

not much older. The cave was self-sealed about 9,000 years ago, having been open only during the waning phase of the Pinedale Glaciation.

Original depositional layers of angular, pebble-sized roof debris should have been of nearly constant thickness on the nearly horizontal floor; disturbance of layers and periodic excavation of depressions indicate occupation by Man prior to 12,000 B.P. and at intervals during the following 3,000 years even though at times the cave was too low for erect posture. Corroborative evidence is supplied by the presence of artifacts and butchered bones. Included in the 44 species of mammals represented in the faunal remains are tundra-adapted Collared Lemmings, the world's oldest record of domesticated dogs, and six extinct species.

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GEOPHYSICS APPLIED TO THE SEARCH FOR GEOTHERMAL ENERGY RESOURCES

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The history of geophysical exploration for geothermal energy resources has been too brief to provide consistent guidelines for the explorationist. Considerable uncertainty exists as to the actual physical nature of geothermal reservoirs. It is, nevertheless, possible to search for geothermal fluids in an approximate way, geophysically, even in the absence of well defined conceptual models. Many techniques currently are being adopted, tested, or applied. Among these are: heat-flow measurements, thermal infrared imagery, direct-current and electromagnetic sounding and profile surveys, electrical self-potential surveys, geomagnetic variation soundings, airborne and ground magnetics, gravity, active seismic refraction and reflection surveys, microearthquake measurements, and seismic noise recording. Only those techniques related to present-day elevated temperature or steep thermal gradients are direct methods. Exploration on a crustal scale is now conducted by means of heat-flow measurements, geomagnetic variation and magnetotelluric soundings, and measurement of teleseismic P-wave delays.

Depending upon geologic circumstances, one may be looking for, (1) individual convecting hydrothermal cells or associated effects, such as intensive hydrothermal alteration; (2) bodies of magma (or their partially crystalline derivatives); or (3) permeable fracture systems along which hot fluids may be moving toward the surface from sources and depths unknown. Although hydrothermally altered rocks have somewhat predictable physical properties, their presence may or may not be significant, depending upon the age of the alteration. The search for magma is at present hampered by a serious lack of fundamental data on the physical properties of magma.

BROKEN OIDS, CEREBROID OIDS AND 'SPASTOLITHS': IMPLICATIONS FOR UNUSUAL SALINITIES IN THE MIDDLE CARBONATE INTERVAL, BELT SUPERGROUP (LATE PRECAMBRIAN), NORTHWESTERN MONTANA

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Syn depositional breakage of calcareous ooids is often at least 100 times greater in modern hypersaline than normal marine settings (Halley, 1974). This interpretation is further supported by the presence of abnormally high breakage and other synsedimentary deformation in several ancient oolite sequences closely associated with evaporites that range in age from Mississippian through Triassic.

Ooids from numerous beds within and among measured sections of the Middle Carbonate Interval of the Belt Supergroup display breakage frequencies from nearly 0% to over 25%. Accompanying breakage in some

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horizons are other types of penecontemporaneous cracks resembling "spastoliths" (Carozzi, 1963). Both may indicate appreciable expansion-recrystallization processes affecting the rock. Furthermore, Belt oolites are frequently interbedded with calcite- and sediment-filled hopper fabrics containing calcite- and sediment-filled hopper fabrics attributable to both primary disruption fabrics attributable to both dissolution of evaporite minerals. Throughout the section, tinctive nodular and laminar microspar-filled evaporites as calcitized evaporites.

Horizons displaying ooid breakage and/or oolite dissolution have been identified in several tens of kilometers. These features represent episodes of widespread elevated salinities during middle Belt times, and represent a salinization heretofore unrecognized within Belt carbonate horizons.

STRATIGRAPHY OF THE SOUTHERN BEAVERHEAD RANGE COUNTIES, IDAHO

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A 3,000-metre-thick incomplete sequence of Paleozoic rocks has been mapped in detail in the Beaverhead Range where it is unconformably underlain by Precambrian rocks. The formation is capped by a remnant of Triassic rocks. The formation, more than 600 m thick, is overlain by Ordovician Kinnikinic Quartzite. West of Coeur d'Alene Devonian Jefferson Formation unconformably overlies Kinnikinic. Unnamed Upper Devonian and Lower Devonian siltstones and limestones and the Upper Mississippian Middle Canyon Formation, which together with the Lower Canyon Formation, are overlain by a folded and faulted massive limestone of the Scurry Formation. More deformed Mississippian limestone and Scurry Canyon Formations overlie the Scurry and Scurry Mountain and range from 150 to 200 m thick. The most Mississippian Formation, the Big Snowy Formation, is interbedded shales and sandy and conglomeratic shales. The Big Snowy Formation is overlain conformably by Pennsylvanian upward into Pennsylvanian and Lower Permian limestone, and dolomite totaling about 975 m thick. The most 95 m of the Paleozoic sequence is interbedded limestone and shale of the Permian Phosphoria Formation overlain by a 36-m remnant of the Lower Triassic Flat-lying Tertiary volcanics and conglomerates overlain by older rocks.

GEOLOGY AND MINERALIZATION OF THE TUBUTAMA MOUNTAIN

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The Tubutama area is located 75 km south of and lies within the Sonoran Desert section of the physiographic province. The major topographic feature is a WNW trending fault block mountain. A Precambrian schist core is overlain by schists and phyllites of the Precambrian