt," was drilled to a depth of 4,010 feet and produces 400 gpm from the pre-Cretaceous Madison Group at a temperature of 67° C.

BLACK HILLS AREA

Table 1 summarizes the water-temperature data for wells in the Black Hills area. The mean annual temperature in the area is about 8°C. Wells tap aquifers in the Pahasapa Limestone, Minnelusa Sandstone, Opeche Formation, Minnekahta Limestone, and the Spearfish Formation. Records are available for 42 wells ranging in depth from 300 feet near the outcrops to 4,900 feet 36 miles east of the outcrop area. Although the water temperatures range from 11° to 67°C, water temperatures of 26 wells ranging in depth from 300 to 1,300 feet are below 16°C.

The relatively low temperatures and geothermal gradients in the Black Hills area may be due to the

 TABLE 1.—Summary of water-temperature data for wells in the Black Hills area, South Dakota

Principal source	Well depth range (ft)	Temperature range (°C)	Average geothermal gradient (°C per 100 ft)	
Spearfish Formation	400-560	11-17	1. 3	
	360-680	11-22	. 9	
	640-1 310	12-16	7	
Minnelusa Formation	300-4, 900	11-60	. 7	
Pahasapa Limestone	460-4, 110	11-67	1. 0	
Average of 42 wells			0. 9	

rapid downward movement of recharging waters in very porous formations such as the Pahasapa Limestone. Or, it may be that, for many wells, a meaningful temperature gradient cannot be calculated because they are uncased holes which allow mixing of water from several aquifers (Cox, 1962).

CONCLUSIONS

Temperatures of water flowing from deep artesian wells are related to the depth of the well and the volume of discharge.

In South Dakota, temperatures and geothermal gradients at wells in the Dakota Sandstone generally decrease eastward from a "hot water belt" west of the Missouri River. The abnormally high geothermal gradients (Levorsen, 1958, p. 401) in the western area may be due to deep high-temperature recharge to the Dakota Sandstone. Throughout most of the eastern part of the State, where temperature gradients are more nearly normal, the Dakota Sandstone rests on crystalline basement rock and is not recharged from below.

The relatively low geothermal gradients computed for the Black Hills area may be due, in part, to rapid downward movement of recharging water in very porous formations.

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