

sea. Though perched water was a more persistent source, at times was not available. In historic time we have been in a period of relatively high precipitation. All of the swamp, lake, spring run, and sinkhole deposits that severe aridity existed throughout the peninsula during the last 16,000 years. There is no evidence of significant temperature intervals that in other temperate regions represent annual intervals. There is no evidence of significant temperature intervals on the peninsula during the last 16,000 years. The changes in the pattern and frequency of precipitation from hurricanes, thunder storms and frontal systems. Geographic and paleontological data from Silver Run, Ichatuckahatchee River, Devils Den Sinkhole, Warm Mineral Springs, Little Salt Springs and from lakes and peat bogs of Florida are reviewed in relation to establishing a chronology of hydroclimatic events.

TECTONICS AT A GREENSTONE-GNEISS BOUNDARY, LAKE WINNIPEG
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 On Manitowishong Island, L. Winnipeg, a "greenstone" felsic volcanic pile with mafic and intermediate rocks, is overlain by volcanogenic sandstones, arkosic sandstones and conglomerates, spinifex-textured gabbro bodies and cherts. The pile is overturned and dips steeply to the east. It is separated from a granite/gneiss terrane to the east, by a zone of quartz-diorite and tonalite gneiss (3.0 and 2.9 b.y.), which is overlain by granites (2.7 b.y.); metasediments, and intrusions related to the volcanic pile, are recognized. This zone trends NNW-SSE and is characterized by intense differential vertical movement, characterized by steeply plunging, upright, isoclinal folds, which invariably indicate that the granite/gneiss has moved upwards relative to the greenstone. Pervasive kink- and box-folds about steeply plunging folds and associated structures, overprint the isoclinal folds. The occurrence of one set of kinks, and a general right-lateral slip surface that compression on a NNE-SSW axis caused bend-glide folding of an earlier NNW-SSE foliation. The gneiss zone is, in part, a remobilized crystalline basement to the greenstones, at the margin of a wholly remobilized granite/gneiss terrane. The events here, the symmetry, lack of N-S trends, predominance of triple junctions, and abundance of triple junctions in the greenstone belts of the northwestern Superior Province suggest that the belts may be related to zones of interference between cells of Kenoran granite/gneiss through a general stress field whose major horizontal compressive axis is N-S to NNE-SSW. The hypothesis can be checked by the structural geology.

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METHOD FOR EXTRACTING ENERGY FROM "DRY" GEOTHERMAL RESERVOIRS
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 In response to the developing energy crisis, a small group at the Los Alamos Scientific Laboratory has been working for the past two years

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on the development of a practical, yet economical and environmentally acceptable, method of extracting thermal energy from the numerous regions of the earth's crust containing hot—but essentially dry—reservoirs at moderate depths. Temperatures in excess of 160°C at depths less than 6 km appear to be suitable for present development.

In this man-made system for extracting geothermal energy, the heat contained in an accessible body of hot rock would be convected to the surface by the flow of pressurized water, not unlike natural geothermal systems. As proposed, a large fractured region of hot rock created by hydraulic fracturing—would be interconnected to a heat exchanger (or heat engine) at the surface by a pair of drilled holes forming a closed convective circulation loop. The water in the earth loop would be maintained as a liquid throughout by pressurization at the surface, greatly increasing the amount of heat that can be removed for a given pair of holes, when compared to the flow of steam.

Preliminary experiments and analyses indicate that thermal stresses resulting from the cooling of the hot rock in such a reservoir may enlarge the initial fracture system so rapidly that its useful lifetime will be greatly extended beyond the planned 10 to 15 years provided by the original reservoir.

ORDERING OF TRANSITION METAL CATIONS IN OLIVINES

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A number of recent X-ray and Mössbauer studies of Fe-Mg olivines have shown that the larger cation, Fe²⁺, has a slight preference for the smaller and more distorted of the two nonequivalent octahedral sites, M1 ($K_D^* = 1.00-1.30$). In contrast, Fe-Mg pyroxenes exhibit strong ordering in the opposite sense with Fe²⁺ preferentially occupying the larger and more distorted M2 octahedral site ($K_D^* = 0.03-0.20$). In order to better understand the factors controlling site preferences of transition metal cations such as Fe(d⁶), Co(d⁷), and Ni(d⁸) in olivines and pyroxenes we have anisotropically refined the structure of Ni_{1.03}Mg_{0.97}SiO₄ (a=4.740, b=10.166, c=5.944Å) to R=0.029. Least squares site refinement indicates significant ordering of Ni on M1 (0.767 Ni + 0.233 Mg σ = ± 0.002) and of Mg on M2 (0.737 Mg + 0.263 Ni) ($K_D^* = 9.22$). This amount of Ni-Mg ordering is rather surprising in view of the small size difference (0.03Å) between the two cations and suggests that crystal field effects strongly favor Ni in the M1 site of olivine in agreement with the prediction of Burns (1970). Comparison of X-ray and optical spectral data for Fe-Mg and Ni-Mg olivines and pyroxenes indicates that site preferences of Fe, Co, Ni and Mg depend on a balance between (1) site distortion, (2) cation/site size and (3) the effect of M-O distance on CFSE. Only the site distortion criterion predicts the observed cation ordering in all cases. The fact that the oxygen ligands surrounding the M1 site of pyroxenes have a net negative charge, whereas those surrounding M2 are neutral on the average, is an additional effect which may enhance the ordering of transition metal cations on the M2 site of pyroxenes. $*K_D = [Mg(M2)Tr(M1)] / [Mg(M1)Tr(M2)]$

CONCLUSIONS AND COMPOSITION OF PYRITES FROM FISSION-ALTERED ROCKS OF THE BOULDER BATHOLITH, MONTANA
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Pyrites from throughout the Boulder batholith (including the Boulder district) have been studied microscopically. Micron-sized inclusions of sulfides, oxides, and silicates exhibiting a variety of textural relationships, with respect to the host pyrite. These inclusions occur either as single grains or as aggregates of two or more minerals. More inclusions are observed in fission-thermally altered rocks as opposed to unaltered rocks. The most common sulfide inclusions consist of chalcocopyrite or cubanite with pyrrhotite, and chalcopyrite plus or minus pyrrhotite. We believe these inclusions represent a phase captured during pyrite growth. The compositional field of cubanite solid solution is broad. It has been dropped, chalcocopyrite or sulfur-poor cubanite may be formed. Annealing experiments indicate that these inclusions are stable at about 325 °C.

About 500 probe analyses of pyrites and chalcopyrites made for Fe, S, Ni, Co, and Cu. Pyrite compositions vary between x = 1.88 to 2.11. Pyrites from fission-altered rocks tend to have more sulfur and less iron than unaltered rocks. The Ni, Co, and Cu contents of pyrites show average values: Ni - 0.15 weight % (range 0.08 to 0.22%), Co - 0.51% (range none to 5.7%), and Cu - 0.16% (range 0.08 to 0.22%). Chalcopyrite inclusions are slightly iron-poor and contain average values are: Ni - 0.14% (range 0.008 to 0.22%), Co - 0.1 to 1.3%.

⁸⁷Sr/⁸⁶Sr ANALYSES OF ULTRAMAFIC ROCKS, BASAL GNEISS, AND GRANULITES

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A zoned olivine-chlorite-tremolite vein within a granulite zone of the Norwegian Caledonides gives a chalcocopyrite age of 38 m.y., supporting geological evidence that granulites are formed by reaction of the ultramafic rock with siliceous solutions emanating from the country rock. The initial ⁸⁷Sr/⁸⁶Sr of the vein is 0.7093. Present day values of the actinolite (0.7077) and chlorite (0.7093) within the vein indicate the fluids contained the ambient ⁸⁷Sr/⁸⁶Sr ratio of the surrounding granulite. In contrast, minerals separated from granulites within similar ultramafic rocks yield the following values:

	Rb (ppm)	Sr (ppm)
Diopside	.008 to .046	23.6 to 112.
Actinolite	.002 to .005	0.13 to 0.
Chlorite	.0012 to .013	0.05 to 1.

⁸⁷Sr/⁸⁶Sr values of these minerals are well below the ⁸⁷Sr/⁸⁶Sr ratio of the olivine-chlorite-tremolite vein. Values of some of the diopsides are lower than those of the actinolites.