

THE PETROGRAPHY, MINERAL CHEMISTRY AND ONE-ATMOSPHERE  
PHASE RELATIONS OF BASALTS FROM SITE 395 - LEG 45: DSDP

Michael A. Dungan,<sup>1</sup> Philip E. Long<sup>1</sup> and J. M. Rhodes<sup>2</sup>

<sup>1</sup>NRC-Resident Research Associate  
NASA Johnson Space Center  
Houston, Texas 77058

<sup>2</sup>Lockheed Electronics, Company, Inc.  
Houston, Texas 77058

## Abstract

Two distinct textural-chemical types of basalt are present in Holes 395 and 395A. Aphyric basalts are of two types; those with only olivine on the liquidus and those that are multiply saturated with plagioclase and olivine. Porphyritic basalts contain phenocrysts of plagioclase, olivine, spinel and clinopyroxene. There is a discrepancy between the compositions of phenocryst phases in the rocks and olivine and plagioclase produced in one atmosphere melting experiments. The disequilibrium mineralogy is suggestive of the mixing of primitive and evolved magmas, prior to eruption, in shallow magma chambers.

## I. Introduction

We present petrographic descriptions, microprobe data on the compositions of mineral phases and preliminary experimental results on samples recovered from Site 395 as part of a collaborative characterization of Leg 45 basalts. A companion paper (Rhodes et al., 1977) provides major and trace element analyses of the same suite of samples. Rhodes et al. (1977) have defined a chemical stratigraphy essentially equivalent to that deduced from shipboard investigations. Two distinct textural-chemical basalt types occur at Site 395.

Four units of basalt lacking macroscopic phenocrysts are referred to as aphyric basalts (Units A<sub>1</sub> and A<sub>2</sub>, Hole 395; A<sub>2</sub>, A<sub>3</sub> and A<sub>4</sub>, Hole 395A). These basalts do contain microphenocrysts of olivine and plagioclase and occasional larger crystals that are inferred to be xenocrysts. The aphyric units appear to consist almost entirely of pillow lavas. This is reflected petrographically in fine-grained textures, that typically range from glassy or spherulitic in pillow rims to pillow interiors with intersertal textures. Rare, more coarse-grained basalts with an intersertal-intergranular texture suggest small, local cooling units larger in scale than the average pillow.

The other group of basalts are porphyritic, containing abundant (10 - 30 percent) phenocrysts and microphenocrysts of plagioclase, olivine and clinopyroxene plus minor spinel microphenocrysts. The porphyritic basalts also occur in both holes as pillowed and massive basalt (Units P<sub>1</sub> and P<sub>2</sub>, Hole 395; Units P<sub>2</sub>, P<sub>3</sub>, P<sub>4</sub> and P<sub>5</sub>, Hole 395A) and as a sill near the bottom of Hole 395A (Unit A<sub>4</sub>'). Consequently, groundmass textures vary from glassy to ophitic.

Rhodes et al. (1977) have shown that these basalts have very similar trace element signatures (e.g., their chondrite normalized REE patterns are essentially parallel) indicative of origination from the same source. However, they also have found that simple crystal fractionation of the phenocryst mineralogy is not entirely sufficient to account for chemical differences among or within the various basalt types.

Our purpose is to relate the mineralogy and mineral chemistry of the various flow units to the bulk chemistry of the various units and the crystallization histories of the individual samples. Integration of these data with the results of experimental melting studies of two representative samples from the major basalt types provides some constraints on the fractionation of the units and their interrelationships. Non-equilibrium cooling rate experiments have been undertaken to determine if basalt textures indicative of variations in cooling history can be produced in the laboratory.

## II. Aphyric Basalts

### Petrography

Petrographically, the aphyric basalts include two subtypes that are most clearly distinguished in the rapidly cooled, spherulitic-textured variants. Basalts of Units A<sub>2</sub> and A<sub>4</sub> are characterized by abundant microphenocrysts of olivine. Plate 1a-d illustrates the range of textural variation in unit A<sub>2</sub> corresponding to decreasing cooling rate. The textural relations in Plate 1a are indicative of olivine crystallization prior to plagioclase and pyroxene. This interpretation is confirmed by melting relations determined for 395A-9-2-100-104 (see below) that indicate a narrow temperature interval below the liquidus in which olivine crystallizes alone. In contrast, the more evolved basalts of Unit A<sub>3</sub> were multiply saturated with olivine and plagioclase at eruption.

Evidence for this is found in glassy pillow rinds that contain both phases as microphenocrysts (Plate 1h). The presence of plagioclase on the liquidus of these rocks results in a different textural development. Whereas the basalts in A<sub>2</sub> and A<sub>4</sub> are characterized by radiating fan-spherulites of acicular plagioclase disposed around small glomerophyric aggregates of olivine, the tendency in A<sub>3</sub> is for the development of an intersertal texture (Plate 1f). These petrographic differences correspond to chemical characteristics. Basalts in Units A<sub>2</sub> and A<sub>4</sub> are less evolved (normative olivine = 9.4 - 12.0 %) than those in A<sub>3</sub> (normative olivine = 5.9 - 9.6 %).

We have analyzed the compositions of plagioclase and olivine (Tables 1 and 2) in six and eight of the aphyric basalts respectively. The range of olivine microphenocryst compositions in rapidly cooled rocks is quite narrow at Fo<sub>85-82</sub>. This corresponds to the liquidus olivine in the experimental study of 395A.9.2.100-104 which is Fo<sub>83</sub>. Olivine in the intergranular basalt 395A.64.CC is zoned from Fo<sub>85</sub> to Fo<sub>75</sub>. Laths and needles of plagioclase in rapidly cooled rocks are similarly restricted in composition (An<sub>66-59</sub>) and are also comparable to the experimentally produced plagioclase (An<sub>62</sub>).

Included within the basalts of Unit A<sub>3</sub> there are rare large crystals of olivine and plagioclase. A large anhedral grain in 395A.41.1.107.111 is considerably more magnesian (Fo<sub>88</sub>) than the associated microphenocrysts. A euhedral olivine grain that also includes a grain of chromite in 395A.60.1.66-70 is also anomalously magnesian (Fo<sub>87</sub>). Similarly, a plagioclase grain attached to the anhedral olivine grain in the core 41 sample is An<sub>86-79</sub> and the large anhedral grain figured in Plate 1 g and h (395A.56.3.63-65) is also more calcic (An<sub>83-79</sub>) than the microphenocrysts. On the basis of their more primitive compositions and the typical anhedral and embayed morphologies of these crystals they are inferred to be xenocrysts rather than phenocrysts.

### III. Porphyritic Basalts

#### Petrography

The porphyritic basalts recovered at Site 395 have clearly had, at minimum, two-stage cooling histories in that glassy pillow rinds contain large equant phenocrysts and microphenocrysts. Whereas the majority of porphyritic ocean floor basalts contain olivine and plagioclase phenocrysts ( $\pm$  spinel) on eruption, the porphyritic basalts at Site 395 were triply saturated with the silicate phases plagioclase, olivine and clinopyroxene. An important characteristic of this phenocryst assemblage from the petrographic standpoint is that there is a diversity of crystal morphologies, types of glomerophyric intergrowths and zoning patterns. Additional phases that crystallize at near-liquidus temperatures are sulfide and chromite. That sulfide is an early crystallizing phase is demonstrated by its inclusion in both plagioclase and spinel. Spinel occurs within plagioclase and olivine and in the groundmass but is never intergrown with clinopyroxene phenocrysts. These relations are consistent with one atmosphere phase relations determined for one of the porphyritic basalts (see below). The petrography and chemistry of the four phenocryst phases plagioclase, olivine, spinel and clinopyroxene are discussed separately in order of crystallization sequence.

#### Plagioclase

Plagioclase is the most abundant and most coarse-grained phenocryst phase in the Site 395 porphyritic basalts. We do not have quantitative modal data but a visual estimate for glassy basalts suggests approximately 10-20 percent plagioclase phenocrysts. Plagioclase phenocrysts range in size from large equant grains up to 5 mm in diameter to lath shaped microphenocrysts less than 0.1 mm. In general, as the size of the phenocrysts decreases they become less equant and increasingly elongate. Of the four phenocryst phases present, plagioclase exhibits the most diversity in morphology and zoning patterns.

Several distinct types of zoning patterns have been recognized and examples of some of these have been studied in reconnaissance with the microprobe. We have performed three element analyses (Ca-K-Na) that are corrected for deadtime, drift and background and converted directly to end-member feldspar ternary component.

Plagioclase phenocrysts in phyric Unit P<sub>1</sub> (395-cores 18 and 19; 395A-cores 13 through 15-section 4) are comprised of large equant grains (~3-5 mm) and laths (Plate 2a). Glomerophyric aggregates of plagioclase are common. Although these plagioclases contain chromite inclusions, olivine and clinopyroxene are not typically intergrown with plagioclase glomerocrysts in this unit. The cores of these phenocrysts are distinctive in that the majority of them are very weakly zoned. A weak normal gradient is present near the margin of the core and irregular patches of a slightly more sodic composition are present around concentrations of small melt inclusions (~5-10 μm). However, the typical compositional range of these cores is An<sub>86-83</sub>. The boundary between the cores and rim overgrowths is generally a sharp optical discontinuity that corresponds to a compositional gap or an extremely narrow zone with a very high gradient. The most calcic rim composition recorded for these phenocrysts is An<sub>73</sub> indicating a possible hiatus of ten mole percent An. The rims are normally zoned with narrow (~1-2 μm) oscillatory zones superimposed. The extent of zoning varies from narrow rims on phenocrysts in glass (An<sub>73-63</sub>) to much larger rims on phenocrysts in coarse-grained rocks. The compositional relations in lath plagioclase in these rocks warrant further investigation.

The majority of large plagioclase phenocrysts in the stratigraphically lower porphyritic basalts (core 15 - section 5 through core 33 and cores 61 through 64 - section 1) exhibit similar zoning patterns characterized by large equant cores and sharply zoned more sodic rims. However, the cores have much better developed oscillatory zoning and an apparently larger compositional range (An<sub>85-77</sub>) although more data is needed prior to rigorous comparisons.

Two additional distinctive petrographic types are present in subordinate amounts. These differ from the equant grains discussed above in that they are riddled with elongate melt inclusions. The more abundant of the two types occurs as large equant grains (3-5 mm) containing numerous regularly spaced cavities elongate in the direction of (010). Concentric normal zoning is developed around the melt inclusions. The core-rim boundaries are highly irregular (Plate 2b). We interpret these features as evidence of resorption, followed by reprecipitation. We have not analyzed these grains with the microprobe. The second type of inclusion-filled feldspars occur as laths rather than equant grains and consist of a core that is rimmed by an irregular vermicular zone which is in turn mantled by an inclusion free rim (Plate 2c). One grain of this type has been analyzed and was found to have a sodic core ( $An_{60-59}$ ) with a more calcic rim ( $An_{71-69}$ ).

It must be emphasized that this preliminary and generalized characterization of the plagioclase phenocrysts by no means describes all the textural and compositional variation in plagioclase of these rocks. However, these fragmentary data do establish the presence of several feldspar populations in porphyritic basalts at Site 395 that have had different crystallization histories. A comparison of compositions of the phenocrysts with the one-atmosphere liquidus plagioclase in 395A-27-2-111-116 shows a large discrepancy. Specifically the phenocrysts are considerably more anorthite-rich than the experimental plagioclase ( $An_{86}$  vs.  $An_{74}$ ). Thus, the composition of the liquidus plagioclase (in the experimental run) falls within the hiatus between the large phenocryst cores and the more sodic rims.



### Spinel

With a single exception, spinel grains  $>10 \mu\text{m}$  in basalts at Site 395 are restricted to the phyric basalts. The spinels occur as inclusions in plagioclase and olivine phenocrysts (never clinopyroxene) and in the groundmass. Spinel included in olivine and plagioclase are subhedral to anhedral whereas those in the groundmass are generally euhedral with sharp idiomorphic outlines. The latter typically contain melt inclusions. One grain of spinel included within a large melt inclusion in plagioclase in sample 395A.13.1.142-147 is an irregular, web-shaped grain distinct in composition from the general population.

We have analyzed spinels in 14 thin sections from Hole 395A. Most of the sections contain only two or three grains, commonly as inclusions in the same plagioclase phenocryst. Groundmass spinel is rare compared to inclusions. The data are presented on Figure 1 in terms of  $\text{Cr}/(\text{Cr}+\text{Al})$  vs.  $\text{Mg}/(\text{Mg}+\text{Fe})$  and permit several generalizations concerning composition vs. distribution. The spinels in the more evolved phyric Unit  $P_2$  plot with consistently low  $\text{Cr}/(\text{Cr}+\text{Al})$  compared to those in  $P_3$ ,  $P_4$  and  $P_5$ . They also exhibit a somewhat wider range in  $\text{Mg}/(\text{Mg}+\text{Fe})$ . The only analyzed groundmass spinel in  $P_2$  is the aforementioned anomalous web-shaped grain. Note that it has considerably higher  $\text{Cr}/(\text{Cr}+\text{Al})$  than the grains included in plagioclase or olivine. Within Units  $P_3$ ,  $P_4$  and  $P_5$  there is also a tendency for lower  $\text{Cr}/(\text{Cr}+\text{Al})$  in spinels that are included in other phenocrysts as opposed to groundmass spinel. A single grain of spinel is included in an olivine microphenocryst in the very sparsely phyric Unit  $A_3$  (395A.60.1.66-70).

Zoning trends in spinel were investigated by analyzing the cores and rims of several grains. Three zoning trends were recognized. Most of the spinel in contact with the groundmass is zoned to higher  $\text{Cr}/(\text{Cr}+\text{Al})$  with slightly decreasing  $\text{Mg}/(\text{Mg}+\text{Fe})$ . This trend mimics the overall trend of Cr-enrichment from inclusions to groundmass spinel. Some groundmass spinel grains have narrow rims zoned

186

towards titanomagnetite. Microprobe analyses at the grain boundaries are characterized by higher  $TiO_2$  and lower  $Mg/(Mg+Fe)$  with little change in  $Cr/Al$ . A third trend results from reequilibration between spinel and plagioclase. In these grains  $Cr/(Cr+Al)$  decreases with increasing  $Mg/(Mg+Fe)$ . Three spinel grains produced in experimental runs with 395A-27-2-111-116 at  $1200^\circ C$  are plotted in Figure 1 and listed in Table 3. They have compositions quite similar to the natural spinels.

The compositional range defined by the 395A spinels is essentially the same as that encompassed by spinels in oceanic tholeiites of a similar composition (Sigurdsson and Schilling, 1975; Dick, 1977). The aluminous spinels characteristic of the highly magnesian primitive tholeiites found in Legs 2, 3, 37 and in the FAMOUS area (Frey et al., 1974; Sigurdsson and Schilling, 1975; Dick, 1977) have not been found.

#### Olivine

The Site 395 porphyritic basalts contain approximately 5-10 percent olivine phenocrysts and microphenocrysts. As with the plagioclase phenocrysts, a variety of morphologies are present. In addition to the typical euhedral to subhedral, equant phenocrysts and microphenocrysts (3.0 - 0.1 mm) there are anhedral grains including some that are highly irregular. These tend to be amoeboid and in some cases are intergrown with plagioclase laths or equant grains as glomerocrysts. This texture is suggestive of resorption at some stage in the crystallization of the olivine. Skeletal olivine microphenocrysts are present in some rocks where they are typically intergrown with highly elongate plagioclase laths.

The results of olivine analyses in eight porphyritic basalts are presented in Table 2. The olivine phenocryst compositions range from  $Fo_{89}$  to  $Fo_{85}$ .

Microphenocrysts and the rims of grains in glass or in rapidly quenched groundmass are  $Fo_{85-82}$  with the majority at  $Fo_{84}$ . As in the aphyric basalts, iron-enrichment in the olivine with more slowly cooled groundmasses increases over the amount found in glassy rocks. The maximum Fo-content of phenocrysts in these rocks correlates weakly with the  $Mg'^1$  of the bulk rock. For samples with  $Mg/(Mg+Fe^*) = 0.59-0.64$  the most magnesian olivines are  $Fo_{87-88}$  whereas  $Fo_{89}$  olivines have been recorded in basalts with  $Mg' = 0.64-0.67$ . However, these olivines are substantially more magnesian than the olivine recorded in the one-atmosphere experimental study of 395A-27-2-111-116 ( $Fo_{85}$ ). They are also more magnesian than the values predicted by a  $K_D$  of 0.28 as determined by Roeder and Emslie (1970). Thus, there is a discrepancy between the natural and experimental olivine compositions analogous to that recognized for plagioclase.

#### Clinopyroxene

Many of the porphyritic basalts contained clinopyroxene phenocrysts on eruption. Large crystals are rare (i.e., >3 mm) and in general clinopyroxene phenocrysts range in size from 0.2 - 1.0 mm and one smaller than the large associated olivine and plagioclase phenocrysts. Although plagioclase and olivine phenocrysts are abundant (15-20%) in all the porphyritic basalts, the distribution of clinopyroxene phenocrysts is more variable in that they are altogether absent from some of the thin sections examined. The presence or absence of clinopyroxene phenocrysts does not appear to be correlated with bulk chemistry as they are present in all five of the chemically defined porphyritic units and chemical differences among these units are not large in any case.

The clinopyroxene phenocrysts occur attached to plagioclase laths in glomerocrystic aggregates or intergrown in the characteristic bow tie texture interpreted as indicative of coprecipitation (Bryan, 1972). Most of the grains are euhedral, or where complex intergrowths have formed, subhedral. Irregular extinction, a common feature in ocean floor basalt pyroxene phenocrysts (Bryan,

<sup>1</sup> $Mg' = Mg/(Mg+Fe^*)$ ;  $Fe^* = Fe_{total} \times 0.9$ .

1972) and complex zoning patterns (sector zoning and concentric oscillatory zoning superimposed on normal zoning) are present (Plate 2f).

Pyroxene Chemistry

The chemical compositions of phenocrysts and groundmass clinopyroxene in five porphyritic basalts have been investigated in reconnaissance with the electron microprobe (Figures 2, 3 and 4). Spot analyses were taken along linear traverses from core to rim or rim to rim in order to characterize the gross zoning trends and the chemical differences between phenocrystal and groundmass grains. As noted previously by Bence et al. (1976) in a study of pyroxene phyric basalts from Leg 15, the distinction between phenocrysts and groundmass pyroxenes is best made on the basis of chromium-content. Analyses of large phenocrysts in fine-grained matrices indicate a very slight iron-enrichment and calcium-depletion trend from core to rim typical of ocean floor basalt clinopyroxene phenocrysts (e.g.,  $Wo_{44} En_{48} Fs_7$  to  $Wo_{40} En_{48} Fs_{12}$ ). However, over this interval the typical chromium content of the core (1.5-1.1 wt%  $Cr_2O_3$ ) is decreased by a factor of 1/2-1/3 at the rim, alumina undergoes a similar depletion from 3.5-4.5 %  $Al_2O_3$  to 1.5-3.0 wt% and in general titania increases although the trend is diffuse and subject to oscillations. In rocks that contain abundant groundmass pyroxene, the outer rims of the phenocrysts are strongly depleted in  $Cr_2O_3$  (<0.10 wt%) relative to their cores even though the relative iron-enrichment is still meager (e.g.,  $Wo_{40} En_{44} Fs_{16}$ ). The marked tendency for Cr-depletion with progressive crystallization is illustrated in Figure 3. The strong partitioning of chromium into clinopyroxene relative to liquid explains the sharp depletion trend within single grains once closed system crystallization ensues (i.e., subsequent to eruption) clinopyroxene crystallization rapidly depletes residual liquid in chromium.

The chemical zoning in groundmass grains also exhibits iron and titanium-enrichment accompanied by decreases in calcium, aluminum and chromium. Some of the larger groundmass grains are cored by relatively high-chromium compositions indicating that the very marked Cr-decrease characteristic of phenocryst rims and groundmass grains was caused by early growth of abundant grains formed near the surface or after eruption and not entirely by a few large phenocrysts formed at depth. The iron-enrichment trend recorded in groundmass pyroxenes is typical of ocean floor basalts (Bunch and LaBorde, 1976; Bence et al., 1976) although the overall degree of enrichment (and concomitant Ca-decrease) is not as extreme as has been observed in other suites. The most iron-rich rim analyzed thus far in these rocks is  $Wo_{37} En_{37} Fs_{26}$ .

Based on the analyses presently available, it would appear that there is a correlation between the composition of clinopyroxene cores and the bulk composition of the rocks in which they occur. This correlation is most marked in phenocrysts in 395A-15-4-10-15, from the uppermost, relatively evolved porphyritic unit (A<sub>2</sub>-395A). The core of these pyroxenes have the highest Fe/(Fe+Mg) and TiO<sub>2</sub> as well as the lowest Cr<sub>2</sub>O<sub>3</sub> of any of the phenocrysts analyzed which coincides with the bulk rock data indicating higher Fe/(Fe+Mg), Cr<sub>2</sub>O<sub>3</sub> and TiO<sub>2</sub> than the underlying units. These relationships are less clearcut for the remaining four samples although they hold in a general sense. Additional analyses may more clearly establish if pyroxene phenocryst compositions are sensitive to such small bulk compositional differences in their host lavas.

#### IV. Experimental Data

In order to place constraints on possible crystal fractionation histories of the basaltic rocks here studied we have selected two representative samples for determination of their melting relationships and mineral compositions at one

atmosphere total pressure. In addition, we have performed controlled cooling rate experiments in an attempt to quantify the interpretations of textures observed in the natural rocks. Although the results obtained to date are preliminary, they have been remarkably successful and allow us to draw several significant conclusions regarding the evolution and emplacement of the basalts drilled on Leg 45.

### Experimental Techniques

Both the melting and cooling rate experiments were performed using the technique described in detail by Donaldson and others (1975) in which samples are suspended on a Pt wire loop in a vertical tube furnace. Oxygen fugacity in the furnace was controlled by mixing H and CO<sub>2</sub> and was measured using a modified Sato cell calibrated against the Ni-NiO buffer curve. In all the melting experiments f<sub>O<sub>2</sub></sub> was maintained at 0.5 (± 0.1) log unit below the QFM buffer curve. The temperatures of the furnace are believed to be accurate within 3°C and calibrations were made periodically against the melting point of Au. Approximately 100 mg of finely ground rock powder was used for each run. For the melting runs the sample was held at temperature for 24 to 72 hours, quenched and then removed from the furnace. The cooling-rate runs were held at 1250°C for 24 hours then cooled to 1000°C at rates of 2, 10 and 30°C per hour.

The experimental conditions just described do not duplicate the conditions on the sea floor that obtained at eruption of the Leg 45 basalts; i.e., there was roughly 0.5 kb of hydrostatic pressure on the rocks at that time. Thus the conclusions we draw from these experiments are based on the assumptions that a small pressure difference and the presence of sea water in the natural environment does not significantly alter the phase relations, textures, or mineral chemistry relative to the experimental system. We feel these assumptions are valid based on available high pressure data for basaltic system (see

Fujii et al., 1977, this volume). The specific  $f_{O_2}$  values were chosen on the basis of a compilation of estimated T versus  $f_{O_2}$  data for deep sea basalts by Haggerty (1976).

### Experimental Results

The results of the melting experiments are presented in Figure 5 and the compositions of plagioclase and olivine are tabulated as a function of temperature in Table 5. The representative aphyric sample, 395A.9.2.100-104, has olivine ( $Fo_{83.1} - 82.9$ ) as the liquidus phase at about 1205°C. Between 1200° and 1175°C olivine ( $Fo_{81.5} - 79.0$ ) is joined by plagioclase of the composition  $An_{66.5}$  to  $An_{61.5}$  (at 1175°C). Clinopyroxene appears at about 1160°C, and at that temperature the plagioclase composition varies from  $An_{61.6} - 64.7$  whereas the olivine composition varies from  $Fo_{77.9} - 74.5$ .

A sample from the phyric unit (395A.27.2.111-116) chosen for its relatively high Mg/(Mg+Fe) has plagioclase as the liquidus phase at about 1235° ( $An_{75.9} - 75.4$  at 1225°C). Olivine begins to precipitate at about 1215°C. At 1200°C the olivine composition is  $Fo_{86.1} - 84.7$  and the plagioclase composition is  $An_{75.9} - 72.4$ . In addition, a very minor amount of spinel is present at 1200°C. The compositions of three grains are plotted on Figure 1 and one analysis is included in Table 3. The Cr/(Cr+Al) and Mg/(Mg+Fe) of these spinels agree very well with those occurring naturally in sample 395A.27.2.111-116 suggesting that the experiments have successfully reproduced the oxygen fugacities that existed during the crystallization of the Leg 45 basalts. That the spinels are not simply relict grains is demonstrated in a similar set of experiments on Leg 46 basalts in which the spinel composition varied consistently over a range of temperatures (increasingly greater Cr/(Cr+Al) with lower temperature) indicating

that the spinels do in fact readily equilibrate with a silicate melt under the conditions at which the experiments were conducted.

At 1175°C plagioclase ( $An_{69.8-68.4}$ ) and olivine ( $Fo_{83.5 - 84.6}$ ) are joined by clinopyroxene but no spinel was observed. At 1160°C clinopyroxene coexists with plagioclase ( $An_{72.4 - 65.4}$ ), olivine ( $Fo_{79.1 - 76.3}$ ) and silicate liquid plus very small, highly reflective grains (either spinel or titanomagnetite). At 1150°C only a small amount liquid (10-20%) coexists with olivine, plagioclase and clinopyroxene.

#### Cooling Rate Experiments

The linear cooling rate experiments for the sample from the aphyric unit (395A 9 2 100-104) produced textures ranging from spherulitic (30°/hr cooling rate) to slightly porphyritic (2°/hr cooling rate). At the 30°/hr cooling rate, spherules of olivine dendrites are set in residual glass which makes up the bulk of the charge. In the 10°/hr run olivine and plagioclase occur as oriented arrays of acicular crystals. In parts of the charge, olivine occurs as skeletal crystals which clearly nucleated early in the crystallization history, whereas in others olivine appears to be interstitial to plagioclase laths. The charge also contains small tabular grains of titanomagnetite and euhedral grains of clinopyroxene(?). A cooling rate of 2°/hr produced large skeletal olivines which are set in a fan-spherulitic groundmass of plagioclase, olivine, clinopyroxene and titanomagnetite. A typical example of the large olivine grains is shown in Plate 1e. Olivine also occurs in chain-like morphologies similar to those shown in Figure 2 of Donaldson (1976).

Textures produced by controlled cooling rates of samples from the aphyric unit (395A-27-2-111-116) differ markedly from those of the aphyric unit. A 30°/hr



cooling rate produced radiating plagioclase needles in a very fine grained spherulitic matrix, whereas a 10°/hr cooling rate produced skeletal plagioclase laths in a matrix of plagioclase olivine and clinopyroxene. At a 2°/hr cooling rate plagioclase laths and subhedral olivine are set in a matrix of olivine, plagioclase and spherulitic clinopyroxene.

Discussion of Cooling Rate Experiments

The results of the controlled cooling-rate runs place some constraints on the crystallization history of the aphyric unit. In general, olivine morphology gives some indication as to the relative degree of undercooling at the time of crystallization (Donaldson, 1976). These experiments confirm this in that different cooling rates produce different degrees of undercooling which in turn could be correlated systematically with olivine morphology. Olivine morphologies in various parts of the aphyric unit then, can be used to make some broad inferences about cooling rates. Specifically, the olivine morphology of the 2°/hr run is very similar to the larger crystals in the glassy rinds and in the interior parts of pillows in the aphyric unit. The range of morphologies within the aphyric unit, however, covers the same range of morphologies seen in going from the 2°/hr run to the 30°/hr run. The similarity between the late dendritic crystals in the aphyric unit pillow rinds and those grown in the 30°/hr run is remarkable and suggests that a similar undercooling occurred at the later stages of crystallization in the pillow rinds as was established in the 30°/hr cooling run. That there are also early, more equant olivine crystals in the pillow rinds suggests that crystallization of the pillows took place at increasingly greater degrees of undercooling. A similar conclusion was arrived at by Kirkpatrick (1976) on a calculational basis although it is not clear whether the increasing undercooling stems from an externally imposed thermal

regime or whether it is inherent in the crystallization behavior of the melt itself. The progressive change in texture from the pillow rim inward does suggest, however, the general importance of the imposed thermal regime and that the effective average relative undercooling was substantially greater at the pillow rind than in the pillow interior.

Based on these experiments, cooling rates ranging from 2 to 30°/hr can maintain the respective undercoolings necessary to produce the range of textures observed in the aphyric unit. It is important to note, however, that the experiments on the aphyric unit sample (395A-9-2-100-104) were conducted under conditions that involved considerable superheat prior to beginning the controlled drop and this may effect the resulting textures. Additional experiments are planned which will involve minimal superheat, but at the present time caution needs to be exercised in attempting to make direct comparisons between the cooling rates of our experiments and those which may have produced the textures observed in natural basalt samples.

The cooling rate experiments for the phyric basalt are less easily interpreted because it is clear that the natural magmas were charged with crystals at the time of eruption whereas the experimental charges were crystallized continuously from an entirely melted sample. Thus the effect of having a large number of crystals present prior to onset of relatively rapid cooling has not been realistically modeled in these experiments.

## Summary and Conclusions

Two distinct textural-chemical types of basalts are present in Holes 395 and 395A. Aphyric basalts are predominantly relatively evolved compositions that occur almost entirely in pillowed units ( $A_2$ ,  $A_3$  and  $A_4$ ). Unit  $A_2$  is the least chemically evolved ( $\approx 10$  percent normative olivine) and has olivine alone on the liquidus. This conclusion is derived from the presence of olivine microphenocrysts in very rapidly chilled samples and one atmosphere melting experiments that indicate liquidus olivine at  $\approx 1200^\circ\text{C}$  followed by plagioclase at  $\approx 1175^\circ\text{C}$ . The compositions of olivine and plagioclase in the experimental charges agree within one percent of those determined in the rocks. Aphyric units  $A_4$  and  $A_3$  are progressively more evolved than  $A_2$  and have microphenocrysts of both olivine and plagioclase suggesting multiple saturation. Those compositions are quite similar with the exception of xenocrysts of olivine and plagioclase in Unit  $A_3$  that are more forsteritic and anorthitic respectively than the microphenocrysts.

The porphyritic basalts are far more complex. Integration of bulk chemical data from Rhodes et al. (1977) with the petrographic, phase chemical and experimental studies in this paper suggest that many of the olivine and plagioclase phenocrysts present in these did not crystallize from melts with the composition of the rocks but from a more primitive basalt. Specifically many of these phenocrysts are either more forsteritic or anorthitic than the experimentally determined olivine and plagioclase in a relatively magnesian porphyritic basalt. In addition, the various types of zoning patterns (including reversed) recognized in plagioclase phenocrysts within individual samples indicates that a mixed population is present. In all the porphyritic basalts from which we have microprobe analyses of olivine, several of the phenocrysts are more magnesian than the composition predicted by the work of Roeder and Emslie (1970). Spinel and clinopyroxene follow plagioclase and olivine in the crystallization sequence of the porphyritic basalts. Spinel crystallization is terminated with the

incoming of clinopyroxene. As the spinel and clinopyroxene compositions appear to correlate with the bulk chemistry of the samples in which they occur, they are probably not exotic to these rocks.

Presently we are examining the possibility that mixing of evolved magmas with more magnesian primitive magmas can account for the major and minor element characteristics of these lavas as well as their disequilibrium mineralogy. Preliminary data on basalts from Leg 46 indicates many of the same mineral chemistry anomalies and bulk chemical characteristics.

References:

- Bence, A. E., Papike, J. J. and Ayuso, R. A. (1976) Petrology of submarine basalts from the central Caribbean: DSDP leg 15. J. Geophys. Res., v. 80, p. 4775.
- Bryan, W. B. (1972) Morphology of quench crystals in submarine basalts. J. Geophys. Res., v. 77, p. 5812-5819.
- Bunch, T. E. and LaBorde, R. (1976) Mineralogy and compositions of selected basalts from DSDP Leg 34. In: Initial Reports of the Deep Sea Drilling Project, v. 34, p. 263-275.
- Dick, H. S. B. (1977) The geochemistry of chrome spinel in FAMOUS basalts and a model for magma genesis at an ocean ridge. Contrib. Mineral. Petrol. (in press).
- Donaldson, C. H., Usselman, T. M., Williams, R. J. and Lofgren, G. E. (1975) Experimental modeling of the cooling history of Apollo 12 olivine basalts. Proc. Lunar Sci. Conf. 6th, p. 843-869.
- Donaldson, C. H. (1976) An experimental investigation of olivine morphology. Contrib. Mineral. Petrol., v. 57, p. 187-213.
- Frey, F. A., Bryan, W. B. and Thompson, G. (1974) Atlantic Ocean floor: geochemistry and petrology of basalts from Legs 2 and 3 of the Deep-Sea Drilling Project. J. Geophys. Res., v. 79, p. 5507-5527.
- Fujii, T., Kushiro, I. and Hamuro, K. (1977) Melting relations and viscosity of an abyssal olivine tholeiite (this volume).
- Kirkpatrick, R. J. (1976) Towards a kinetic model for the crystallization of magma bodies. J. Geophys. Res., v. 81, p. 2565-2571.
- Rhodes, J. M., Blanchard, D. P., Rogers, K. V. and Brannon, J. C. (1977) The chemistry of basalts from site 395-Leg 45 of the Deep Sea Drilling Project (this volume).

Roeder, P. L. and Emslie, R. F. (1970) Olivine-liquid equilibrium. Contrib. Mineral Petrol., v. 29, p. 275-289.

Segurdsson, H. and Schilling, J. G. (1976) Spinels in mid-Atlantic ridge basalts: chemistry and occurrence. Earth and Planet. Sci. Lett., v. 29, p. 720.

Aphyric Basalts

<u>Unit</u>	<u>Sample</u>	<u>Anorthite-Content</u>	<u>Petrography</u>
A <sub>2</sub>	395-11-2-100-105	59-61	Groundmass needles
A <sub>2</sub>	295-16-3-3-8		
A <sub>3</sub>	395A-41-1-107-111	86-79	Rounded xenocryst in spherulitic groundmass, irregular zoning pattern
	395A-41-1-107-111	64-62	Groundmass needles
A <sub>3</sub>	395A-56-3-62-65	83-79	Rounded xenocryst in glass rind with spherulitic overgrowth
	395A-56-3-62-65	65-63	Groundmass needles and in glass
A <sub>4</sub>	395A-64-CC	68-43	Core to rim zoning in laths in intersertal-intergranular basalt
A <sub>4</sub>	395A-67-CC	66-64	Groundmass needles

Porphyritic Basalts

<u>Unit</u>	<u>Sample</u>	<u>Anorthite Content</u>	<u>Petrography</u>
P <sub>2</sub>	395A-15-5-16-20 (Coarse-grained groundmass)	86-83 81-78 73-71 73-24(Or=0.6-1.1)	Core of large equant phenocrysts (5 grains) Core compositions in phenocryst (1 grain) Most calcic composition of rims on large phenocrysts Range of rim compositions on large phenocrysts
P <sub>3</sub>	395A-17-1-46-55	84-83 71-62	Core of large equant phenocryst Most calcic composition and most sodic compositions from interior to edge of rim on the same phenocryst. Normal zoning with very weak oscillatory zoning
P <sub>4</sub>	395A-27-2-111-116  (fine-grained groundmass)	83-85  73-63  85-79  72-63 83 86 73-71	Reverse zoned core of large equant grain intergrown with olivine in glomerocryst Most calcic and sodic compositions (interior to edge) within rim on the same phenocryst Range of compositions within oscillatory zoned, large equant phenocryst Range of rim compositions on same phenocryst Core of oscillatory zoned microphenocryst Rim of core portion of the same grain Interior of the rim of the same grain
P <sub>5</sub>	395A-28-1-116-120	85-77  72-70 84-69 59-60 69-71	Range of compositions within oscillatory zoned core of large equant phenocryst in glass Range of the rim composition on the same grain (in glass) Range of compositions in small laths of plagioclase (in glass) Core of exotic, partly resorbed grain Rim overgrowth on same grain

Table 2: Microprobe Analyses of Olivine

<u>Unit</u>	<u>Sample</u>	<u>Forsterite Content</u>	<u>Petrography</u>
A <sub>1</sub>	395·11·1·136-138	84.3	Skeletal microphenocryst
A <sub>2</sub>	395·11·2·100-105	83.9-83.8	Core of microphenocryst
		82.1	Rim of same grain
A <sub>2</sub>	395·16·3·3-8	84.1-83.8	Skeletal microphenocrysts
A <sub>3</sub>	395A·41·1·107-111	88.4	Core of large anhedral xenocryst
		84.4	Rim of same grain
A <sub>3</sub>	395A·56·3·62-65	83.7	Microphenocryst in glass
A <sub>4</sub>	395A·60·1·66-70	87.0	Core of euhedral phenocryst containing chromite inclusion
		85.3	Rim of same grain
		83.8	Groundmass
A <sub>4</sub>	395A·64·CC	82.5	Core of granular olivine in intergranular-intersertal basalt
		75.4	Rim of a granular olivine
A <sub>5</sub>	395A·67·CC	85.1-84.0	Skeletal microphenocrysts and groundmass
P <sub>1</sub>	395·18·2·43-46	85.6-85.4	Cores of microphenocrysts (2)
		84.8	Rim of a microphenocryst
		87.0	Core of large phenocryst in glomerophytic aggregate
P <sub>2</sub>	395A·15·4·10-15	88.0-85.6	Cores of large phenocrysts (3)
		81.9-81.3	Cores of microphenocrysts (3)
		75 -58	Rims of microphenocrysts (3)
P <sub>2</sub>	395A·15·5·16-20	87.7-87.5	Cores of large phenocrysts (2), one of which contains a spinel inclusion
		85.7	Core of large euhedral phenocryst
		83.4-81.2	Cores of small grains (3)
		77 -61	Rims of groundmass grains
P <sub>3</sub>	395A·17·1·46-55	87.8	Core of large phenocryst
		85.5-85.0	Cores and rims of microphenocrysts (3 in glass)
P <sub>3</sub>	395A·19·1·132-136	89.2	Core of large anhedral phenocryst
		87.5	Core of large euhedral phenocryst
P <sub>4</sub>	395A·26·2·125-129	87.4-85.3	Cores of variable-sized phenocrysts (4)
		84.6	Rim of microphenocryst in glass
P <sub>4</sub>	395A·27·2·111-116	89.0-86.3	Cores of large phenocrysts anhedral and euhedral (5)
		84.8-84.2	Cores of microphenocrysts (3)
		82.5	Rim of a microphenocryst
P <sub>5</sub>	395A·28·1·116-122	87.9	Phenocryst cores (2)
		84.8-84.3	Microphenocryst cores and rims in glass (5)



Table 3. Microprobe Analyses of Spinel - Leg 45

	1	2	3	4	5	6	7	8	9	10
TiO <sub>2</sub>	2.1	0.53	0.43	0.92	0.55	0.72	0.41	0.47	0.65	0.40
Al <sub>2</sub> O <sub>3</sub>	22.7	36.5	38.9	33.0	31.0	25.9	32.2	26.2	24.8	26.6
Cr <sub>2</sub> O <sub>3</sub>	31.4	26.8	25.2	28.7	32.6	35.1	32.9	39.0	37.6	38.3
Fe <sub>2</sub> O <sub>3</sub> *	12.6	6.4	5.3	6.3	6.7	8.9	5.6	5.5	7.8	—
FeO	19.6	11.8	11.3	13.3	12.5	14.3	11.6	12.9	14.3	16.9
MnO	0.31	0.18	0.17	0.24	0.23	0.28	0.18	0.27	0.28	0.24
MgO	11.4	17.0	17.4	15.6	16.0	14.4	16.9	15.2	14.2	15.1
Total	100.1	99.2	98.7	98.1	99.6	99.6	99.4	99.5	99.6	97.5
FeO Total	30.9	17.5	16.1	18.9	18.5	22.3	16.2	17.8	21.4	16.9

## Cations Calculated on the Basis of 32 Oxygens

Ti	0.403	0.093	0.075	0.166	0.099	0.135	0.073	0.086	0.122	0.075
Al	6.929	10.033	10.568	9.358	8.739	7.587	8.968	7.542	7.269	7.726
Cr	6.427	4.940	4.591	5.458	6.163	6.895	6.144	7.528	7.390	7.452
Fe <sup>+3</sup>	2.359	1.103	0.909	1.115	1.178	1.622	0.975	0.994	1.432	—
Fe <sup>+2</sup>	4.086	2.251	2.150	2.621	2.454	2.894	2.177	2.585	2.918	3.486
Mn	0.068	0.036	0.033	0.049	0.047	0.059	0.036	0.056	0.059	0.051
Mg	4.399	5.907	4.976	5.592	5.702	5.314	5.950	5.531	5.261	5.547
Total	24.671	24.421	24.346	24.426	24.450	24.624	24.371	24.379	24.549	24.336
Cr/(Cr+Al)	0.481	0.330	0.303	0.368	0.414	0.476	0.407	0.500	0.504	0.491
Mg/(Mg+Fe)	0.397	0.634	0.658	0.595	0.606	0.534	0.650	0.603	0.542	0.641

Fe<sub>2</sub>O<sub>3</sub>\* calculated on the assumption of stoichiometry.

- 395A.13.1.142-147 Unusual web-shaped grain which occurs in a large crystalline melt inclusion within a plagioclase phenocrysts. (Unit P<sub>1</sub>)
- 395.15.4.10-15 Core and opposite rims of a spinel grain included within plagioclase. #3 is the composition of the rim facing the interior of the plagioclase, #4 is located at the rim of the plagioclase. (Unit P<sub>1</sub>)
- 395A.27.2-111-116 Core and rim of a euhedral octahedron of spinel included within a fine-grained ground mass. (Unit P<sub>3</sub>)
- 395A.27.2.111-116 Core of subhedral spinel included along a plagioclase-olivine grain boundary in a glomerocrystic aggregate. (Unit P<sub>3</sub>)
- 395A.21.1.87-90 Core and rim of a euhedral grain included within fine-grained ground mass. (Unit P<sub>2</sub>)
- Experimentally produced spinel in 1200°C melting run using powder of 395A.27.2.111-116. p = 1 atmosphere, f<sub>O<sub>2</sub></sub> = 10<sup>-9</sup>.

Table 4. Microprobe Analyses of Clinopyroxenes Leg 45

	1	2	3	4	5	6	7	8	9	10
SiO <sub>2</sub>	49.7	50.4	51.5	50.9	51.4	50.6	51.0	50.7	52.3	50.4
TiO <sub>2</sub>	0.82	0.82	0.65	0.47	0.78	0.34	0.34	1.28	0.66	0.88
Al <sub>2</sub> O <sub>3</sub>	4.6	3.8	1.6	4.2	3.7	4.1	3.7	3.3	1.6	1.5
Cr <sub>2</sub> O <sub>3</sub>	1.2	0.51	0.03	1.5	0.19	1.4	1.15	0.05	0.02	0.01
FeO	5.7	6.5	9.4	4.8	7.6	4.2	4.3	10.3	8.9	15.5
MnO	0.18	0.18	0.25	0.16	0.22	0.14	0.10	0.26	0.30	0.47
MgO	15.7	16.3	17.1	17.0	18.0	17.3	17.5	15.3	17.4	14.0
CaO	20.9	20.2	18.1	21.0	17.8	20.9	21.2	19.4	17.9	16.0
Na <sub>2</sub> O	0.33	0.34	0.23	0.23	0.17	0.28	0.27	0.35	0.25	0.29
Total	99.13	99.05	98.86	100.26	99.86	99.26	99.56	100.94	99.33	99.05
Si	1.850	1.876	1.932	1.863	1.888	1.866	1.875	1.878	1.943	1.933
Al	0.150	0.124	0.068	0.137	0.112	0.134	0.125	0.122	0.057	0.067
Al	0.053	0.043	0.001	0.045	0.047	0.044	0.035	0.022	0.013	0.001
Ti	0.023	0.023	0.018	0.013	0.022	0.009	0.009	0.036	0.018	0.025
Cr	0.036	0.015	0.001	0.043	0.006	0.040	0.033	0.001	0.001	0.000
Fe	0.178	0.201	0.296	0.146	0.235	0.130	0.132	0.319	0.277	0.497
Mn	0.006	0.006	0.008	0.005	0.007	0.004	0.003	0.008	0.009	0.015
Mg	0.868	0.904	0.954	0.929	0.987	0.951	0.959	0.845	0.964	0.800
Ca	0.832	0.805	0.727	0.822	0.699	0.826	0.835	0.770	0.713	0.657
Na	0.024	0.025	0.017	0.016	0.012	0.020	0.019	0.025	0.018	0.022
Total	4.020	4.022	4.023	4.019	4.014	4.025	4.028	4.026	4.012	4.018
Wo	44.3	42.1	36.7	43.3	36.4	43.3	43.4	39.8	36.5	33.6
En	46.2	47.3	48.3	49.0	51.4	49.9	49.8	43.7	49.3	40.9
Fs	9.5	10.5	15.0	7.7	12.2	6.8	6.9	16.5	14.2	25.4

1,2. 395A.15.4.10.15 Core and rim of euhedral, zoned phenocryst.

3. 395A.15.4.10.15 Rim of groundmass grain in contact with mesostasis.

4. 395A.19.1.132-136 Core of zoned phenocryst.

5. 395A.19.1.132-136 Rim of groundmass grain.

6,7,8. 395A.15.4.99-107 Core, intermediate zone and rim of a large clinopyroxene glomerocryst intergrown with plagioclase laths.

9,10. 395A.15.5.16-20 Core and rim of groundmass grain.

Table 5: Compositions of olivine and plagioclase from experimental runs

Temperature (°C)	Sample 395A·27·2·111-116				Sample 395A·9·2·100-104			
	Fo content of olivine		An content of plagioclase		Fo content of olivine		An content of plagioclase	
	Range	Mean	Range	Mean	Range	Mean	Range	Mean
1225			74.8-72.1	73.7 (4)				
1200	86.1-84.7	85.1 (10)	75.4-72.4	74.2 (4)	83.2-82.9	83.1 (7)		
1175	84.6-83.9	84.2 (5)	69.8-68.4	69.3 (3)	81.5-79.0	80.0 (5)	63.4-61.5	62.6 (4)
1160	29.1-76.3	77.4 (5)	71.0-65.4	68.2 (2)	77.9-74.5	76.3 (8)	63.2-58.5	61.5 (4)

Figure 1. Microprobe analyses of spinels in porphyritic basalts (Hole 395A). The presence of melt inclusions in many spinel grains indicates that they are products of basalt crystallization. Many of the grains are included within olivine and plagioclase and the compositions reflect, in part, reequilibration with host grain and crystallization from trapped liquid. Zoning in groundmass grains shows two trends. One exhibits simple CrFe-enrichment while the other shows FeTi enrichment at relatively constant Cr/Al.

Figure 2. Pyroxenes from five porphyritic basalts; Holes 395 and 395A.

Figure 3. Microprobe analyses of clinopyroxene phenocrysts and groundmass pyroxene in five porphyritic basalts showing strong chromium zoning with iron enrichment. Zoning within large phenocrysts shows a very slight iron increase and a decrease in  $Cr_2O_3$ . Chrome-contents of the cores of some grains in subophitic groundmass are  $>1\% Cr_2O_3$  indicating that the drastic chromium depletion in groundmass rims did not occur until some of the groundmass grains had formed. Note crude correlation between bulk rock composition and pyroxene composition.

Figure 4. Microprobe analyses of clinopyroxene phenocrysts and groundmass grains in five porphyritic basalts. The trends demonstrate decreasing alumina and chromium in phenocrysts followed by an increase in  $TiO_2$  in rims and groundmass grains. Note crude correlation between  $TiO_2$ -content of bulk rock and the trend on the Ternary plot.

Figure 5. Graphical presentation of one-atmosphere melting experiment results of two basalts (aphyric and phyrlic). Hole 395A.

## Plate Captions

### Plate 1

- a) Sample 395·11·1·136-138: Olivine microphenocrysts in the rapidly quenched groundmass at pillow margin in unit A . Width across photo = 1.25 mm.
- b) Sample 395·15·2·68-72: (Same scale, Unit A ). A more slowly cooled variant from the same unit that contains equant microphenocrysts in a fitted groundmass of poorly developed fan spherulites of plagioclase needles with associated skeletal groundmass olivine and very fine-grained interstitial plumbose pyroxene.
- c) Sample 395·16·3·3-8: Another still more slowly cooled variant from the same unit (note change of scale-width of photograph = 2.5 mm). Plagioclase needles are pyroxene and interstitial skeletal clinopyroxene are more coarse-grained. Fan spherulites generally radiate from olivine microphenocrysts.
- d) Sample 395·11·2·100-105: (same scale as c) Most coarse-grained textural variant. The texture is intergranular-intersertal and at these slower cooling rates the mineralogy is dominated by skeletal plagioclase laths.
- e) Sample 395A·9·2·100-104: Product of linear cooling rate of 2°C/hr. Skeletal olivine microphenocrysts in a fine-grained groundmass similar to that shown in 1a.
- f) Sample 395A·46·1·61-64: Typical intersertal texture of aphyric basalts in unit A<sub>3</sub> (width across photograph = 2.5 mm).
- g) Sample 395A·56·3·62-65: (same scale as f) Rounded plagioclase xenocryst in glassy pillow rind. Note embayed outline suggestive of resorption.
- h) Detail of the margin of the xenocryst (left) against glass. The spherulitic overgrowth on the xenocryst and single microlite of plagioclase adjacent to xenocryst appears as a black band. Note also olivine microphenocrysts in glass (Photograph width=1.25 mm).

## Plate 2

- a) Sample 395A·15·4·10-15: Typical porphyritic basalt with subophitic groundmass in unit P<sub>2</sub>. Note large plagioclase glomerocrysts consisting of large weakly zoned cores and abrupt rims (width of photograph = 1.1 cm).
- b) Sample 395·18·2·43-46: Plagioclase phenocryst exhibiting complex oscillatory zoning and containing numerous melt inclusions in the core. The majority of phenocrysts in these rocks do not exhibit this type of zoning pattern (width of photograph = 3.5 mm).
- c) Sample 395A·38·1·116-122: An example of the least common distinctive zoning pattern type in the porphyritic basalts. The core is An<sub>59-60</sub> and the rim (darker) is more calcic at An<sub>73-71</sub>. Note vermicular zone in between core and rim indicative of resorption (width of photograph = 2.5 mm).
- d) Sample 395A·27·2·111-116: A complex olivine-plagioclase-spinel glomerocrystic aggregate is illustrated. A single large olivine with a highly irregular outline and numerous melt inclusions is intergrown with plagioclase phenocrysts that exhibit variable zoning patterns (width of photograph = 3.5 mm).
- e) Sample 395A·27·2·111-116: A large euhedral olivine phenocryst contains melt inclusions and two anhedral plagioclase grains (width of photograph = 2.5 mm).
- f) Sample 395A·15·4·10-15: Typical zoning pattern in clinopyroxene phenocryst. (Width of photograph = 1.5 mm).