UNIVERSITY OF UTAH RESEARCH INSTITUTE



EARTH SCIENCE LABORATORY 391 CHIPETA WAY, SUITE C SALT LAKE CITY, UTAH 84108—1295 TELEPHONE 801-524-3422

October 18, 1988

Dr. Daniel Carrier Unocal Geothermal Division Unocal Corporation 3576 Unocal Place Santa Rosa CA 95406

Dear Dr. Carrier:

Attached are the preliminary results of x-ray diffraction and petrographic analyses for 10 drill cutting samples which you submitted in early August.

These samples of the "8"-series group are predominately andesites or porphyritic andesites with the exception of samples 8-10 and 8-18 which contain principally quartz latite or rhyodacite.

Sample 8-18 consists almost entirely of quartz, plagioclase and potassium feldspar. Sample 8-10 contains more primary quartz with some phenocrysts exhibiting dissolution phenomena. Granophyric, axiolitic and spherulitic textures are well developed in sample 8-10 and in some chips of quartz latite in sample 8-12.

Principal hydrothermal alteration products in these rocks are potassium feldspar, quartz, actinolite, chlorite (and mixed-layer chlorite-smectite), epidote, and leucoxene. Calcite, prehnite, wairakite and phengite are important subordinate vein-forming minerals. In the more porphyritic andesites, actinolite appears to replace primary clinopyroxene phenocrysts, especially in samples 8-13 and 8-16. (However, clinopyroxene in sample 8-11 appears relatively unaltered.) Actinolite also occurs as long prisms within quartz veins in sample 8-13. In sample 8-16, actinolite occurs in a large vein with epidote and chlorite and as a replacement of clinopyroxene phenocrysts.

Potassium feldspar veining and flooding is common in sample 8-13. In samples 8-9 and 8-11, some of the primary plagioclase is

replaced by carbonate and some by sericite. Epidote (with or without leucoxene) also commonly replaces plagioclase phenocrysts.

The quartz latite samples (8-10 and 8-18) contain mostly primary potassium feldspar and quartz as devitrification aggregates. Epidote and chlorite appear as the principal alteration products in these quartz latites or rhyodacittes.

Unfortunately, a thin section of sample 8-17 was not made so it is not possible to determine from the diffraction patterns alone whether the amphibole is primary hornblende or secondary actinolite, and whether the 10 angstrom mica is primary biotite or secondary illite and/or phengite. Another problem is the presence of an unidentified 9.40 angstrom peak in diffractograms of samples 8-11, 8-12 and 8-13 which may represent a talc-group mineral not recognized petrographically (although traces of serpentinite are evident).

As I mentioned in our phone conversation, a more detailed geologic and mineralogic report will be prepared after Jeff returns in a few weeks time. I've sent the "17"-series group out to be made into thin sections and hope to look at them fairly soon.

Thank you for this great introduction to hydrothermal alteration mineralogy. Please call me if you have any questions or if you would prefer only XRD data sent initially.

Sincerely,

Shoan lutz

Susan Lutz Manager, X-ray Diffraction Laboratory

SJL:kr

UNOCAL GEOTHERMAL · DANIEL CARRIER MINERALOGY, APPROX. WT.% X (or) RELATIVE ABUNDANCE 10 CUTTINGS SAMPLES 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 1017 | 10 8-9 70 8-18 PRELIM. BULK KRO SAMPLE NO. TRACE - SERPENTINITE 8-9 TR 7 3 4 5 5 20 13 17 43 34 2 8-11 15 17 5 TR TRACE -SERPENTINITE 8 8-12 40 3 TR 16 3 12 21 8-13 13 9 43 1 8-14 2 2 27 5 8-14 20 4 3 10 2 6 1.0 3 8 - 16 ·13 8-17 5 5 4 2 (NO THIN SECTION) 18 56 8-18 14 8 3 COULD BE ACTINOLITE * NOTIS - COULD INCLUDE MM = PREDOMINANT M = MAJOR m = MINORTr = TRACE? = TENTATIVE IDENTIFICATION



SUMMARY OF X-RAY DIFFRACTION ANALYSIS

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SUMMARY OF X-RAY DIFFRACTION ANALYSIS

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UNIVERSITY OF UTAH RESEARCH INSTITUTE, EARTH SCIENCE LABORATORY

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SUMMARY OF X-RAY DIFFRACTION ANALYSIS

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